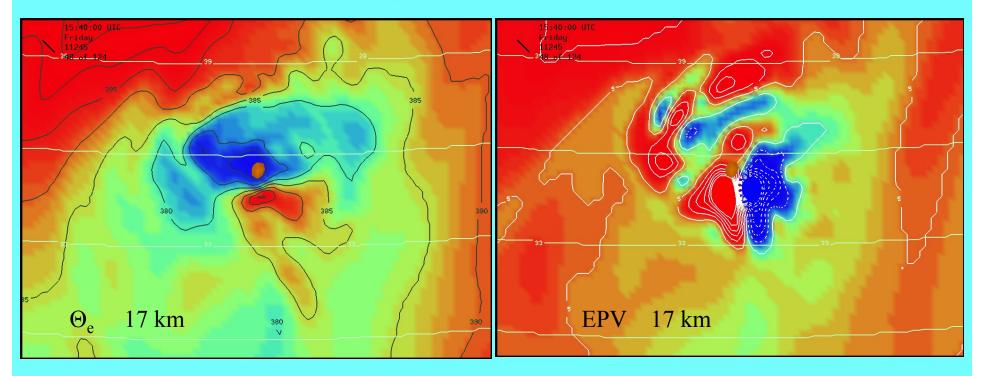
A Modeling Study of STE Near Tropical Cyclones Talas and Ita

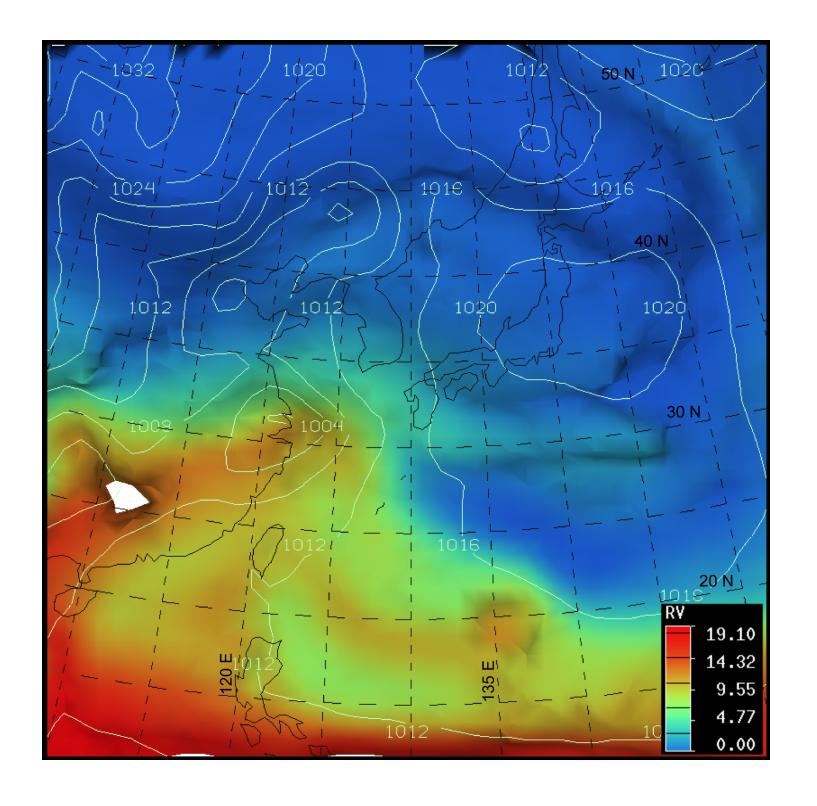


Matthew H. Hitchman and Shellie M. Rowe

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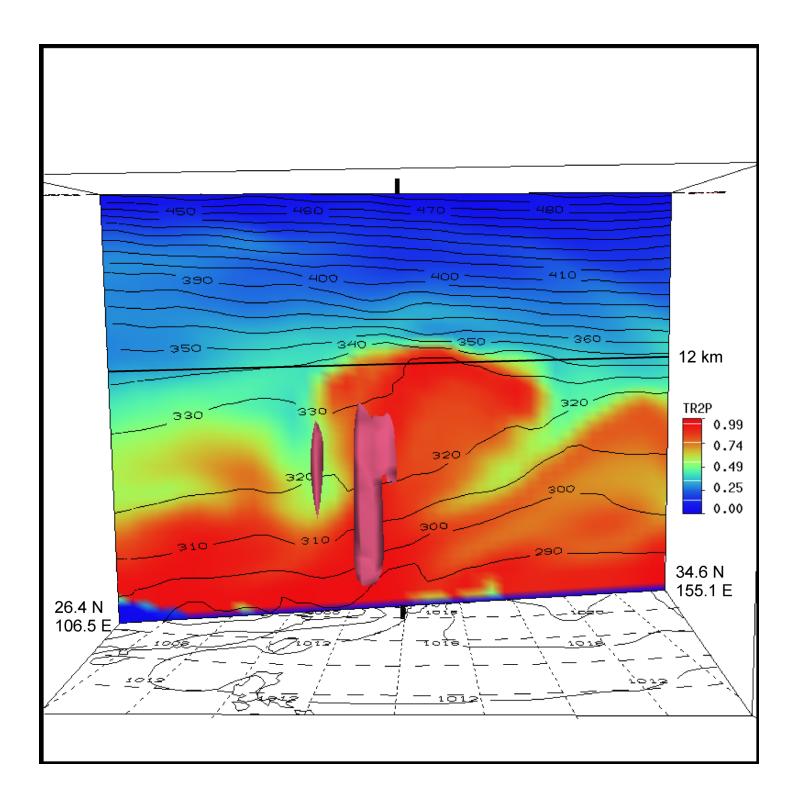
IGAC Composition and Transport in the Tropical Troposphere and Lower Stratosphere
Boulder CO July 20, 2015

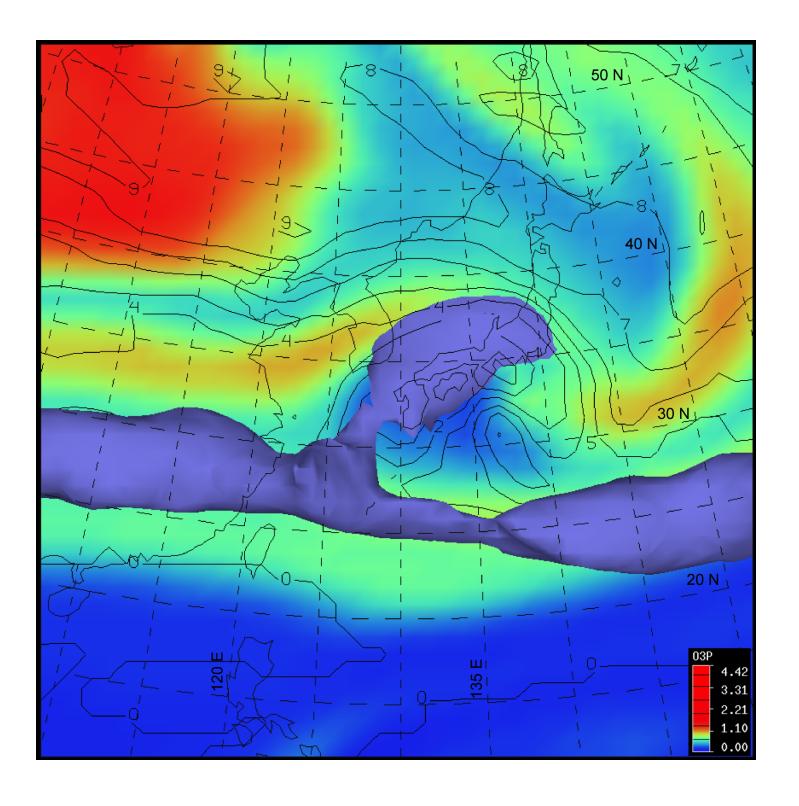
- 1. Example of STE on the periphery of deep convection
- 2. Inertial instability and PV dipoles
- 3. Tropical cyclone Talas, Japan, Sept 2011 CIMSS Satellite loops UWNMS simulation Sept 2-4
- 4. Tropical cyclone Ita, Queensland, April 2014 UWNMS simulation April 10-11

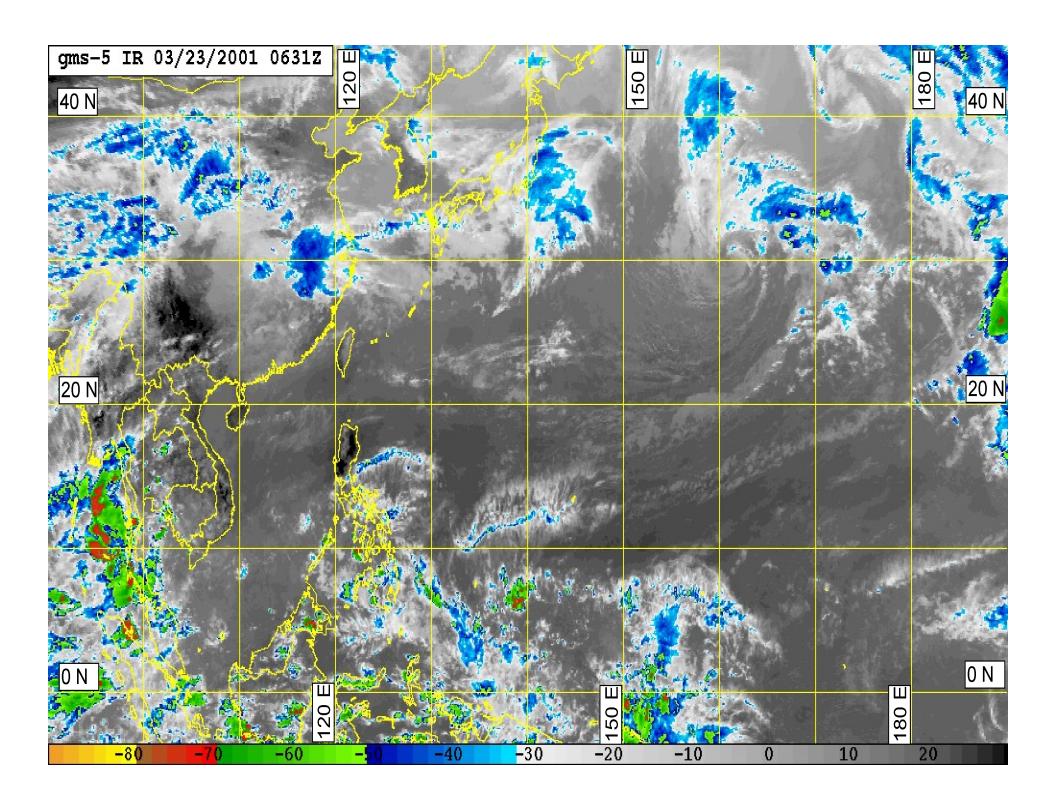


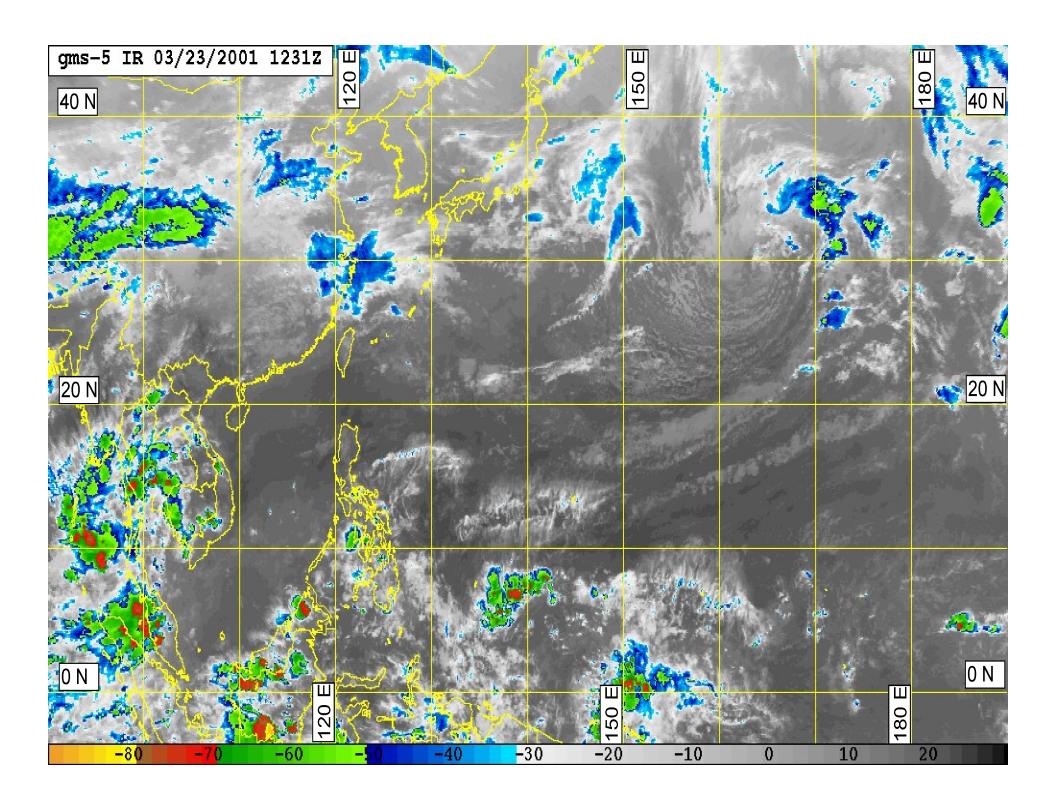
UWNMS forecast for TRACE-P

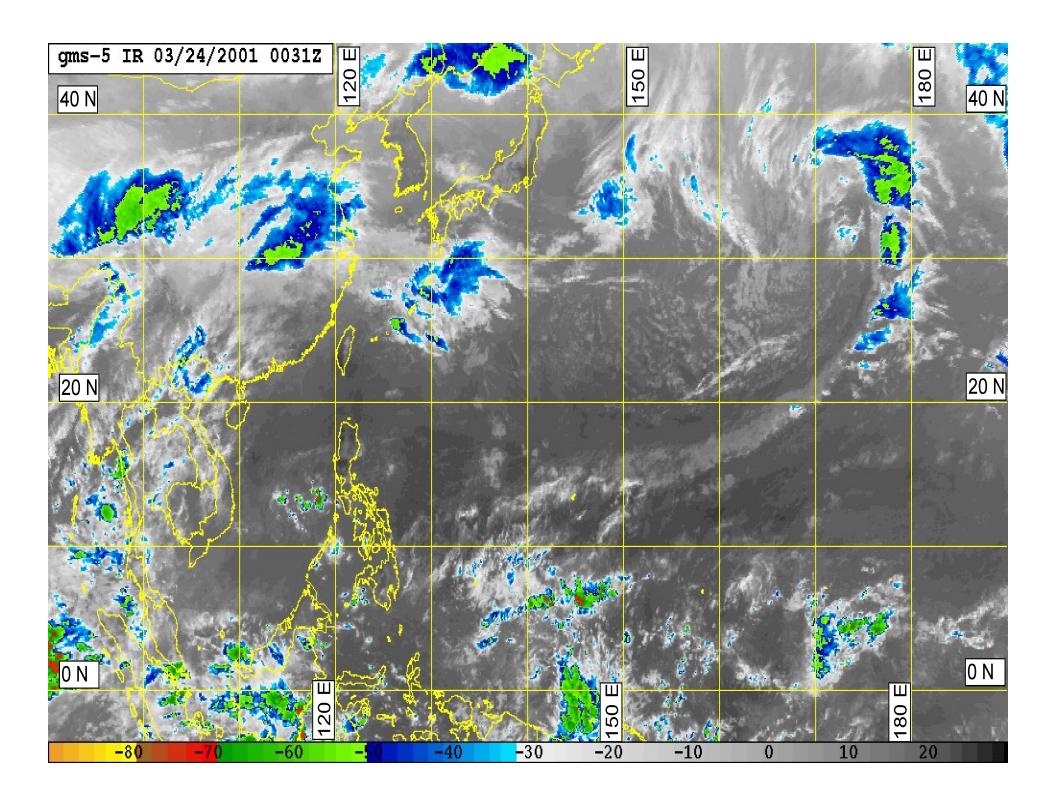
Mar 23, 2001

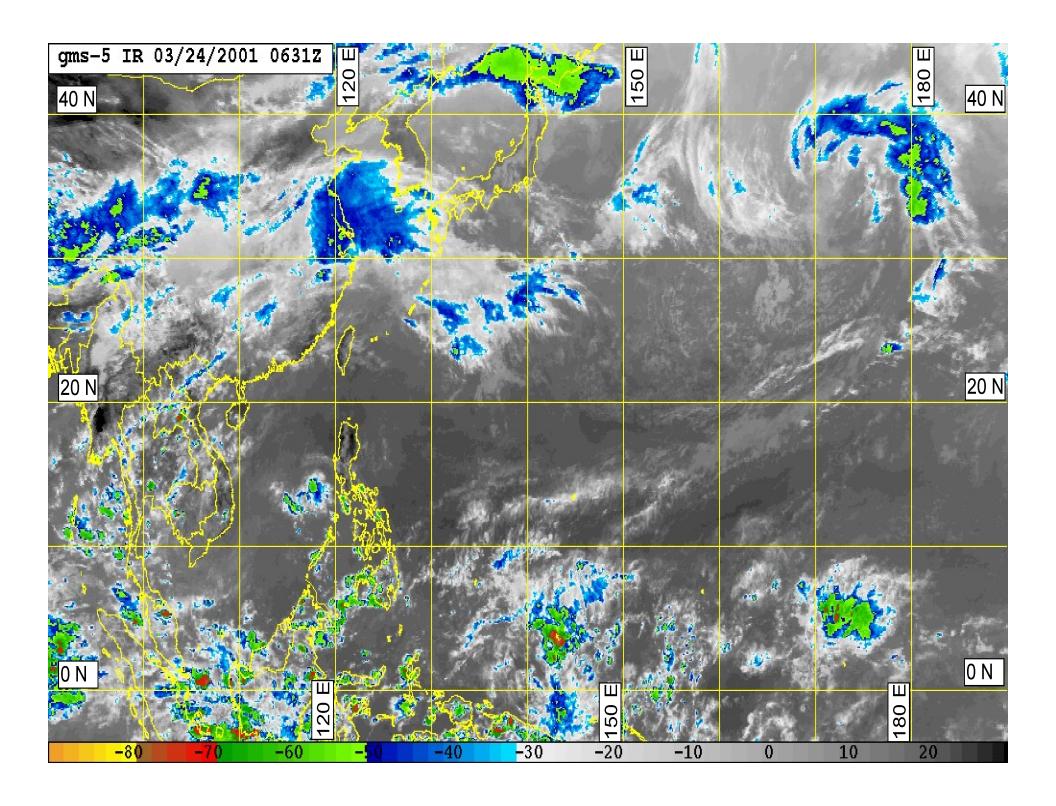


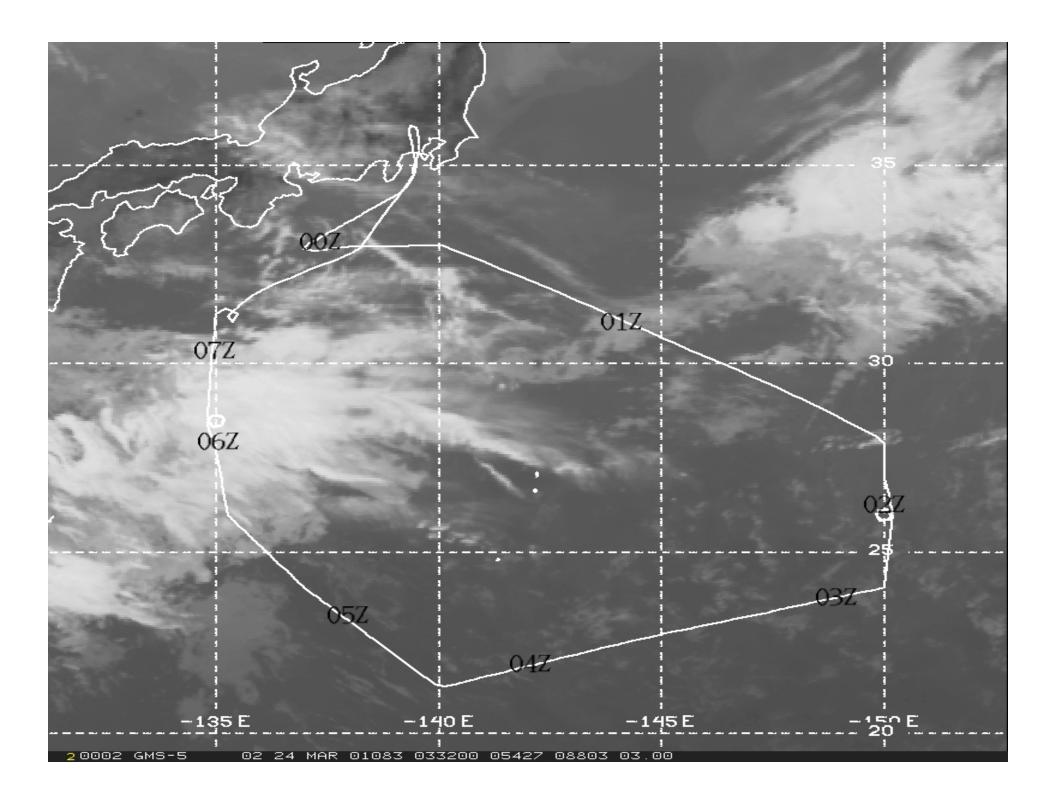


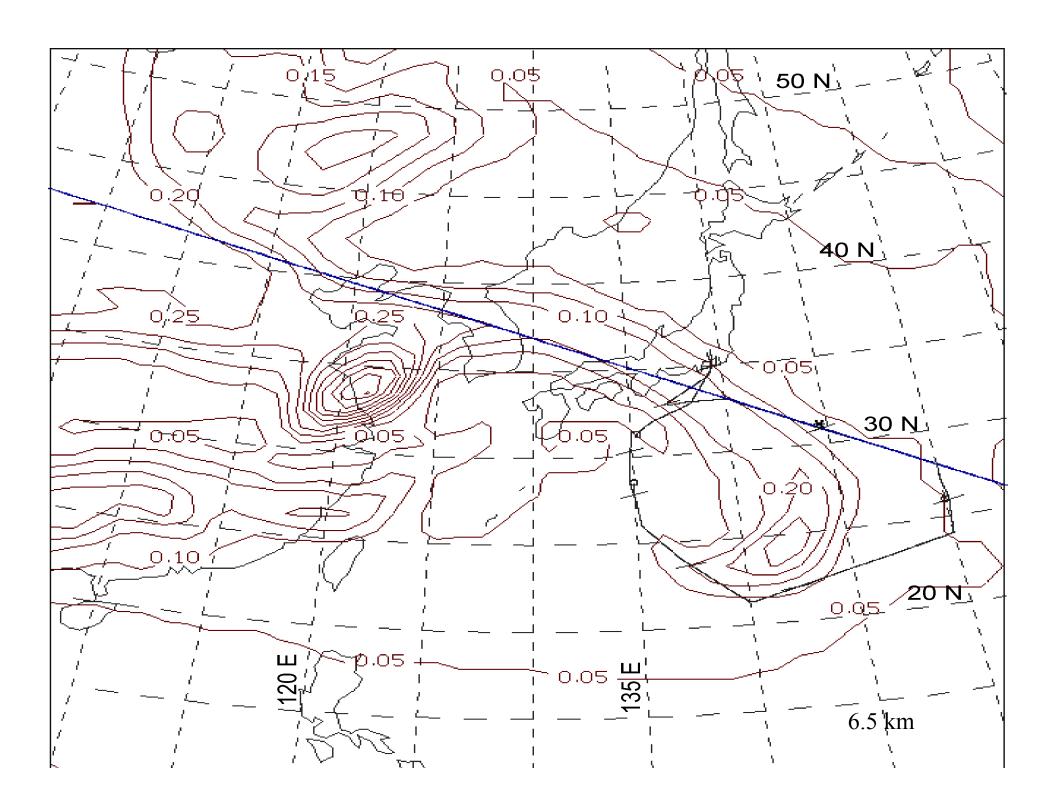


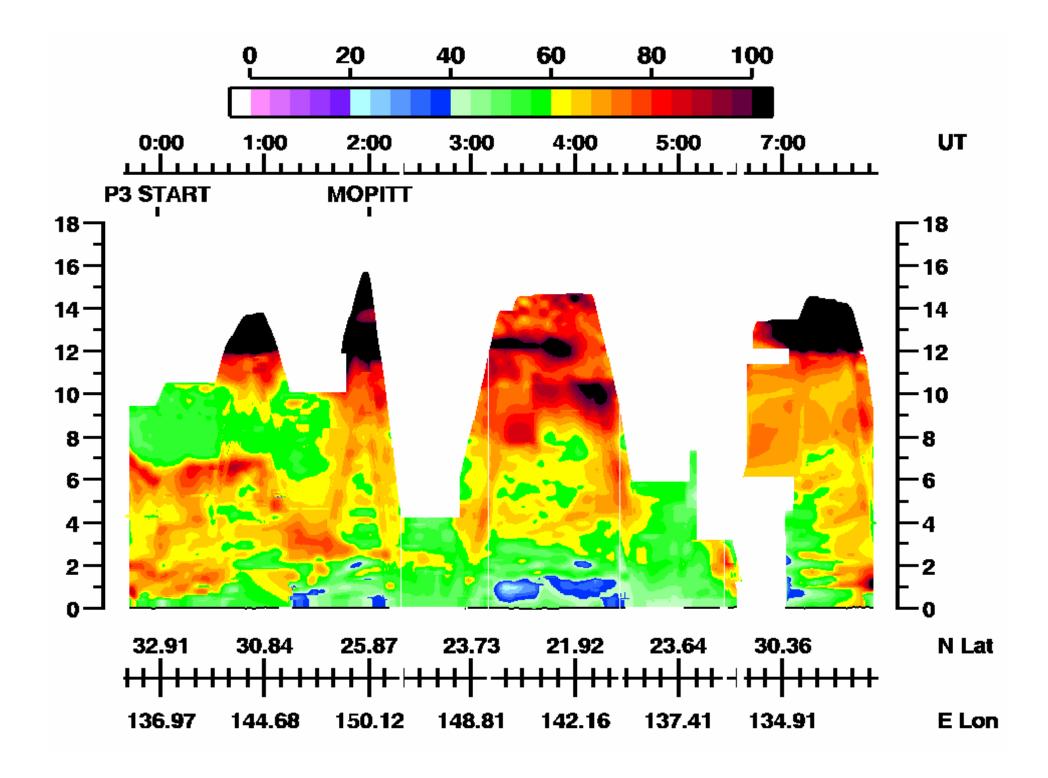


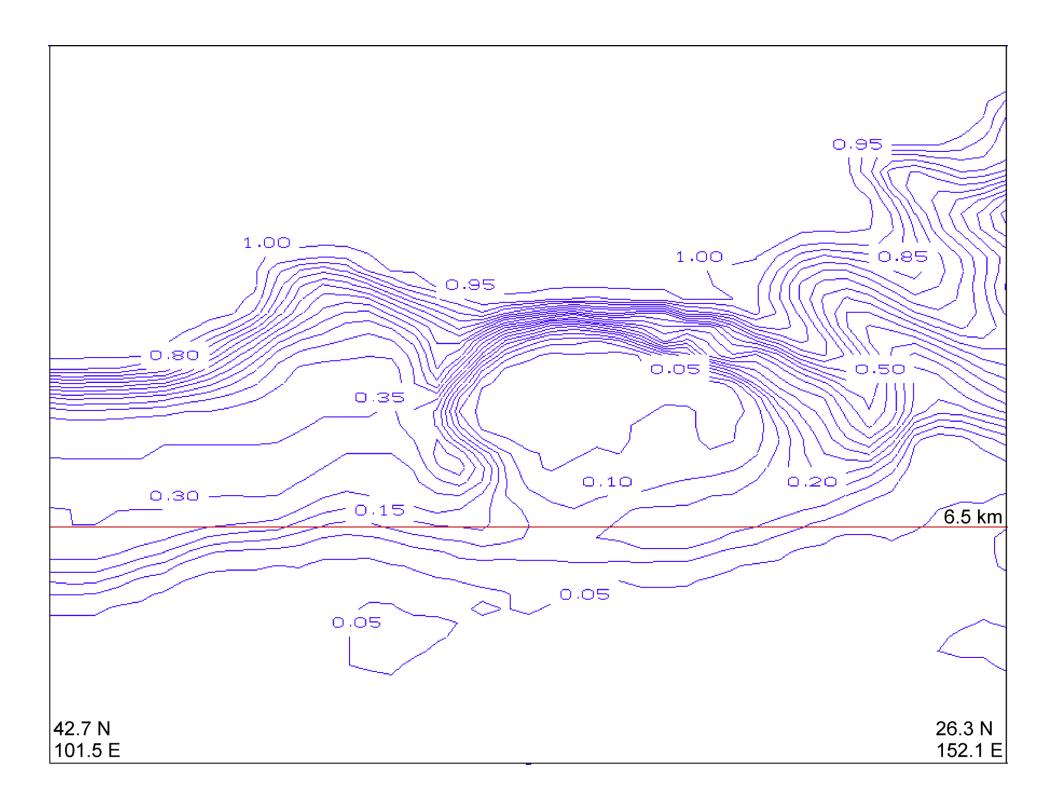


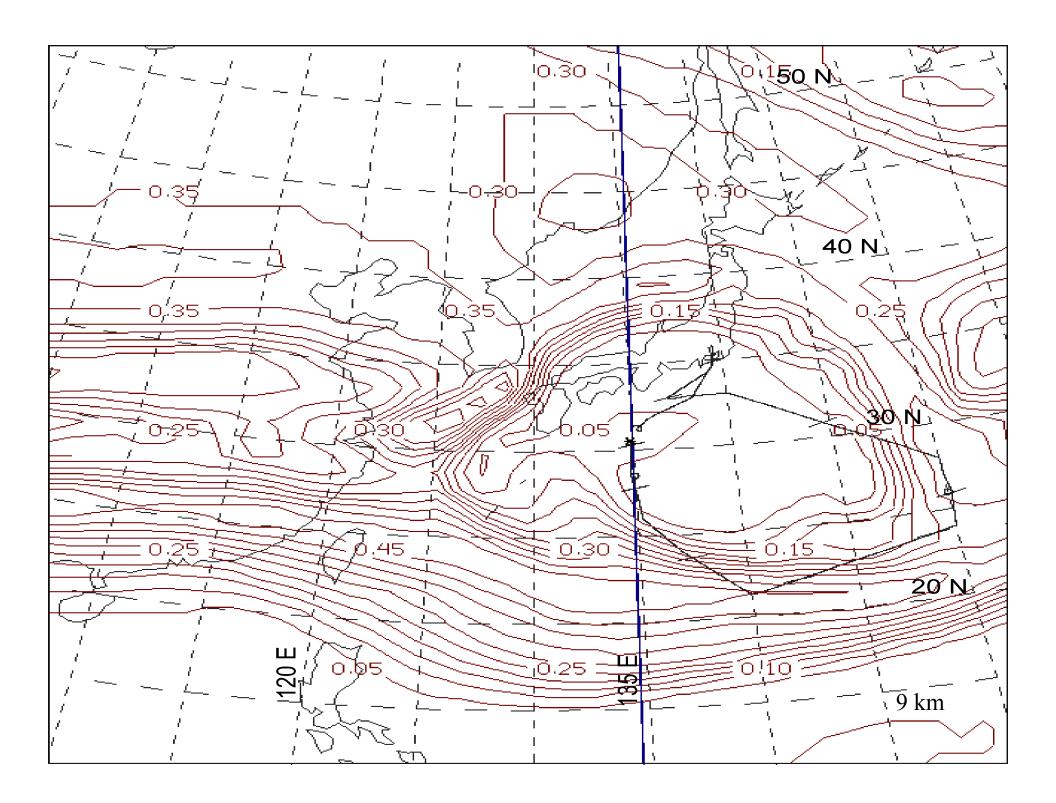


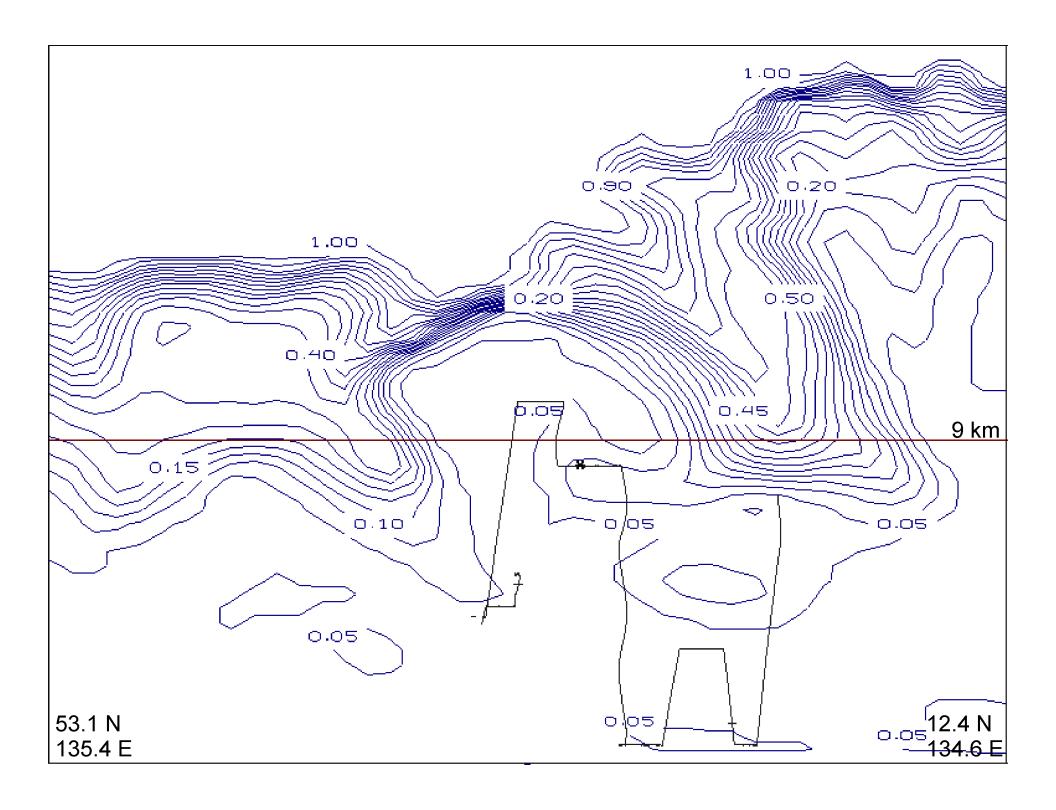




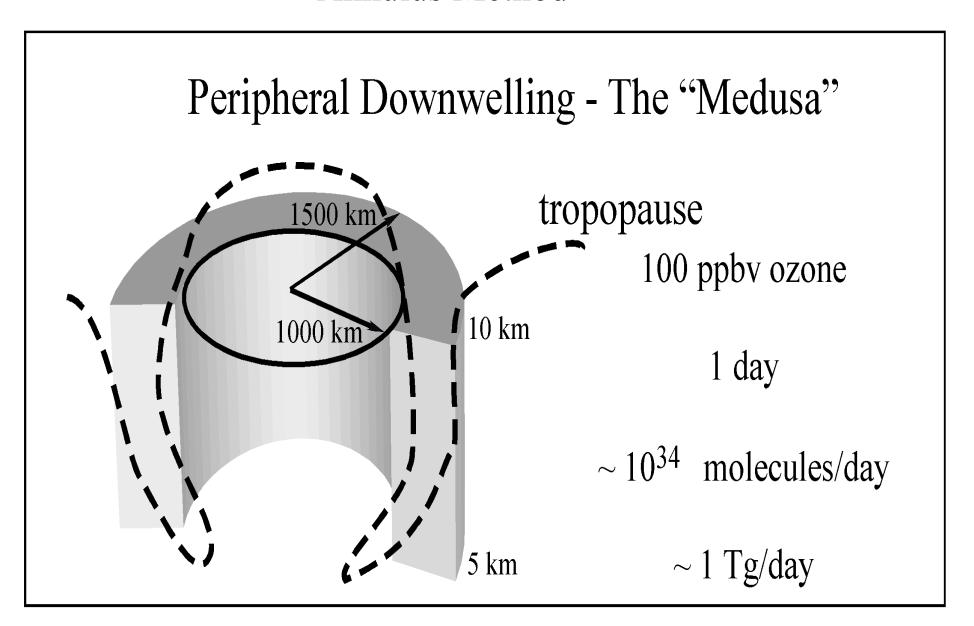






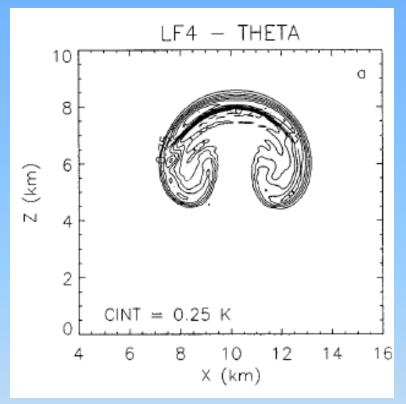


Annulus Method



Hitchman et al. JGR 2004

Wicker and Skamarock 1997 MWR









Inertial Instability

$$\frac{\partial^2}{\partial t^2} \delta s = -f(f + \zeta) \delta s$$
$$f(f + \zeta) < 0$$

$$M = r u + r^2 \Omega$$
, $r = a \cos \varphi$

$$f PV = \frac{1}{\rho} \frac{\partial \theta}{\partial z} f(f + \zeta) = -\frac{1}{r} \nabla_{\theta} M < 0$$

$$EPV = \frac{1}{\rho} \frac{\partial \theta_e}{\partial z} (f + \zeta) < 0 \qquad \theta_e = \theta \ exp \left(\frac{Lw_s}{c_p T} \right)$$

Signature: negative absolute vorticity, PV, or EPV

reversed absolute angular momentum contours

Implies: enhance div/conv, gravity wave excitation,

acceleration of outward and poleward flow

Chagnon and Gray, 2009, QJRMS

$$\frac{\partial q}{\partial t} = \left(f \frac{\partial}{\partial z} - \Lambda \frac{\partial}{\partial x} \right) \left(\frac{B}{N^2} \right)$$
 $\Lambda = dV/dz$ B is the rate of heating

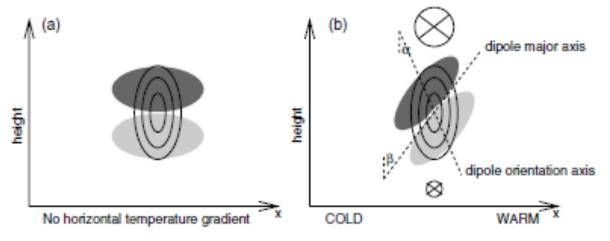
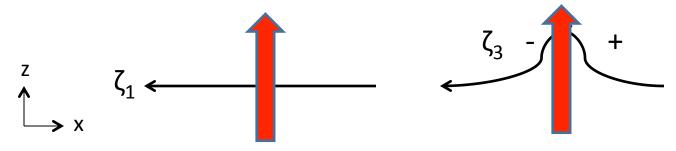


Figure 1. Schematic diagram depicting the orientation of the PV dipole arising from heating applied against (a) a barotropic environment having a background vertical (planetary) vorticity, and (b) a baroclinic environment containing vertical wind shear (directed into the page). Positive (negative) PV anomaly is indicated by light (dark) shading.



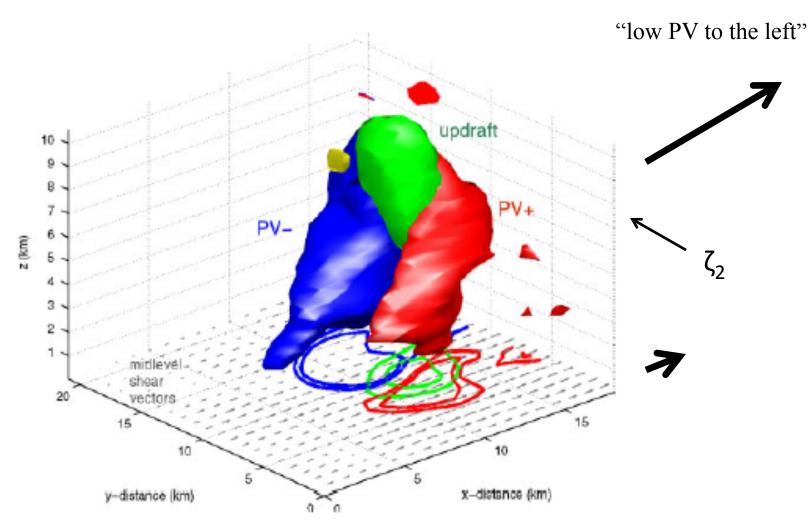
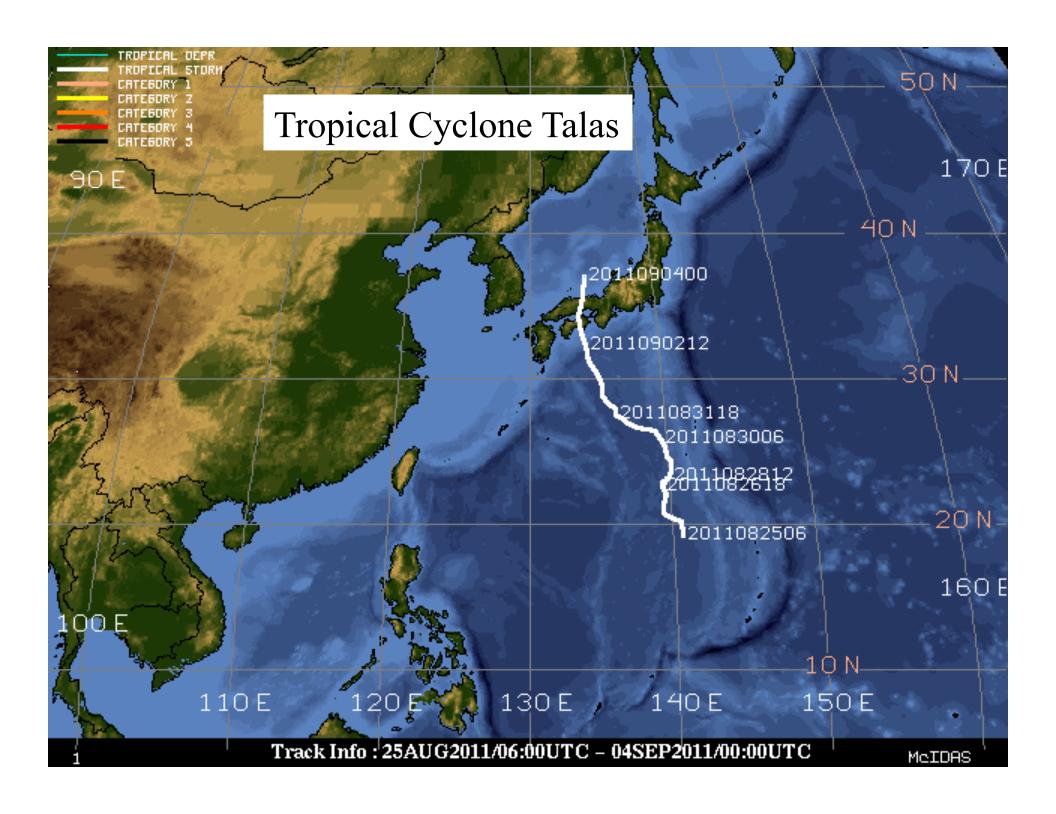
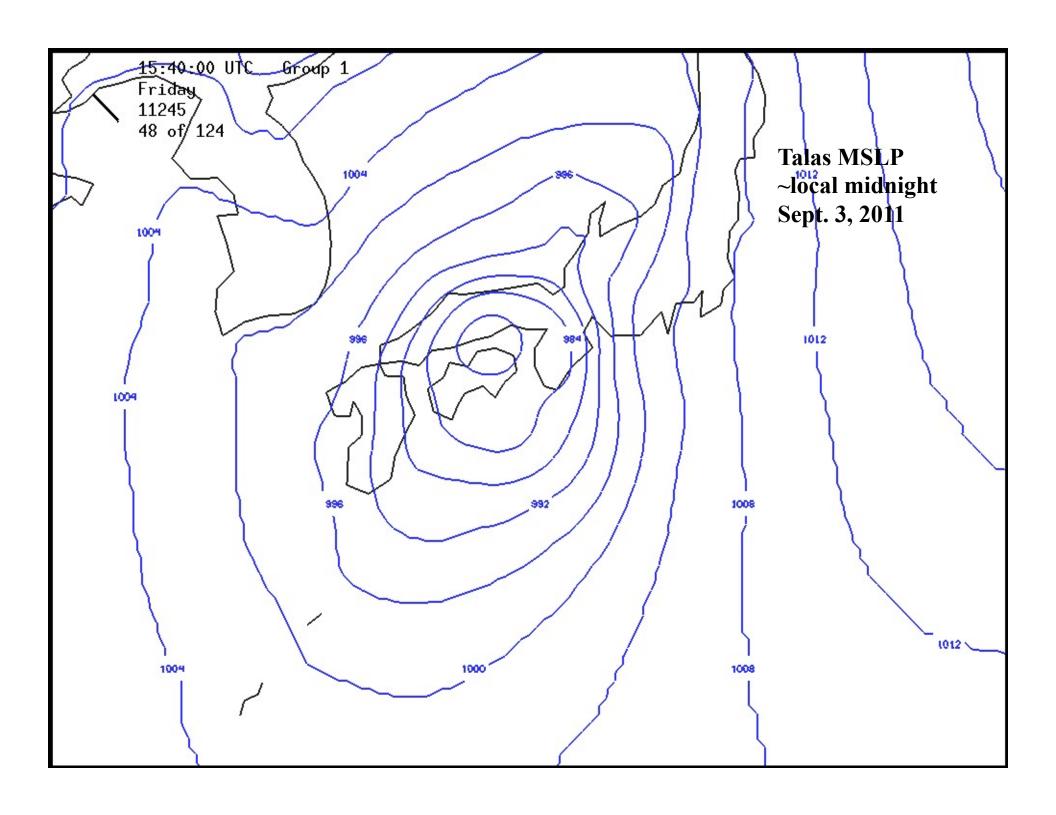
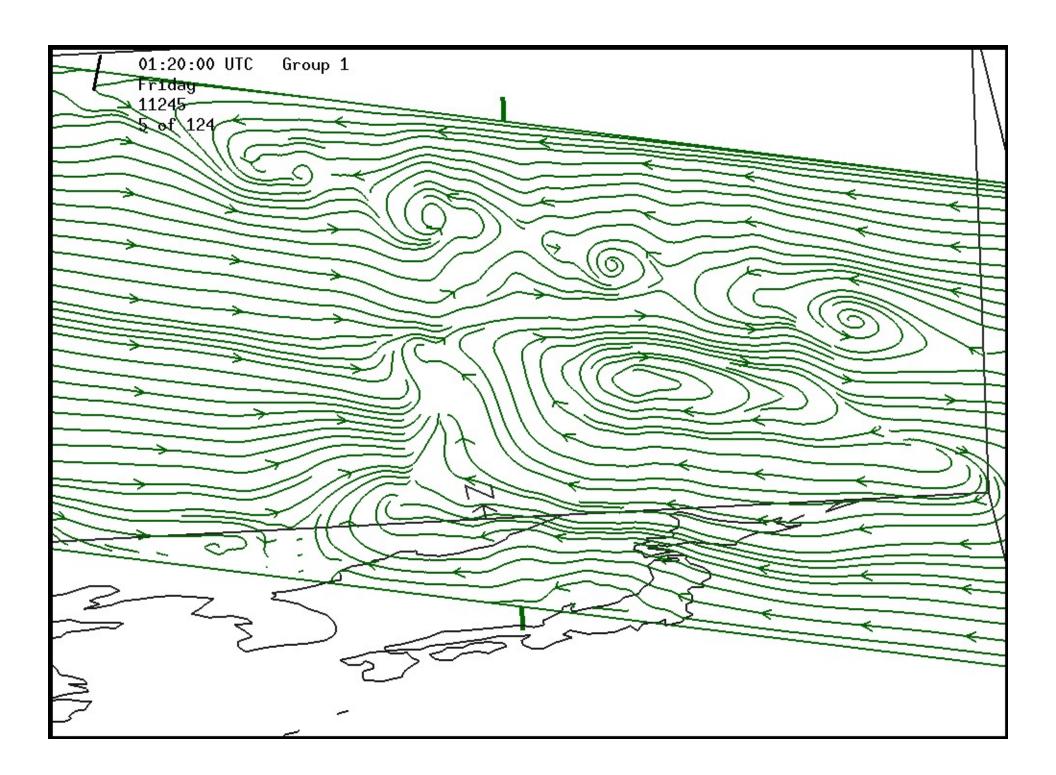


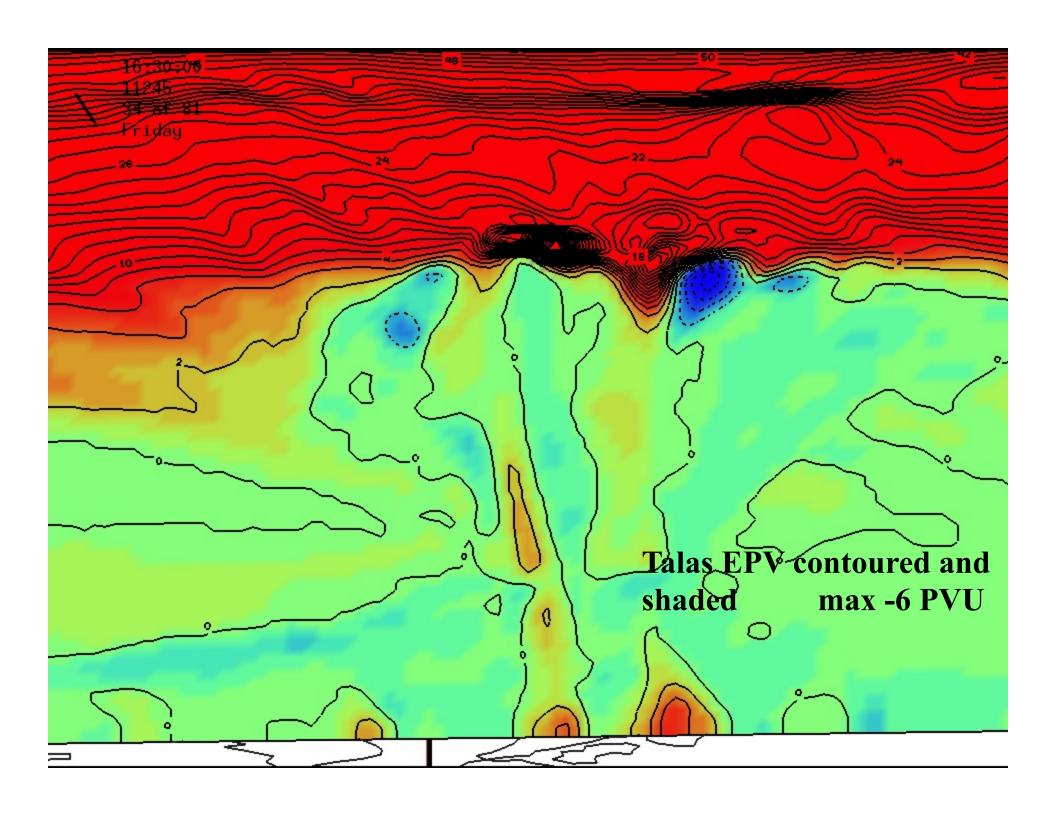
Figure 9. Three-dimensional rendering of the horizontal PV dipole. Plotted are isosurfaces of potential vorticity (in the online figure, the red surface denotes 15 PVU and the blue surface -15 PVU) and vertical velocity (in the online figure, the green surface is for 7 m s⁻¹). For reference, contours of PV (10 and 15 PV) and vertical velocity (5 and 7 m s⁻¹) at an elevation z = 6.5 km are projected on to the lower horizontal plane. Mid-tropospheric shear vectors (averaged from z = 2 to 7 km) are also plotted in the lower horizontal plane. A colour version of this figure is available online. This figure is available in colour online at www.interscience.wiley.com/journal/qj

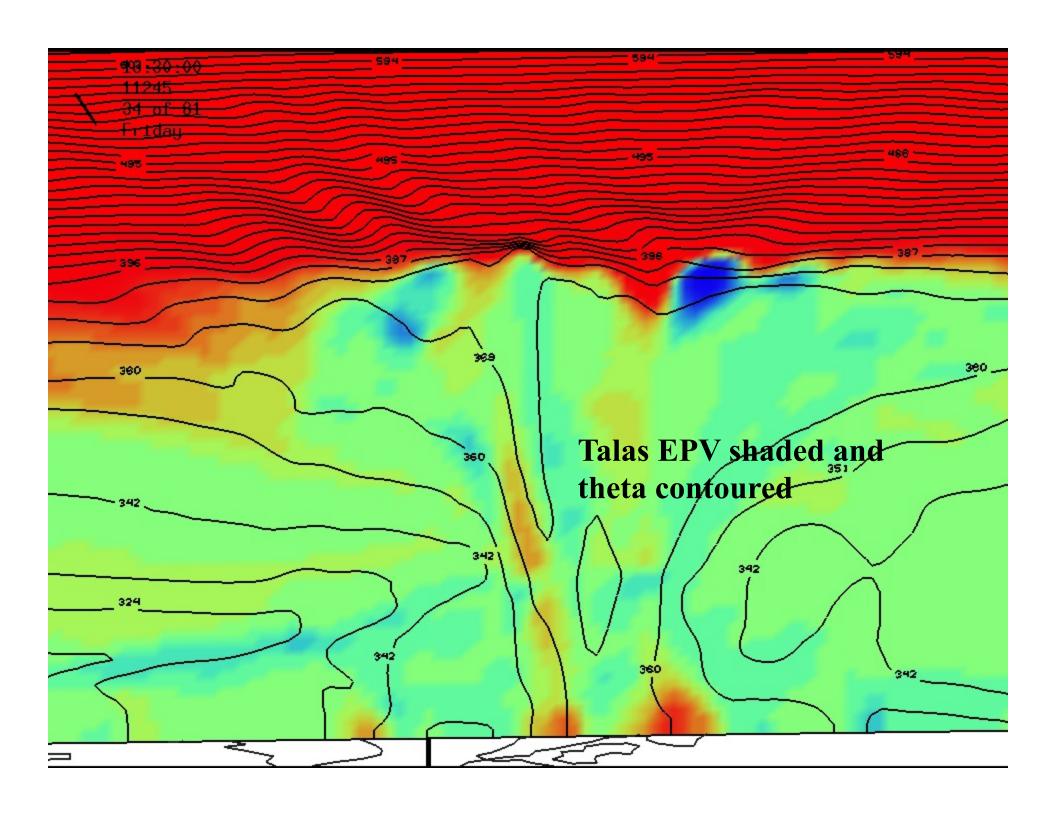


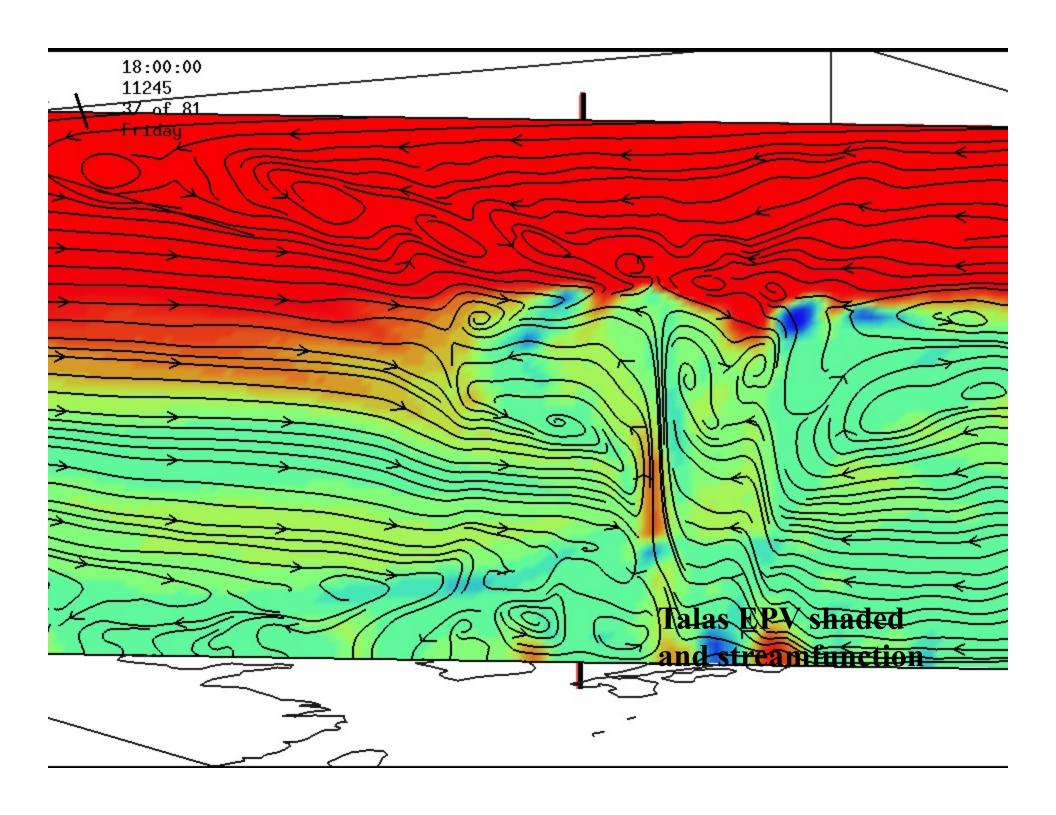
- 1. Enhanced IR film loop
- 2. Water vapor film loop
- 3. Upper level winds loop
- 4. UWNMS simulation model top 26 km vertical resolution ~300 m horizontal resolution 20 km 2.5° ECMWF initialization run 48 hr

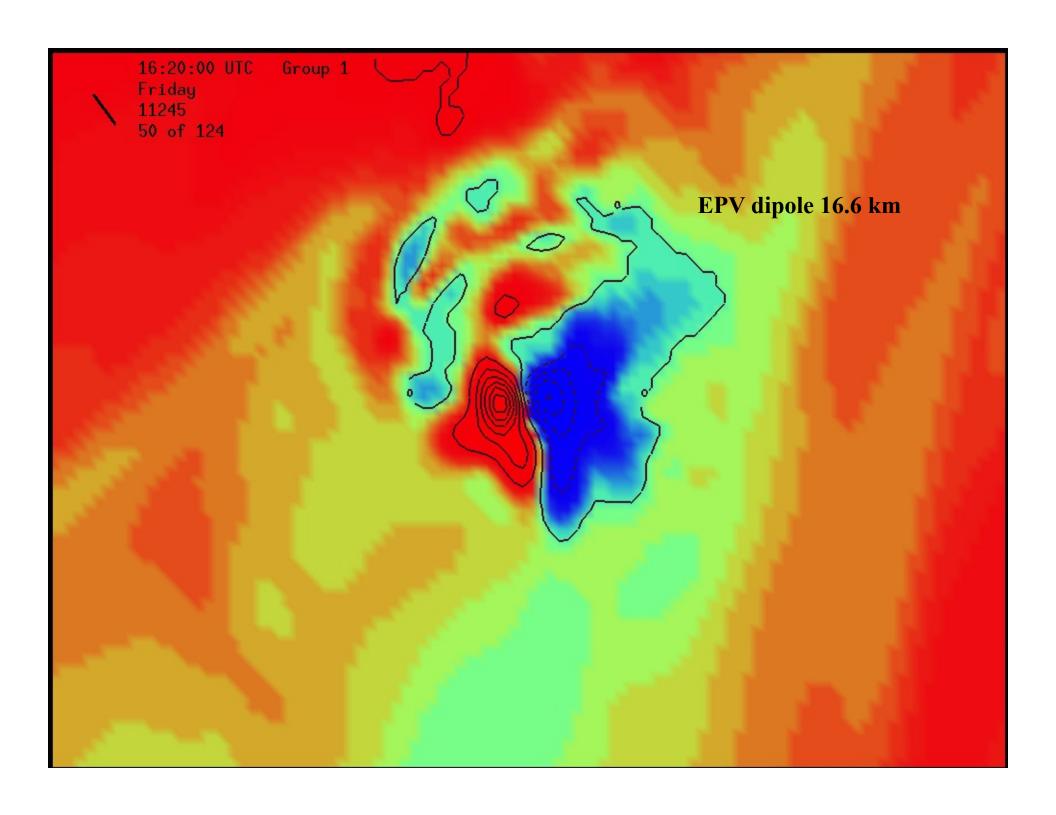


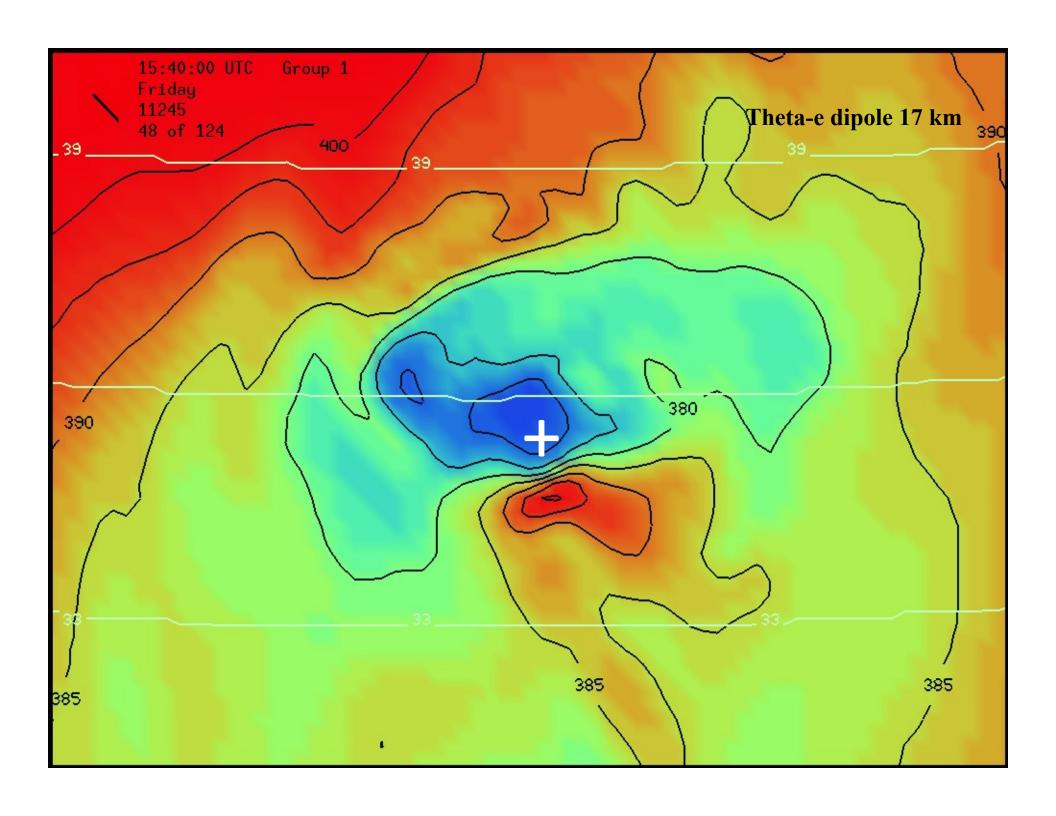


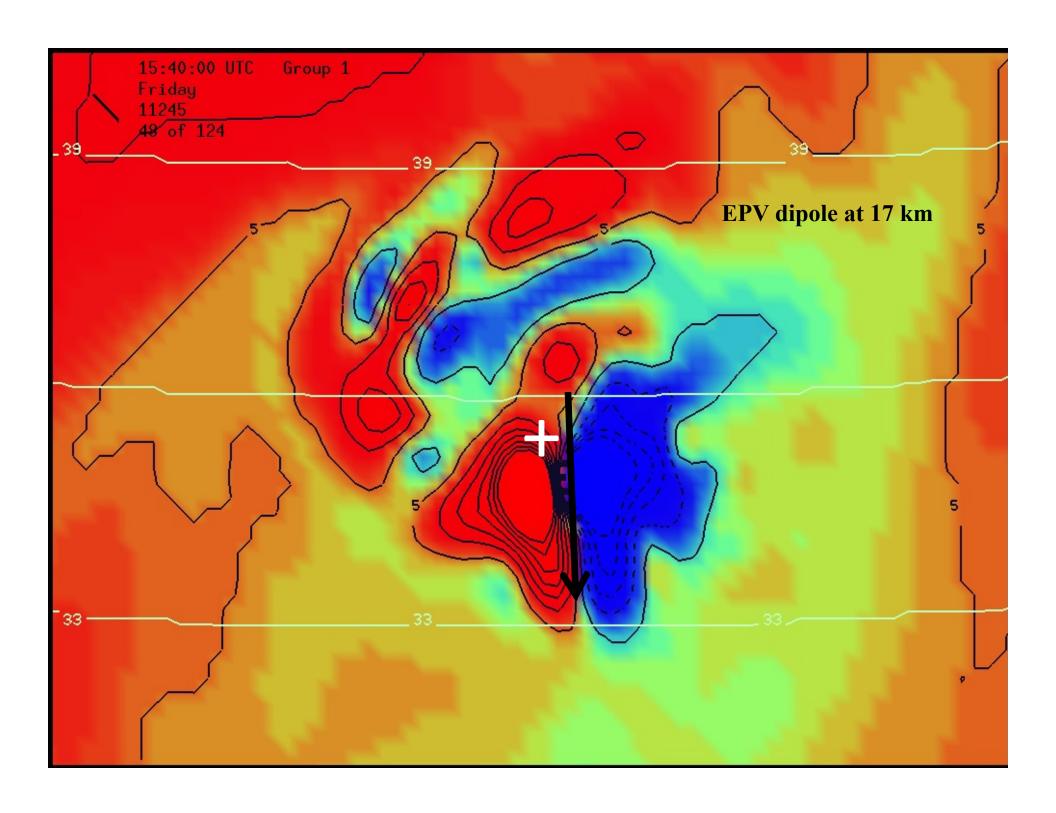


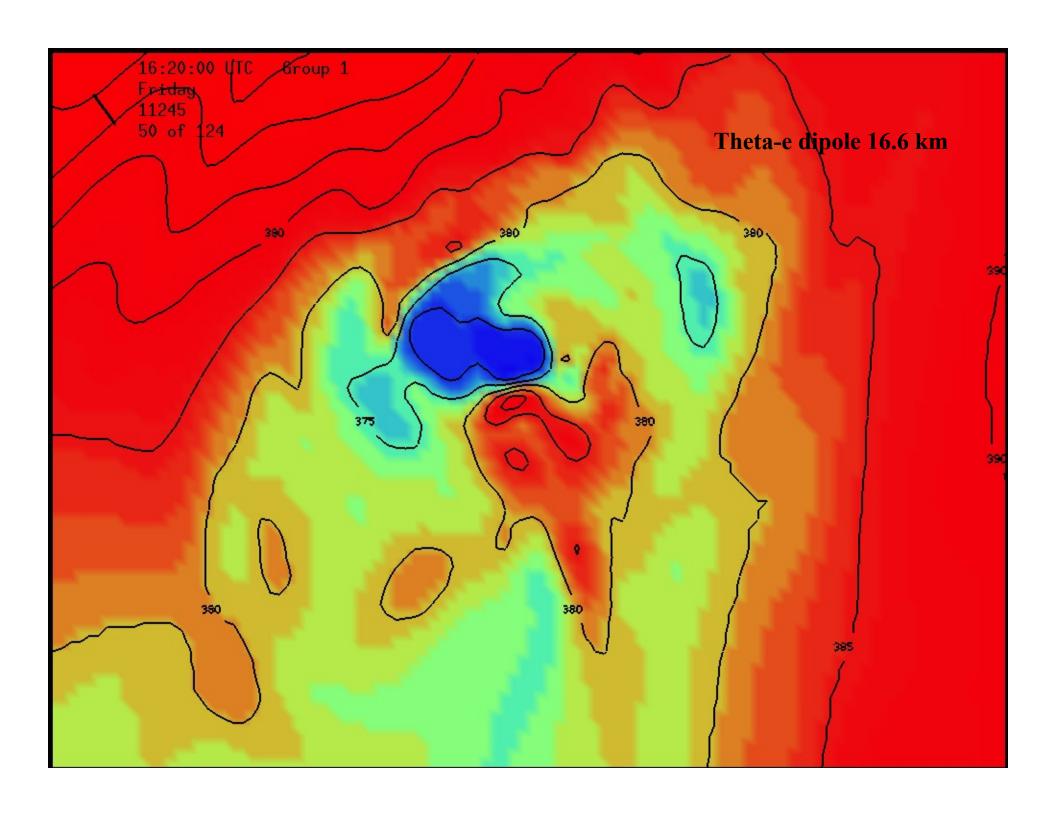


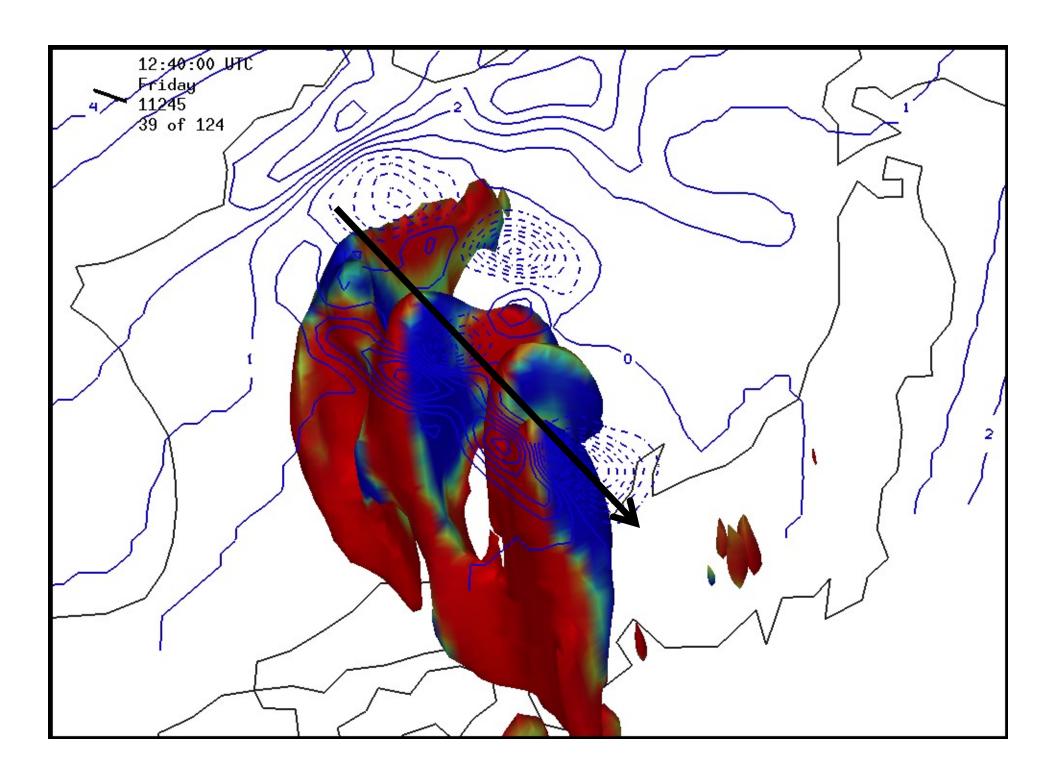


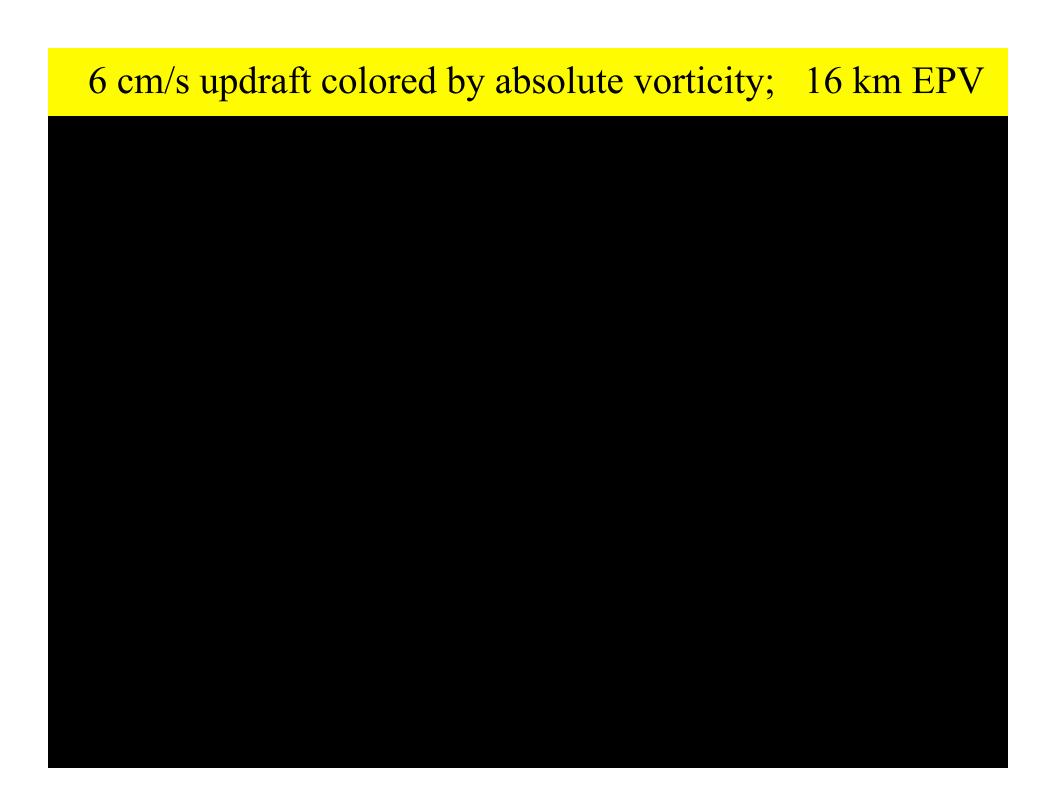


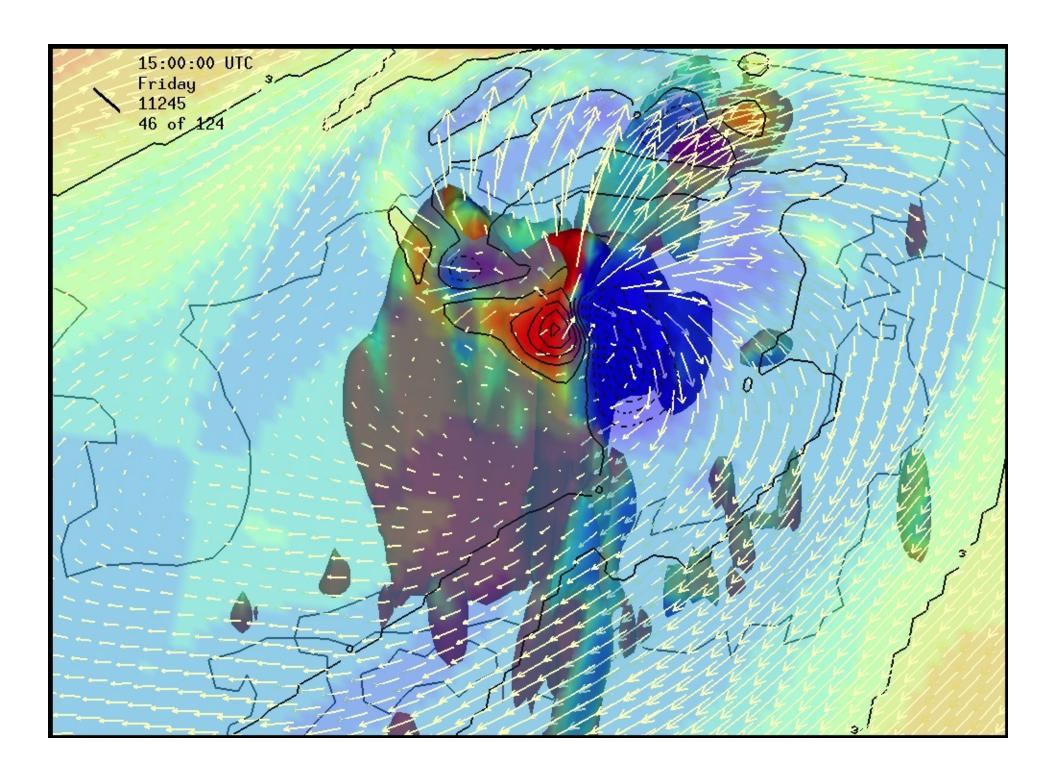


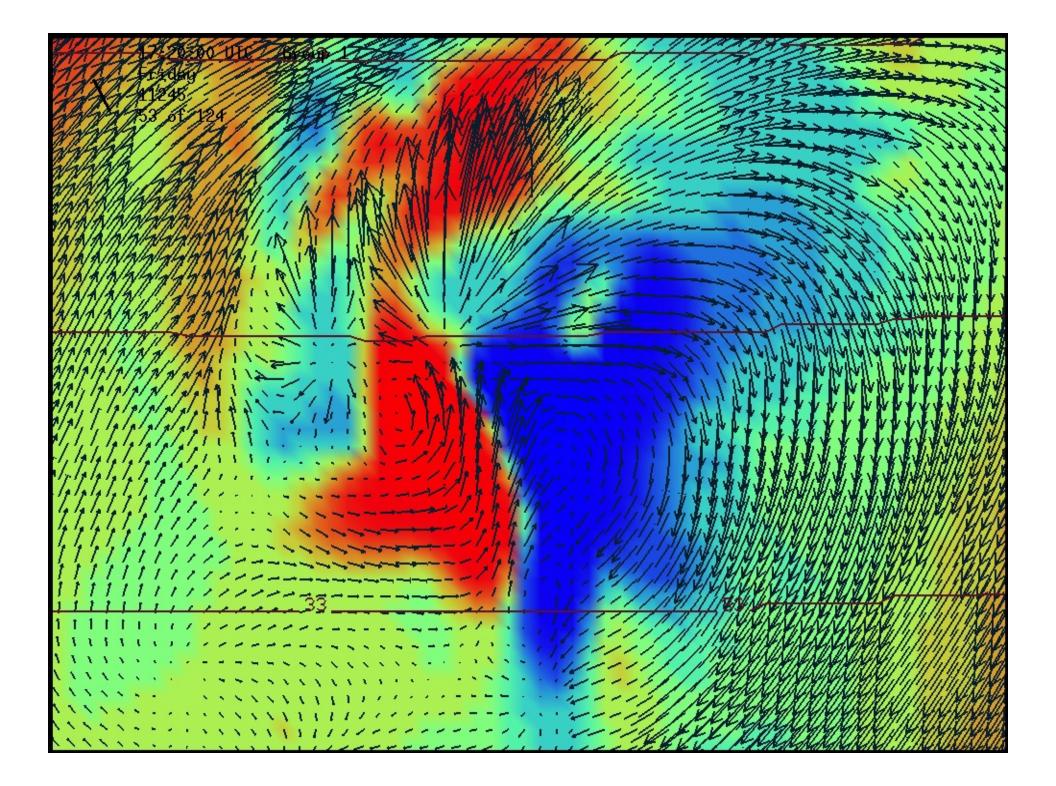


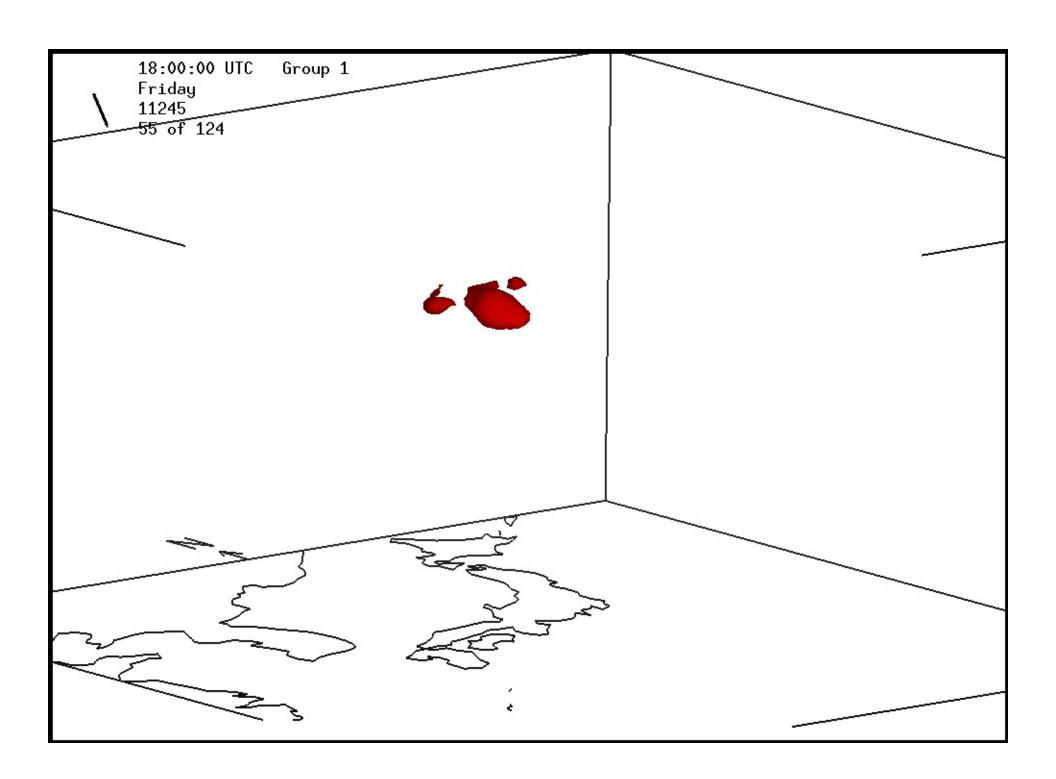


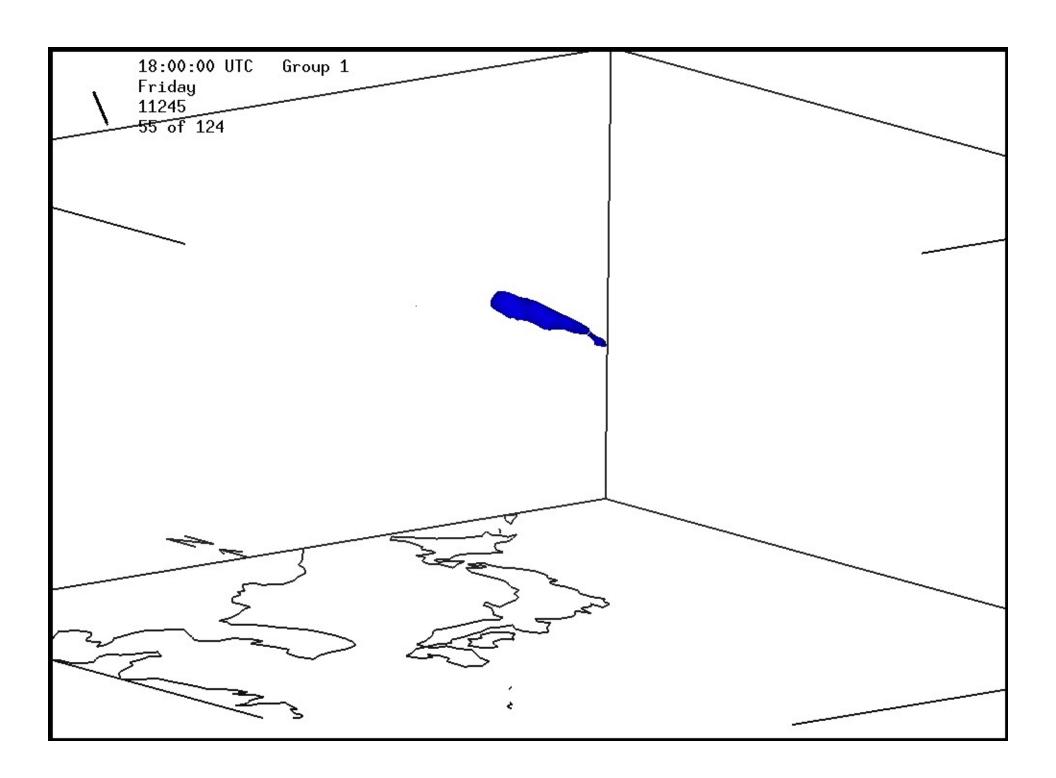


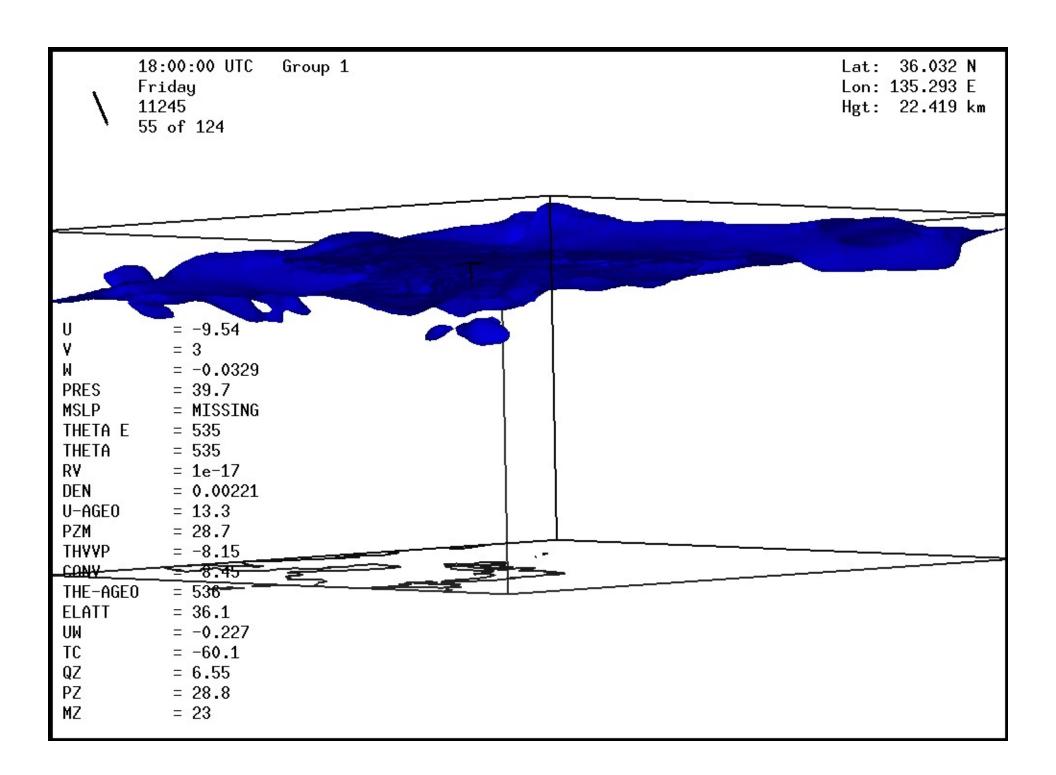


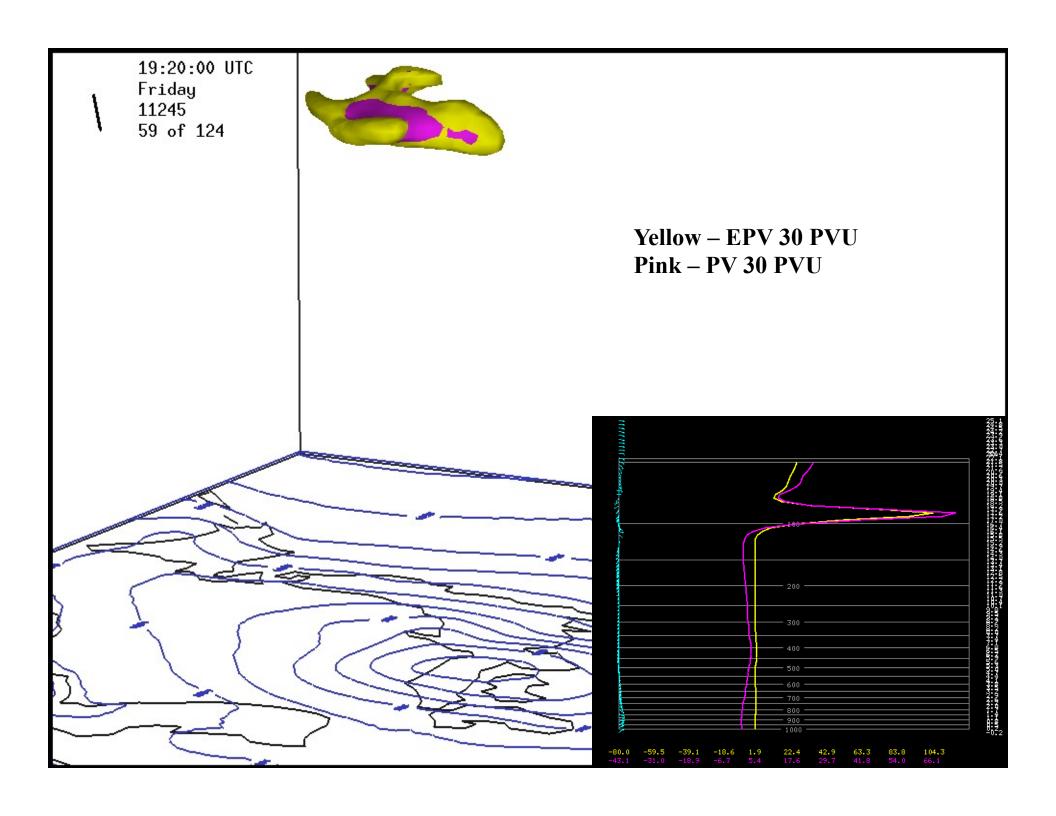


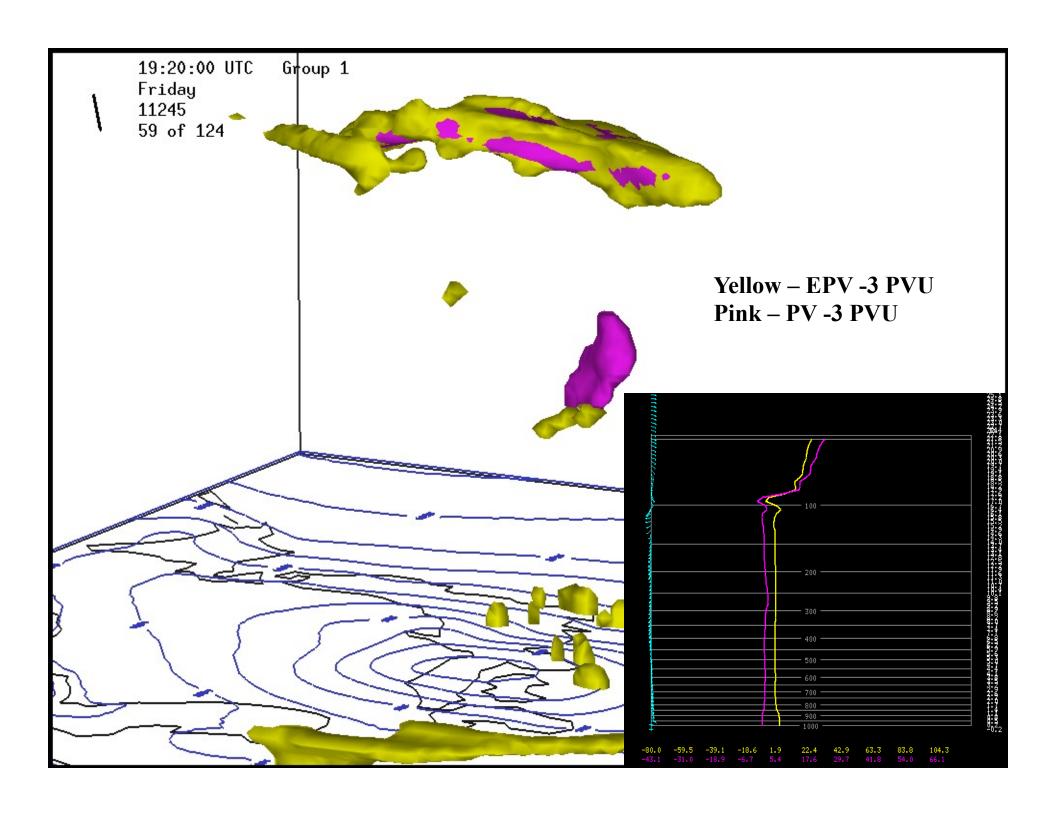


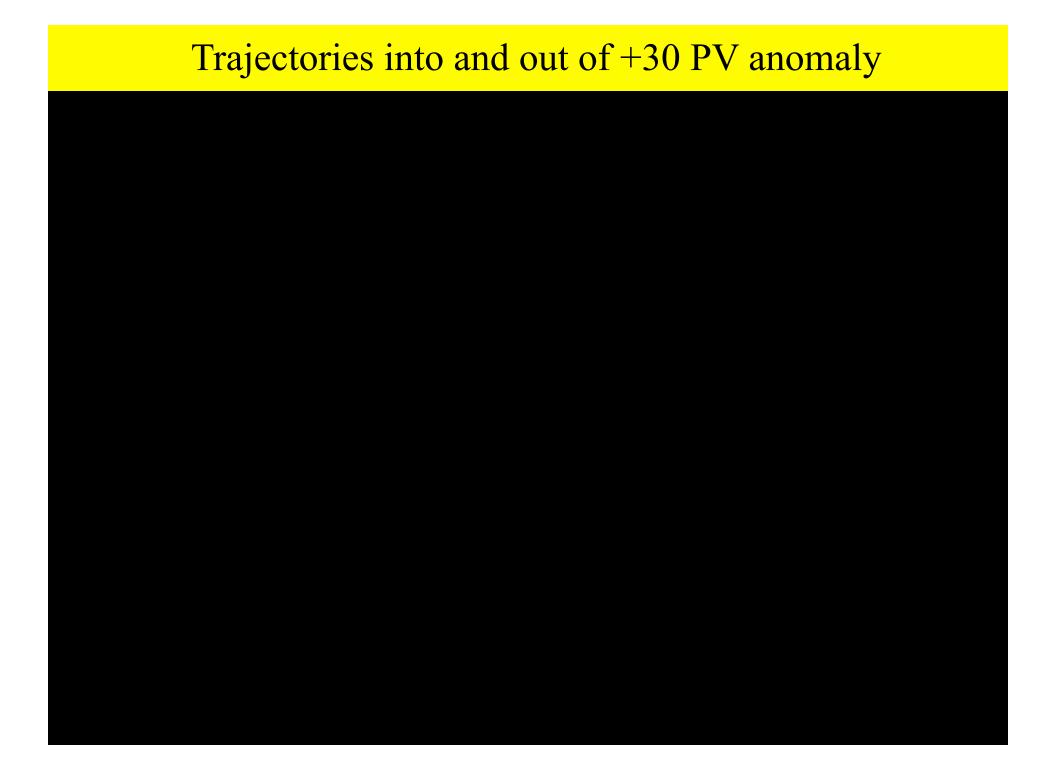


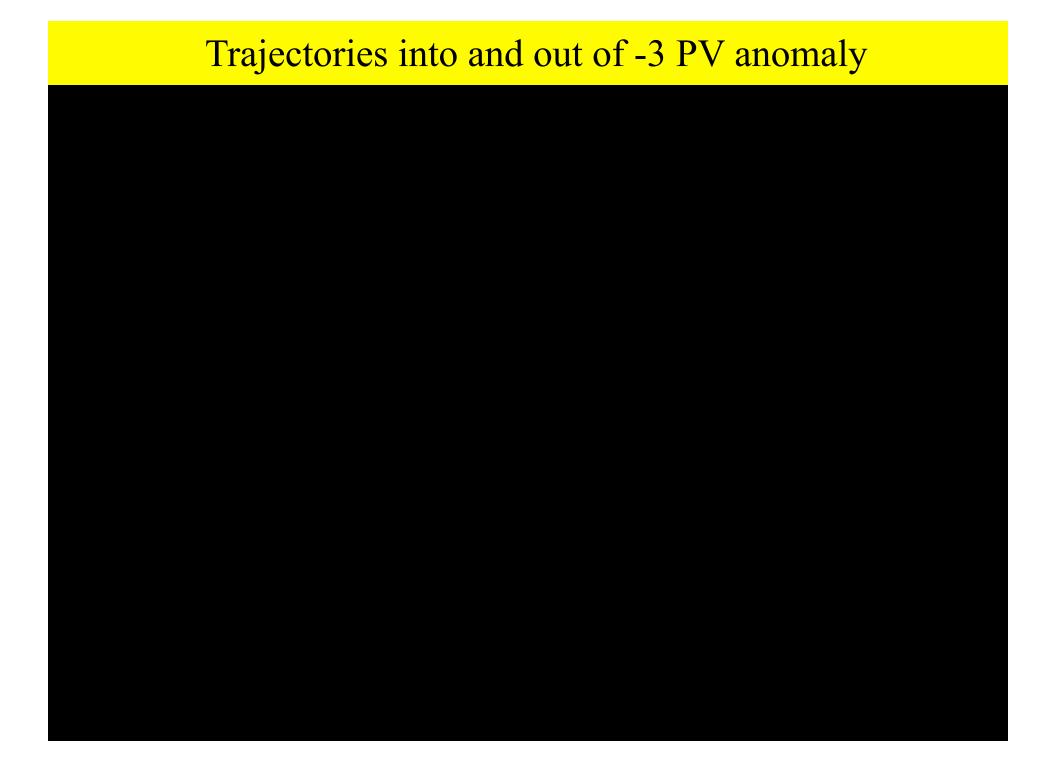


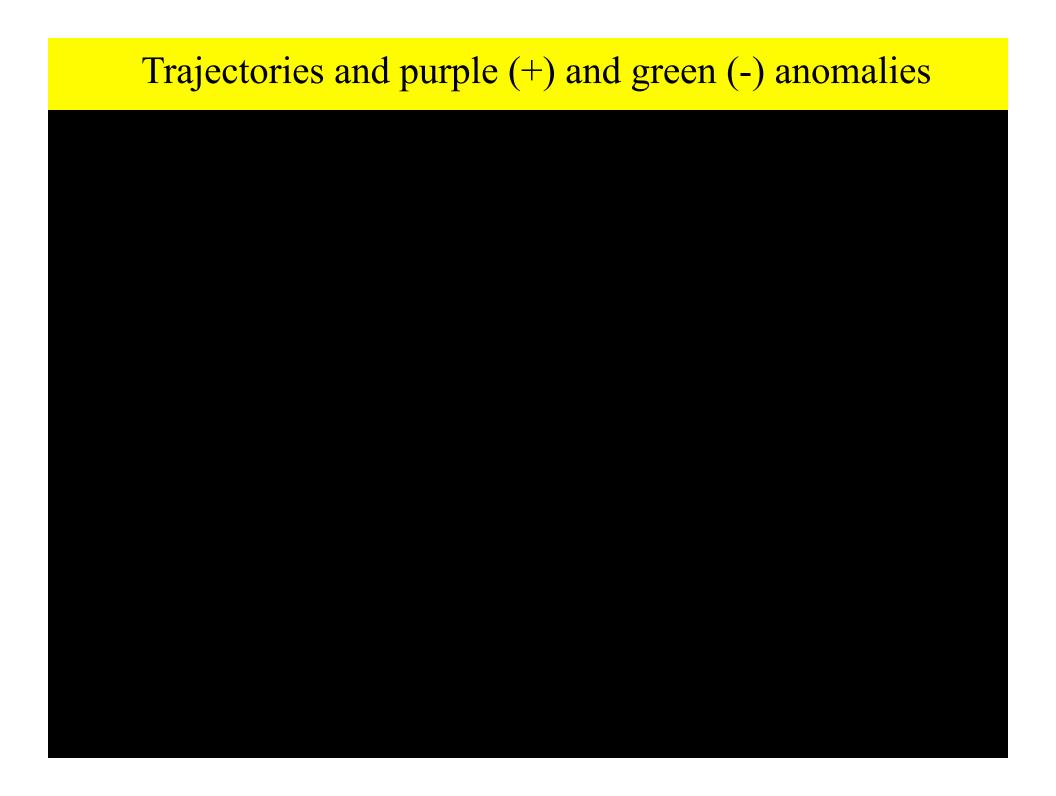


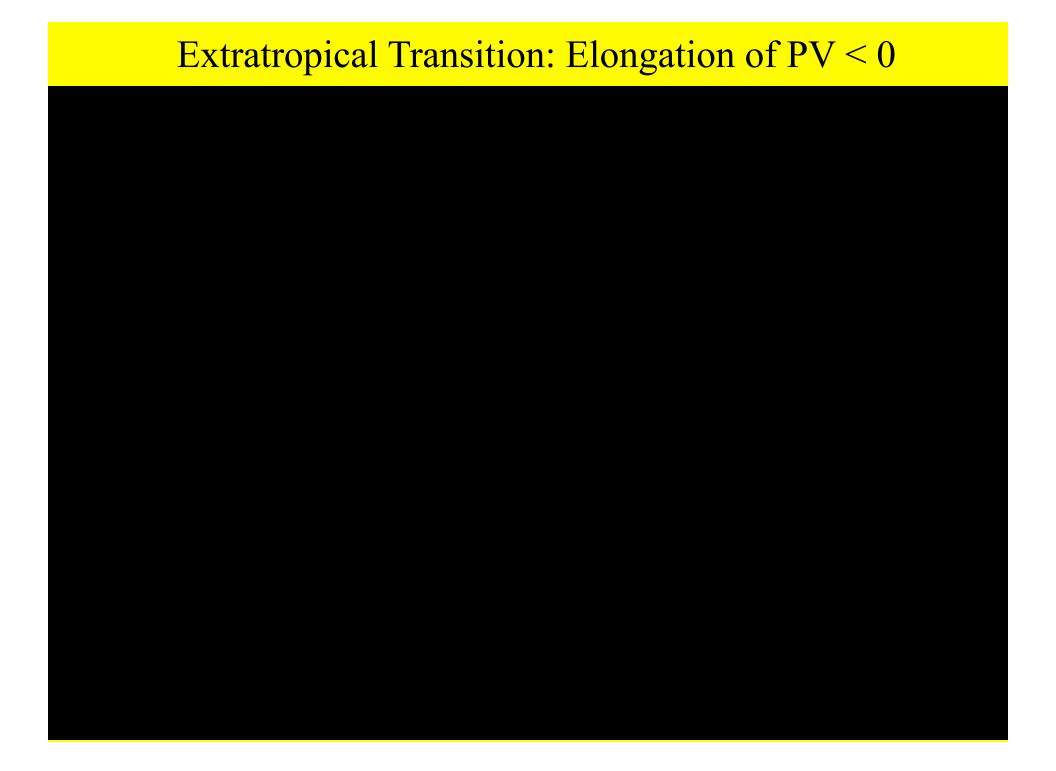


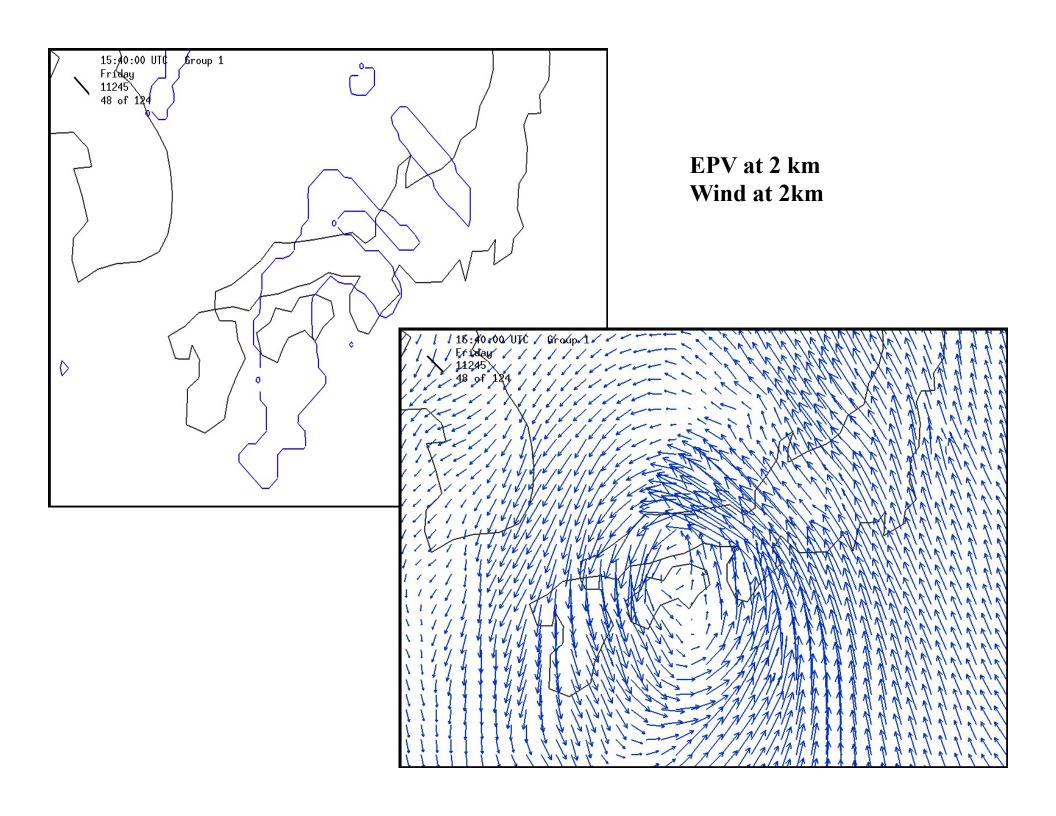


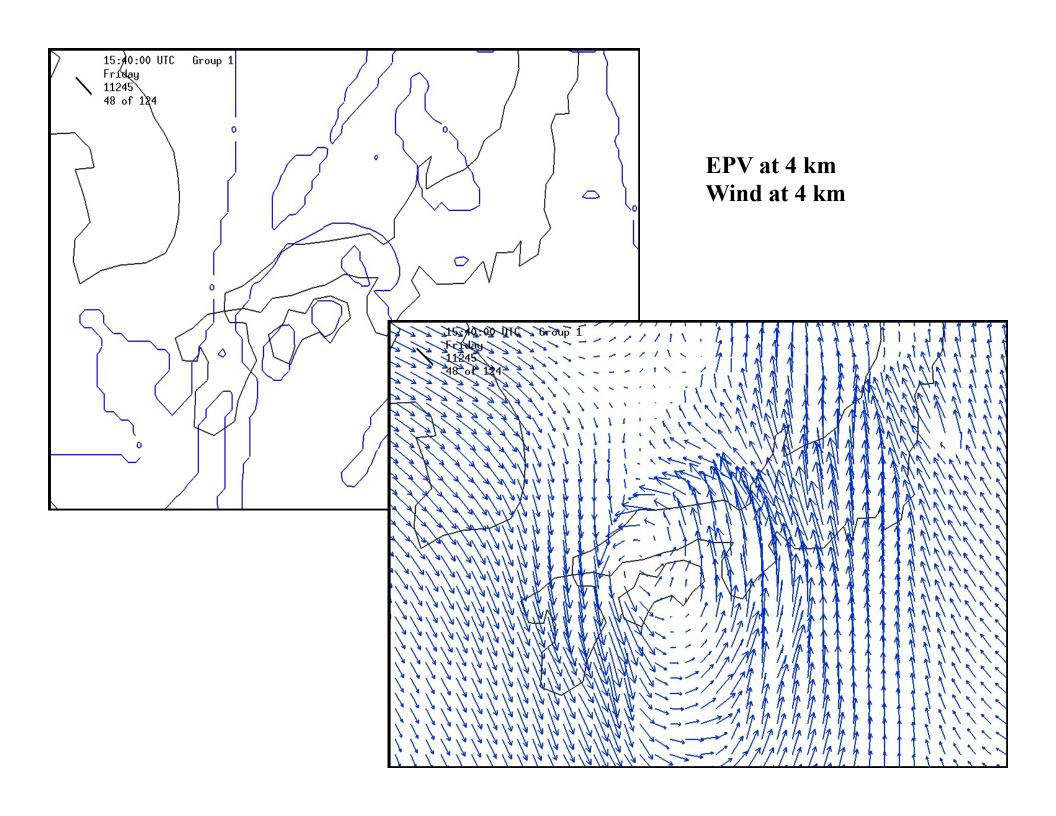


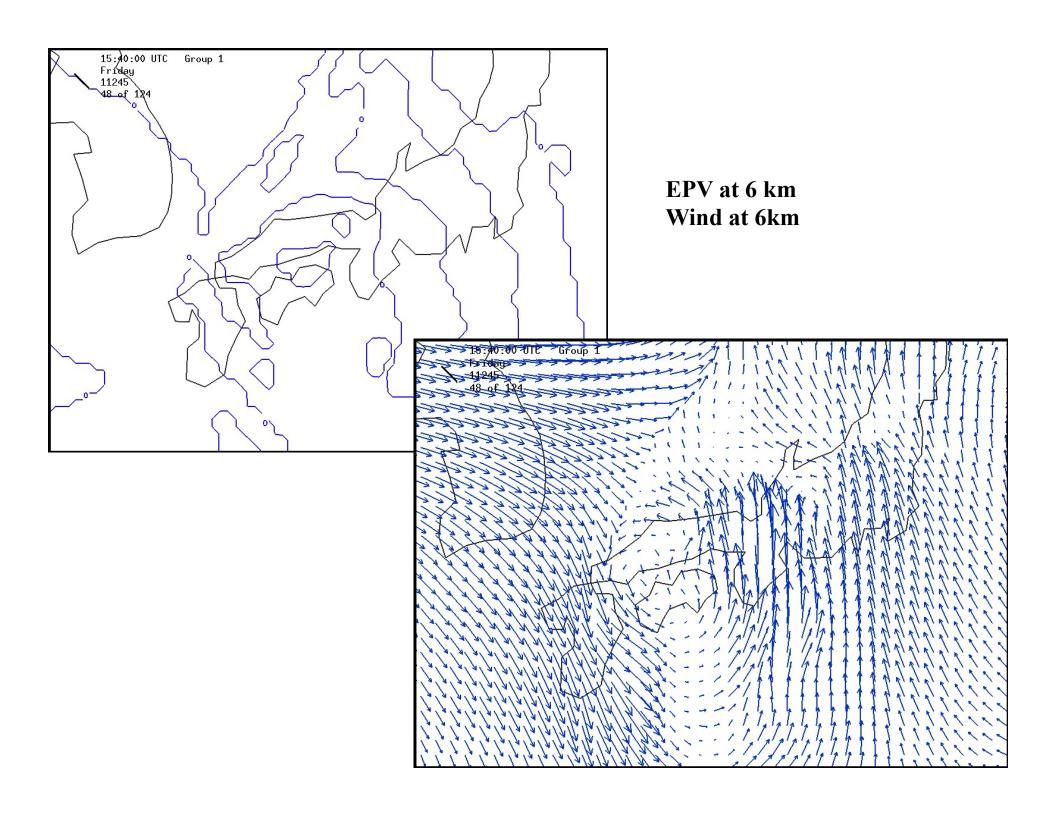


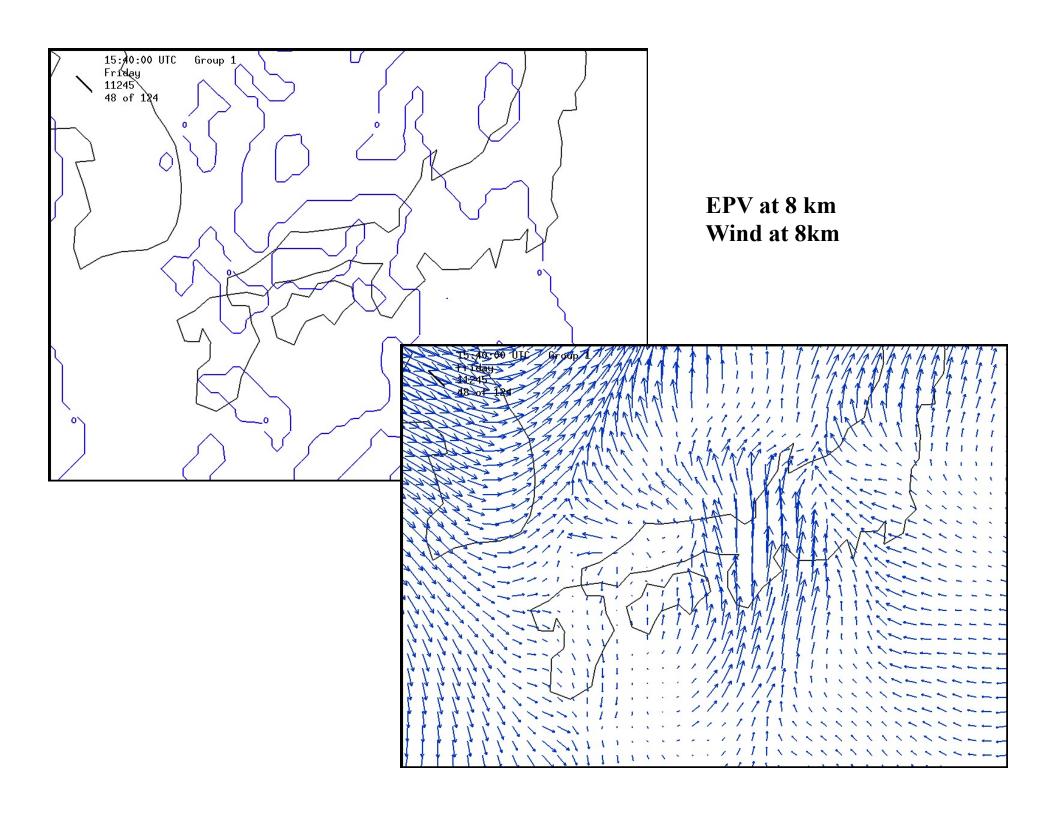


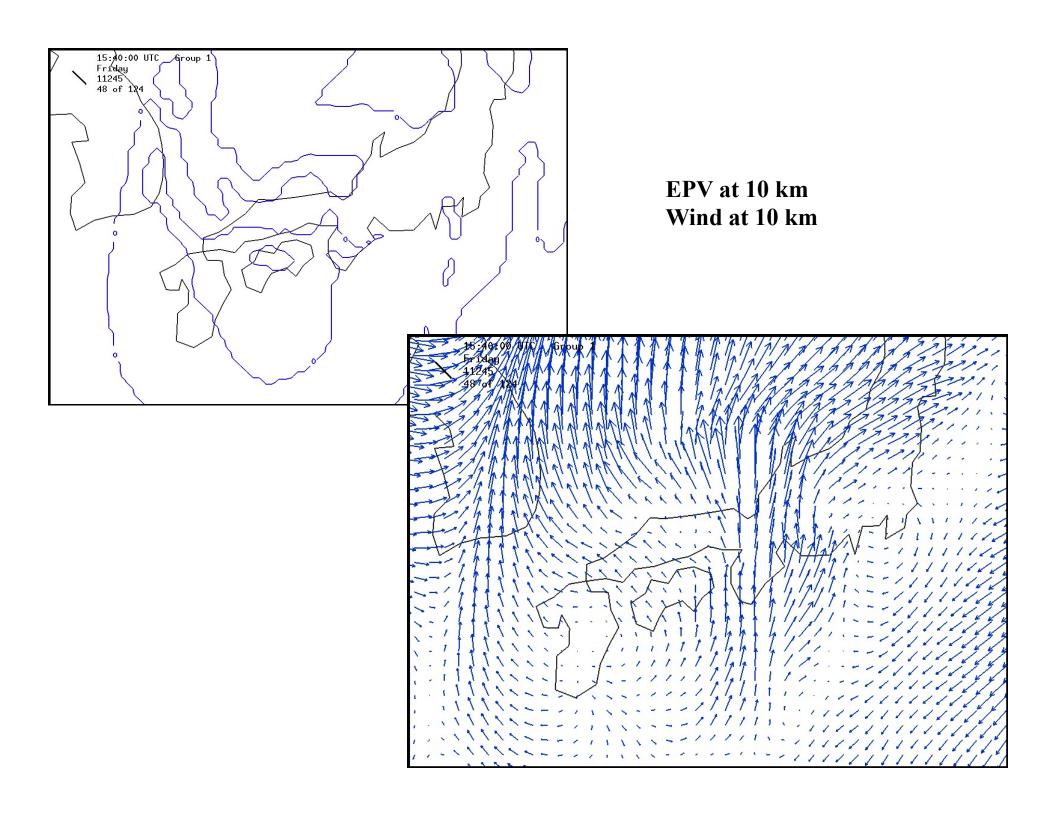


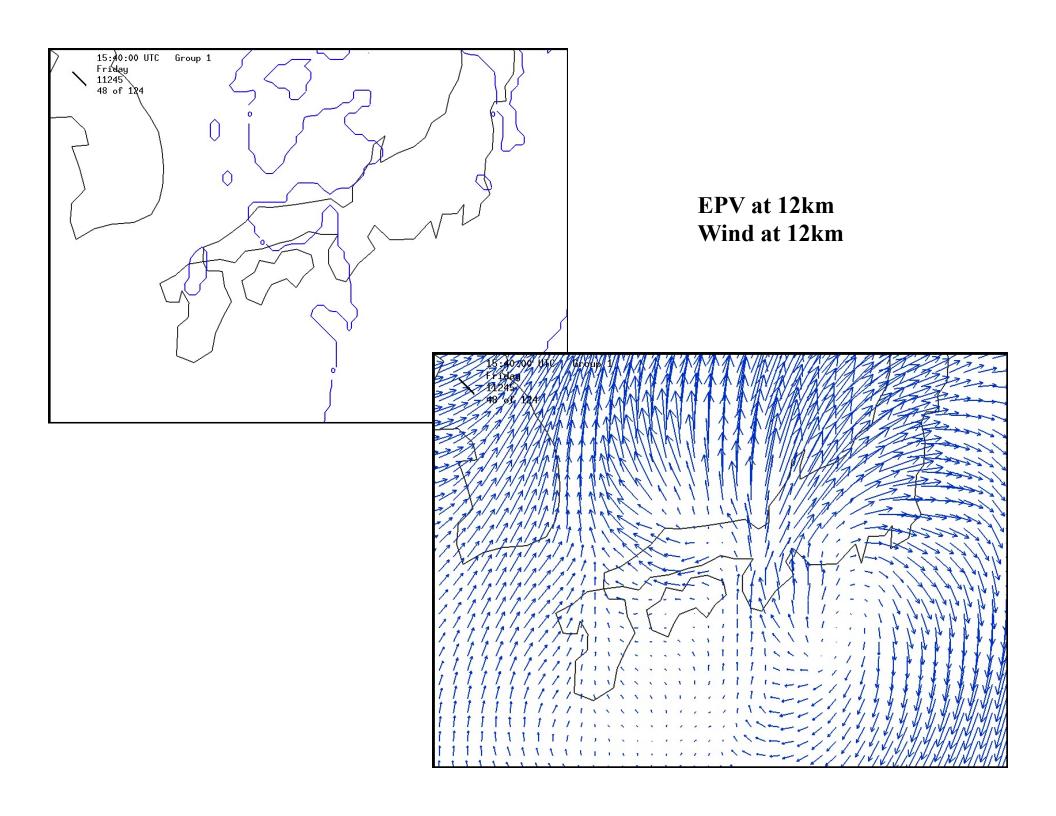


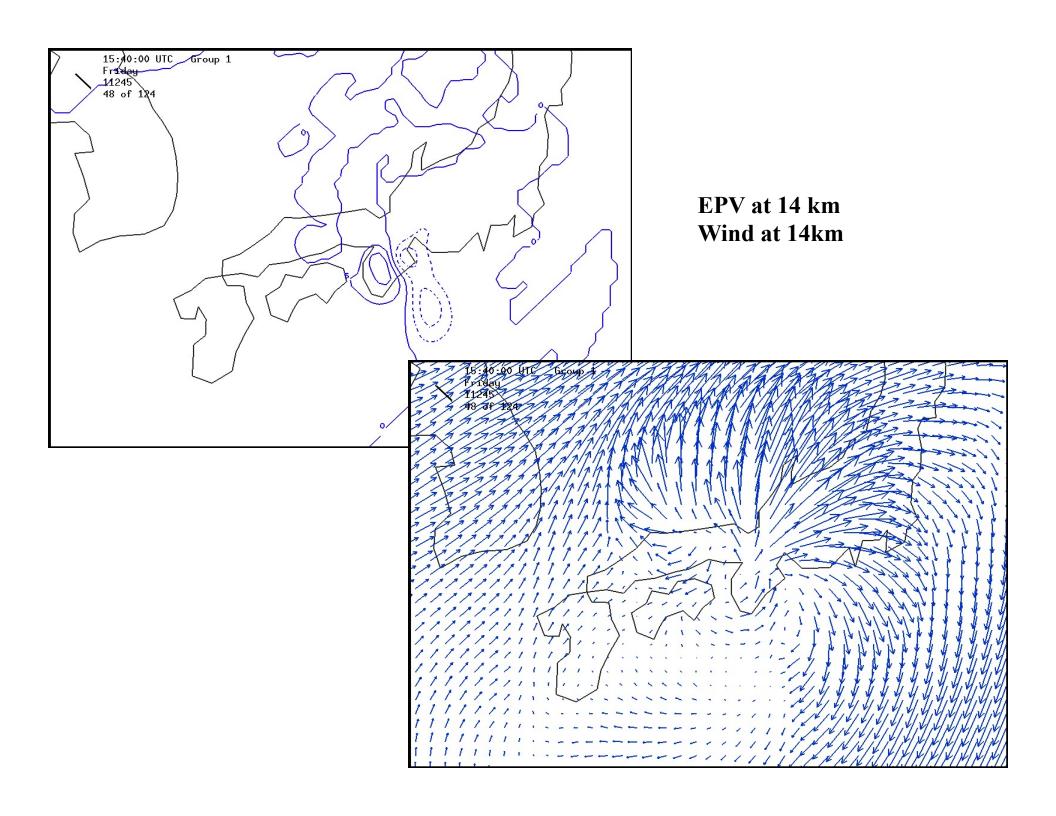


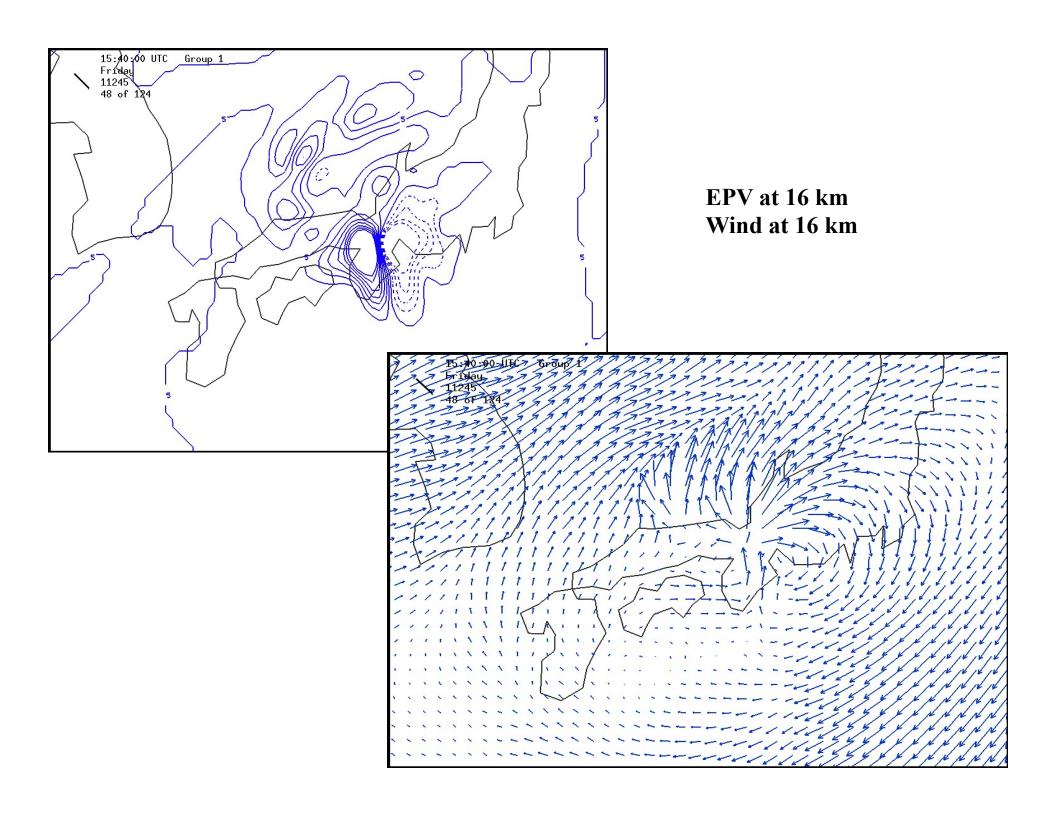


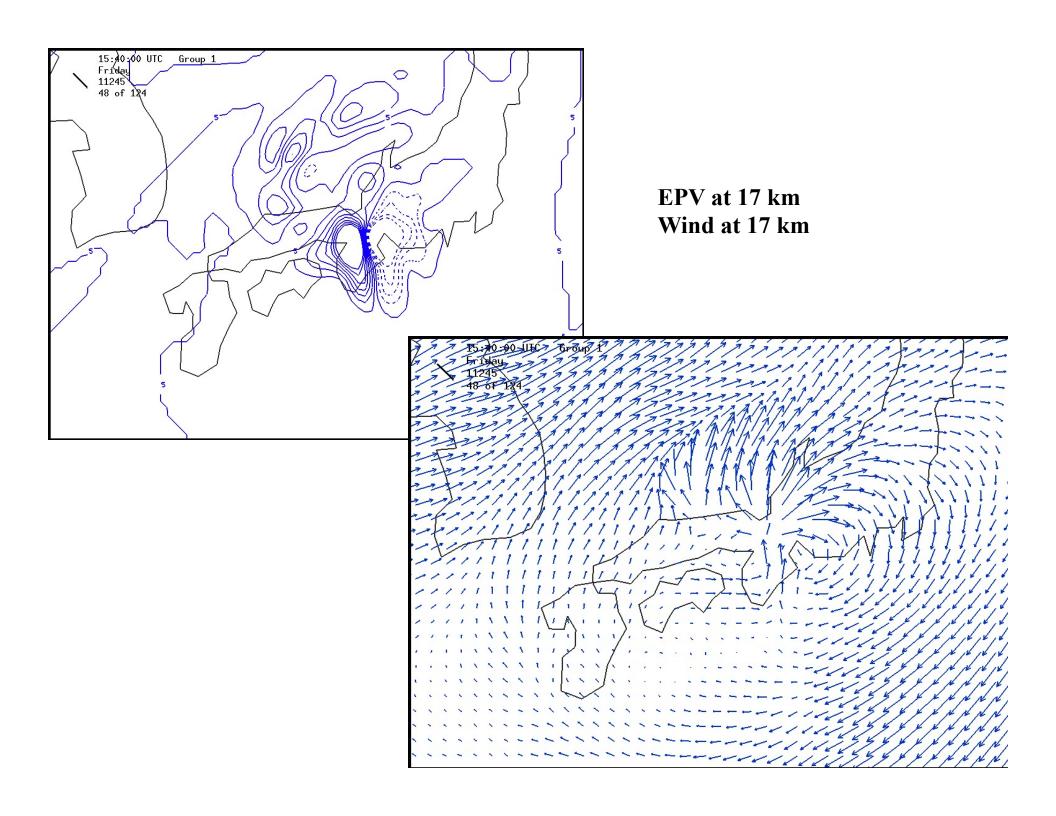


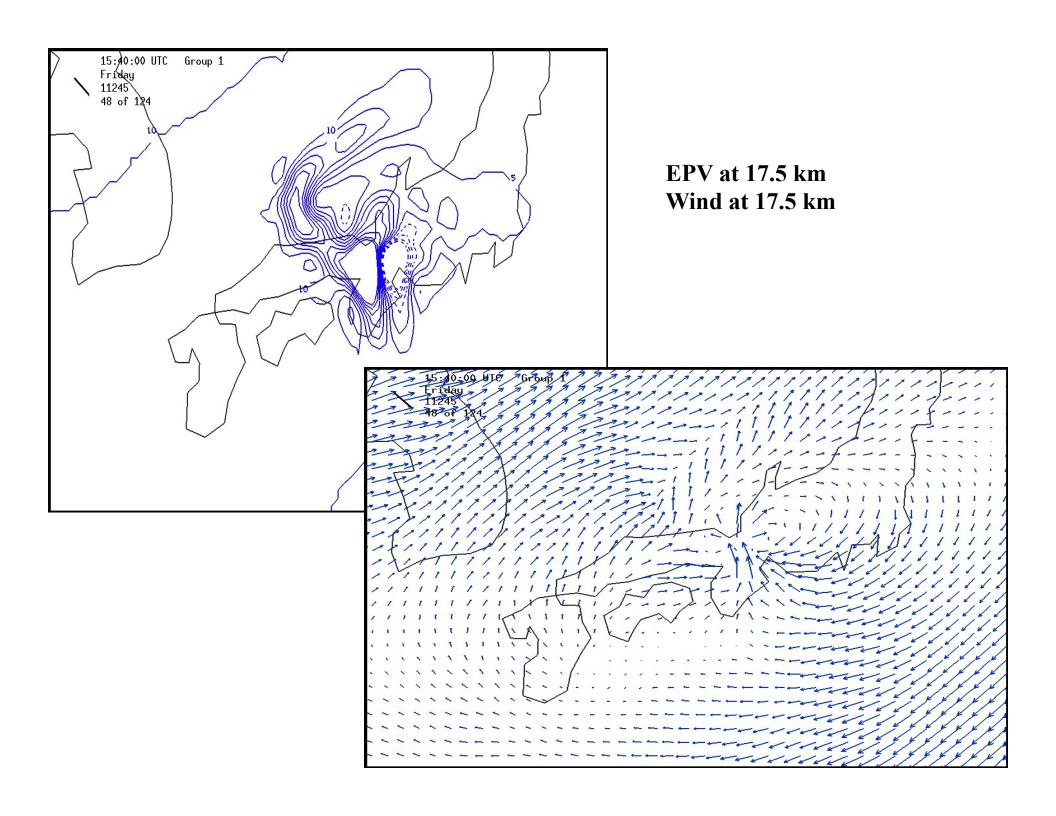


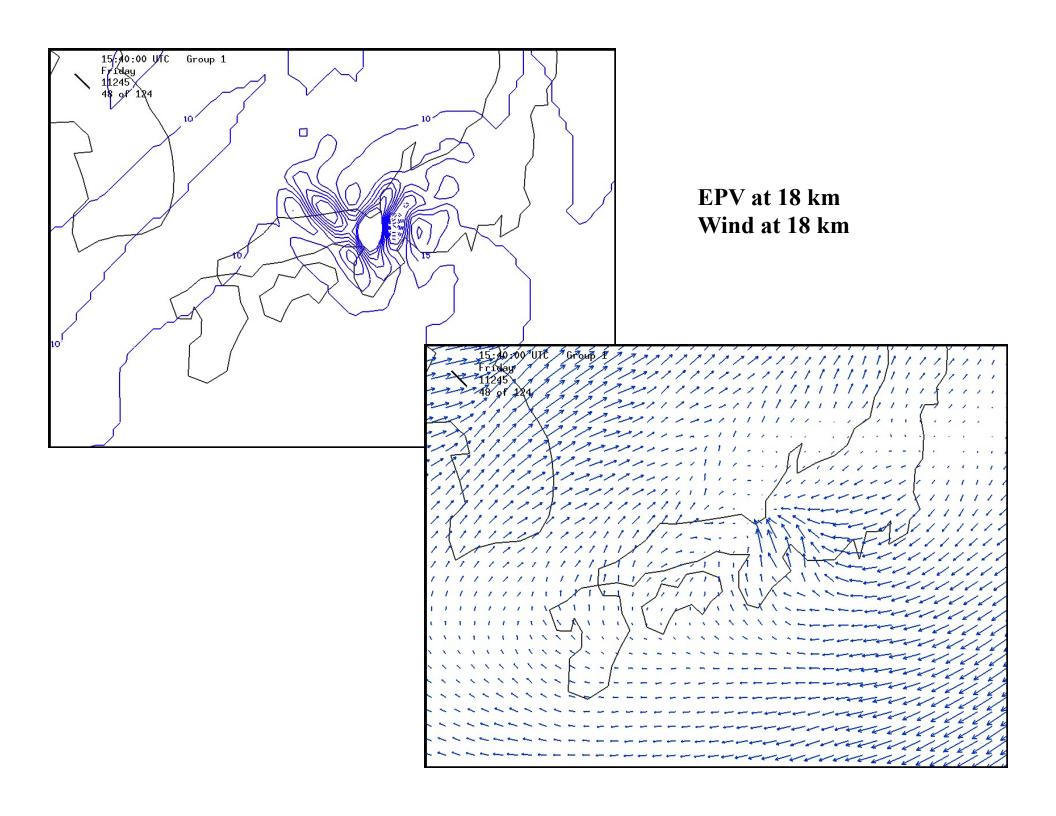


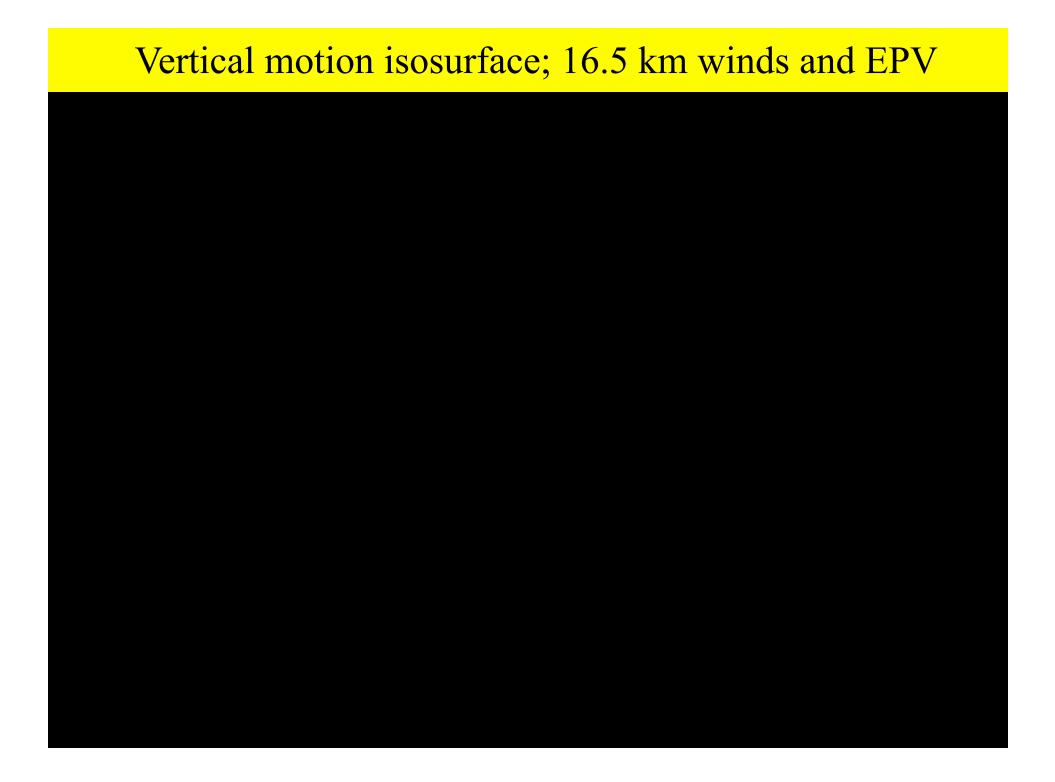


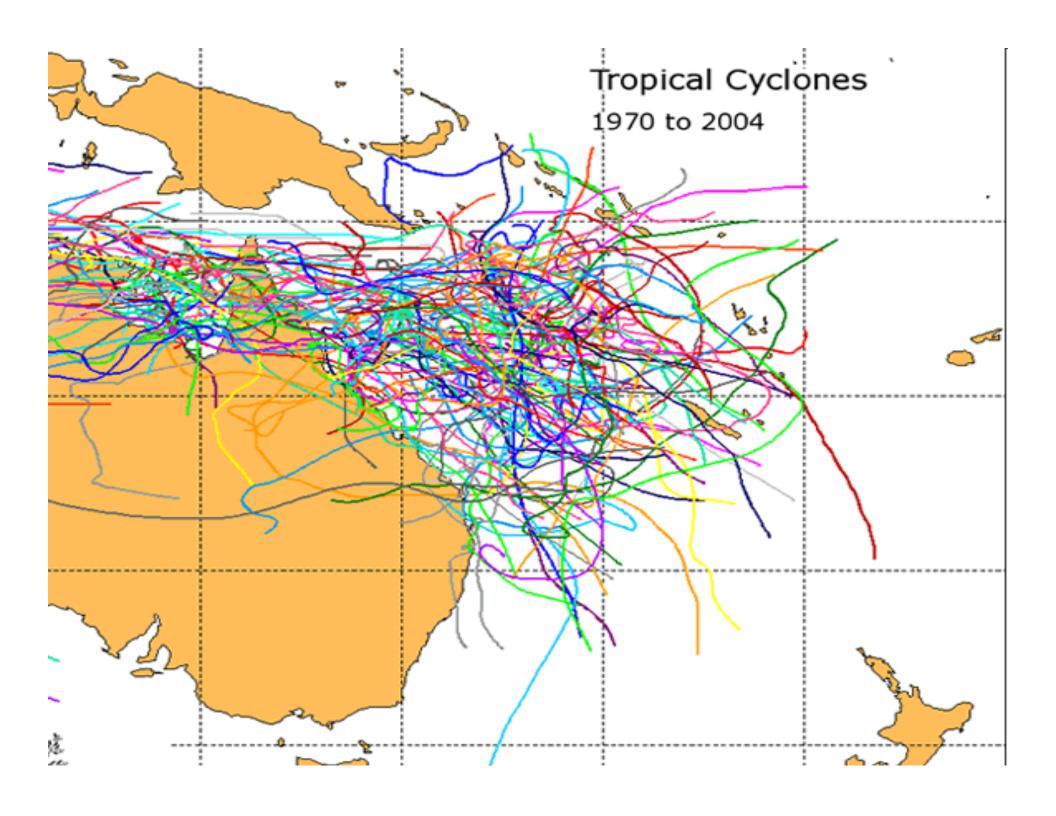


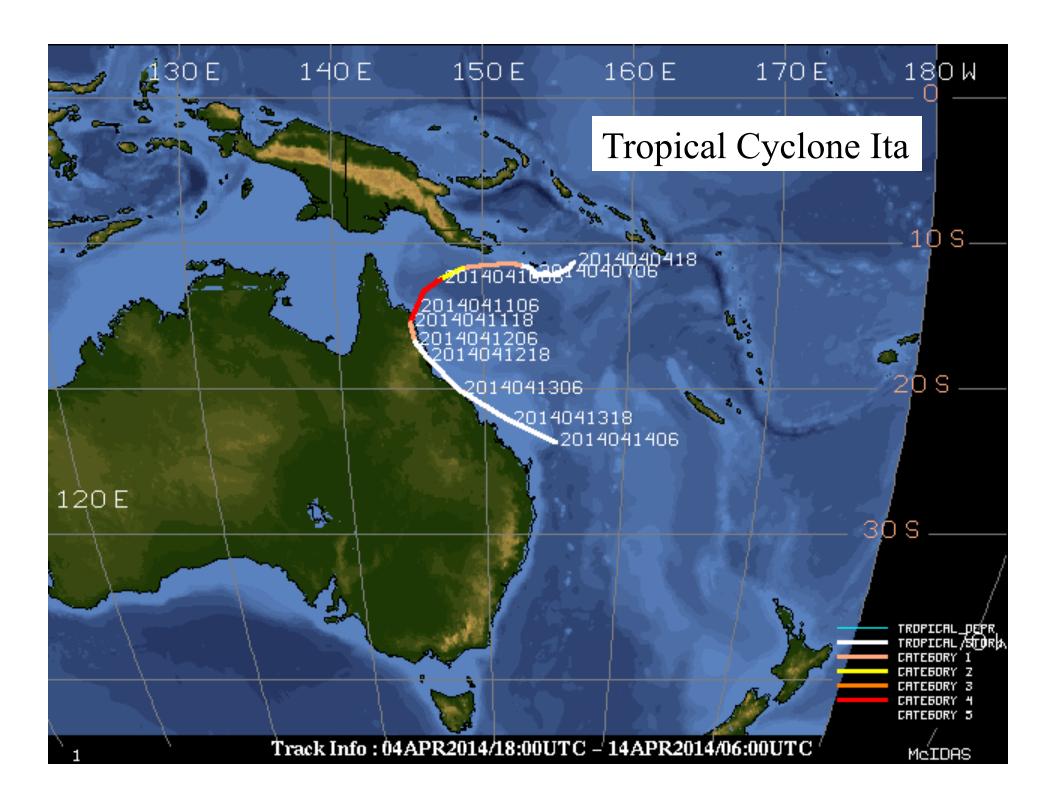


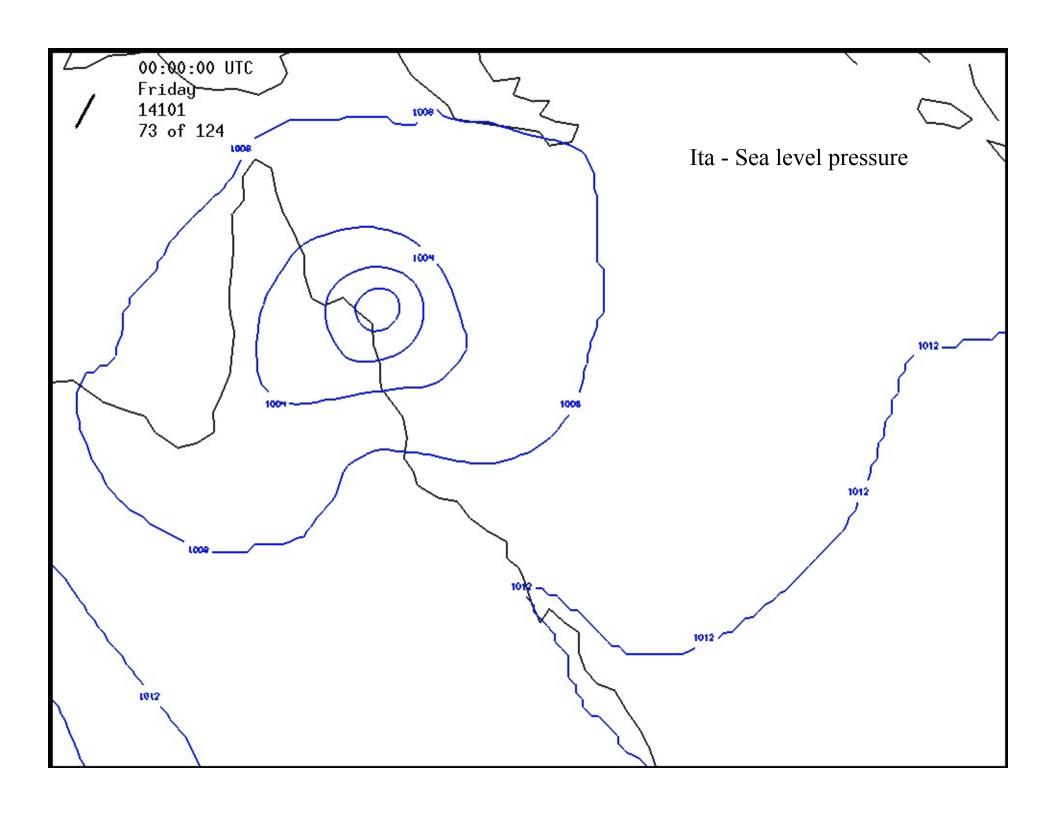


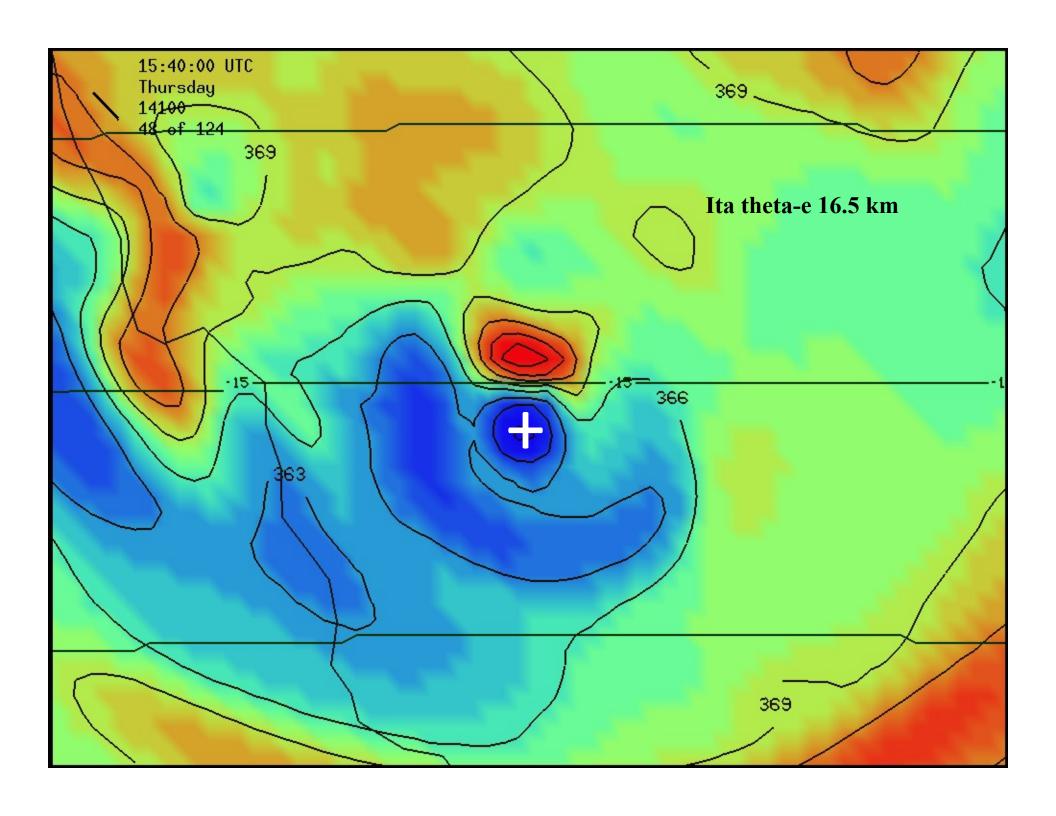


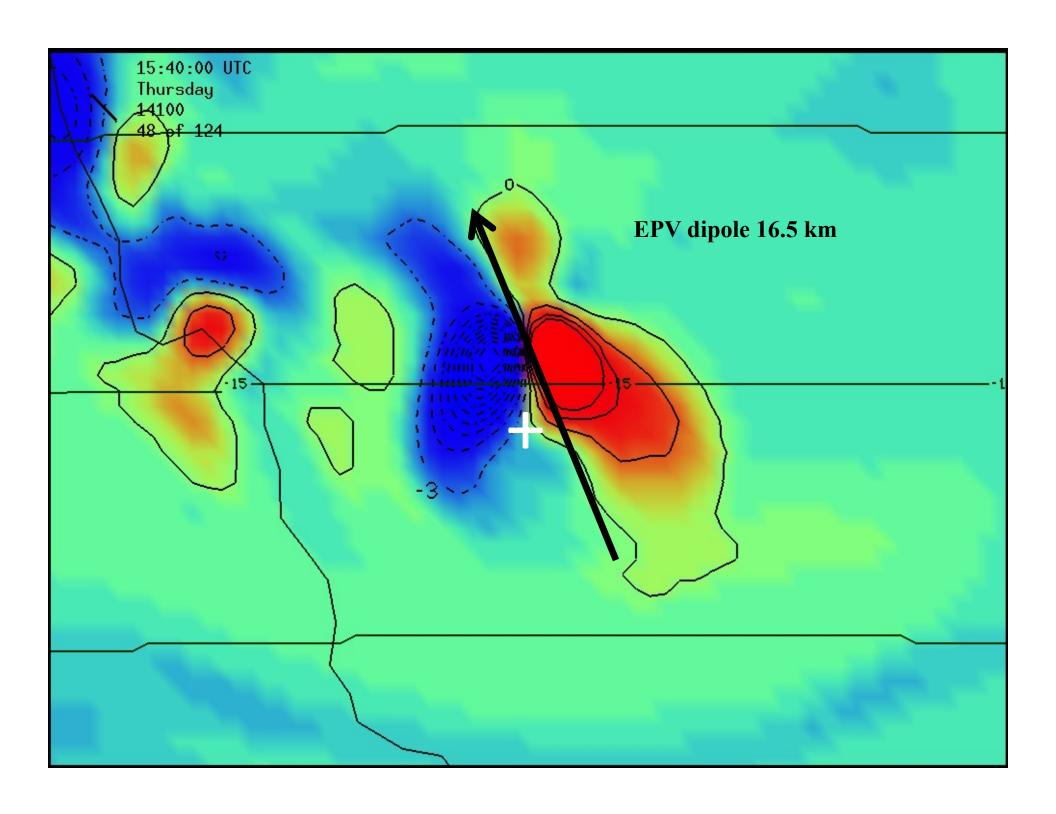


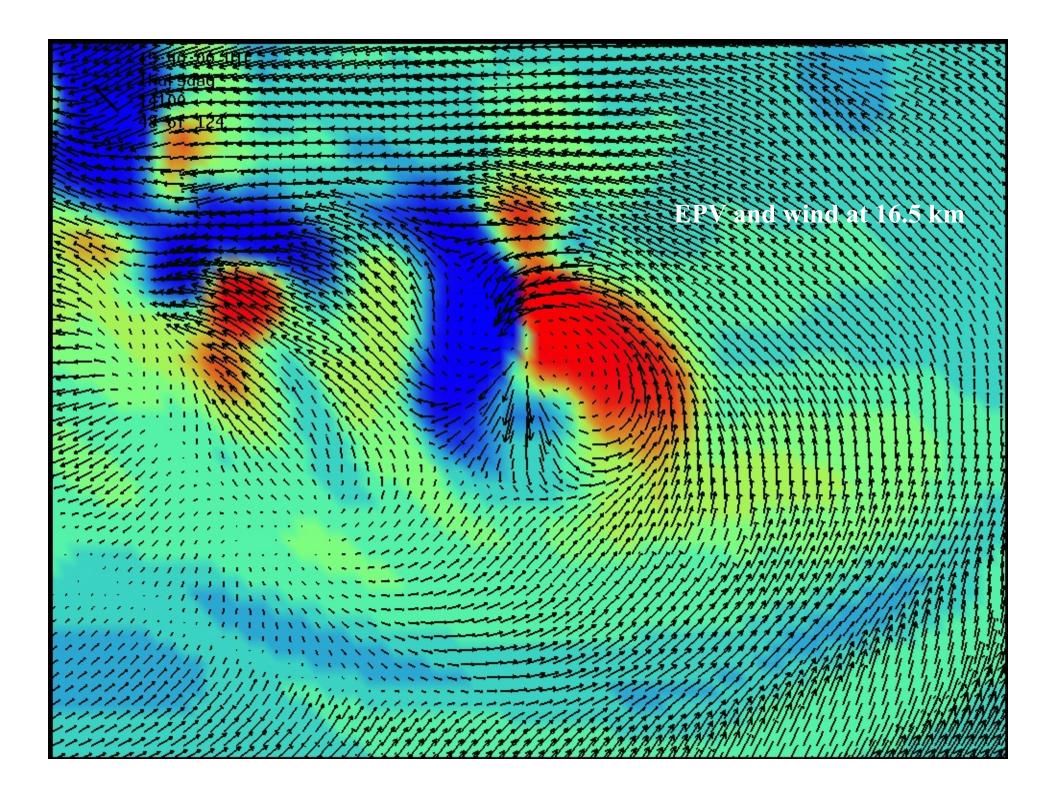


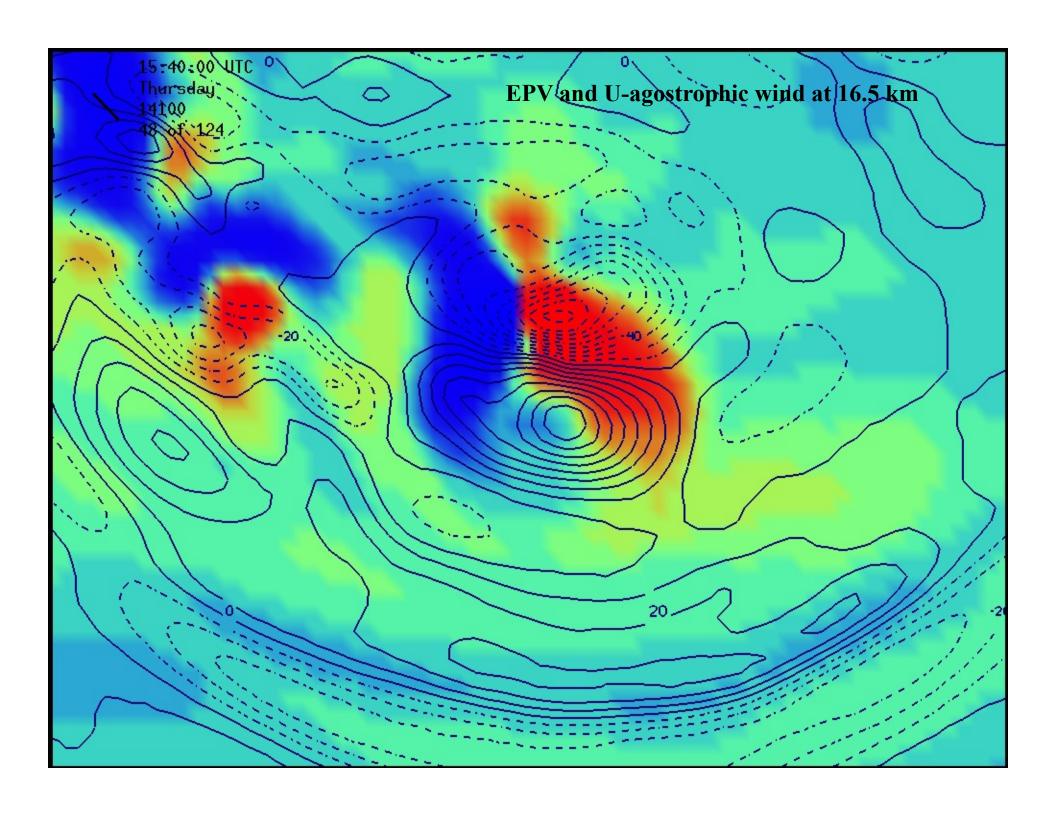




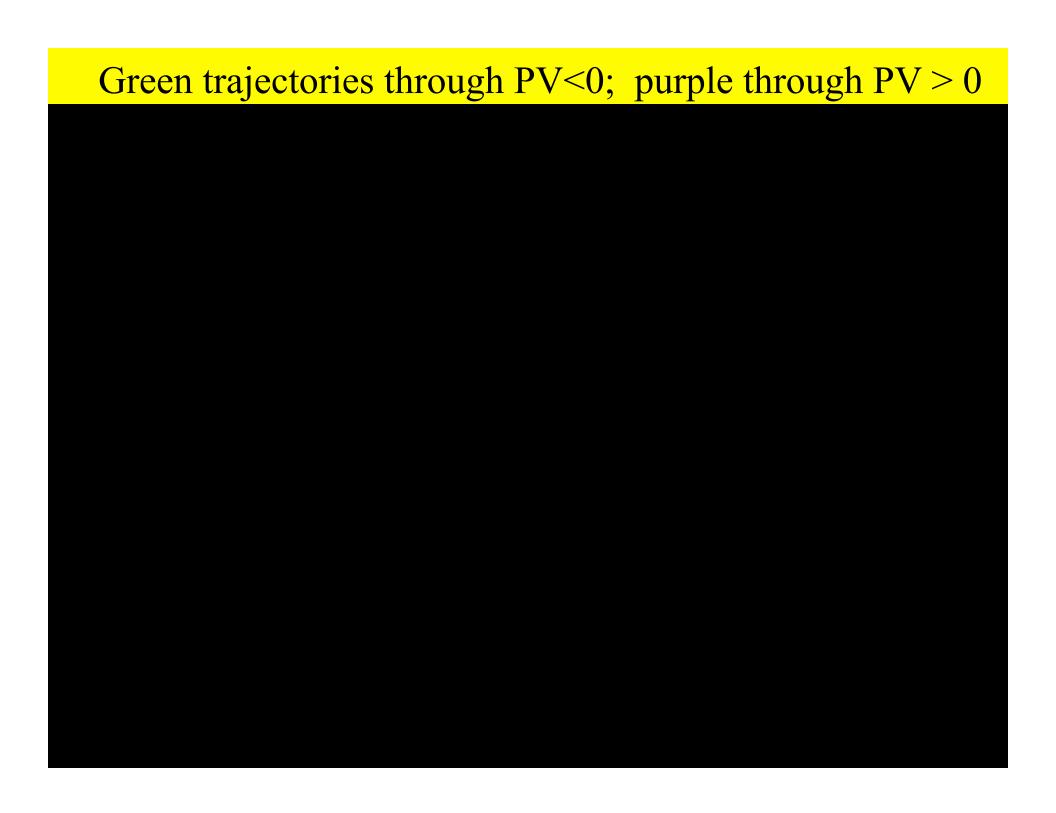


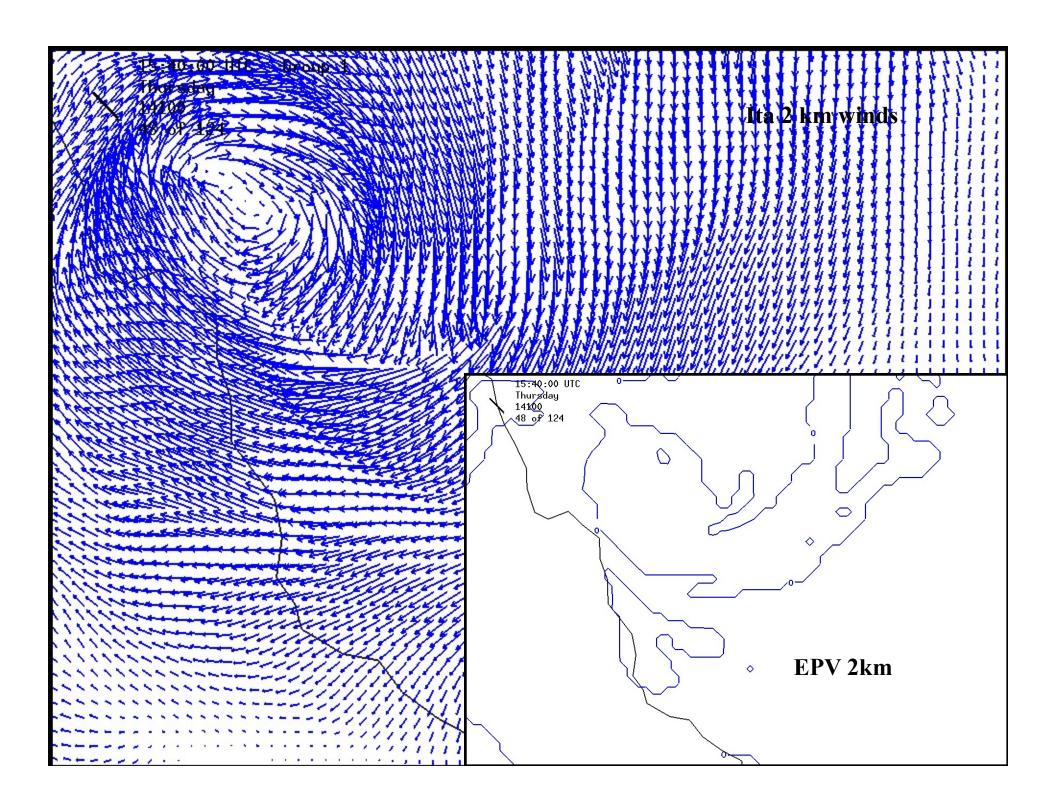


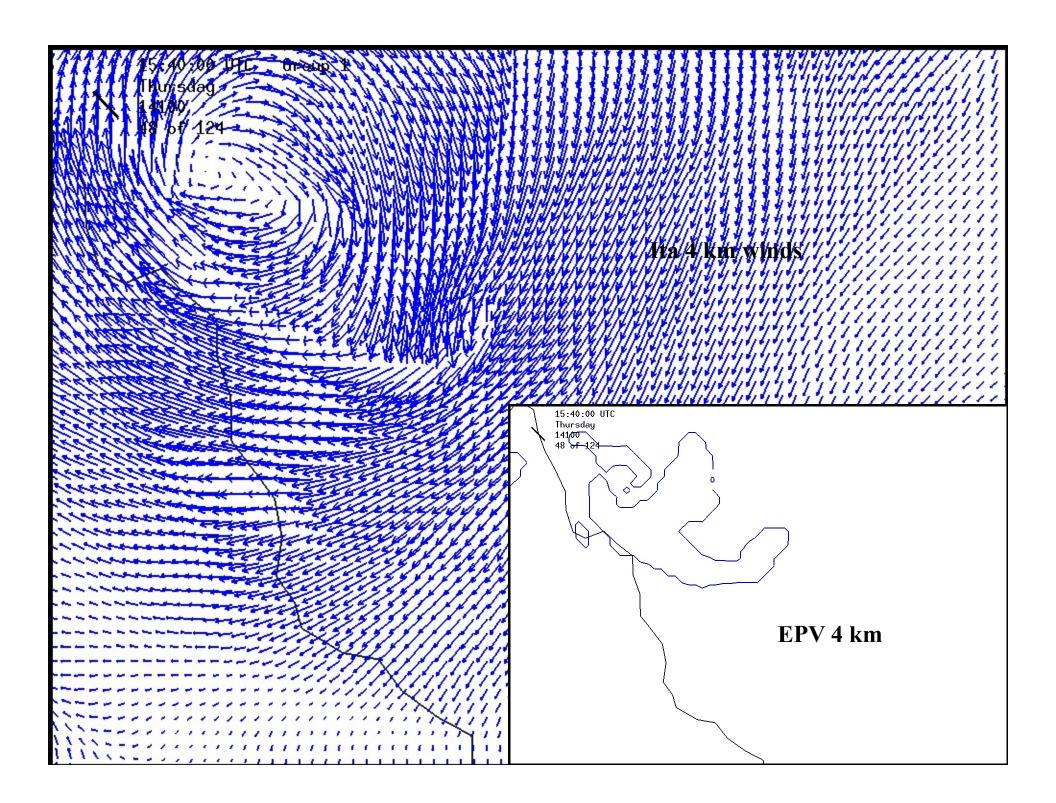


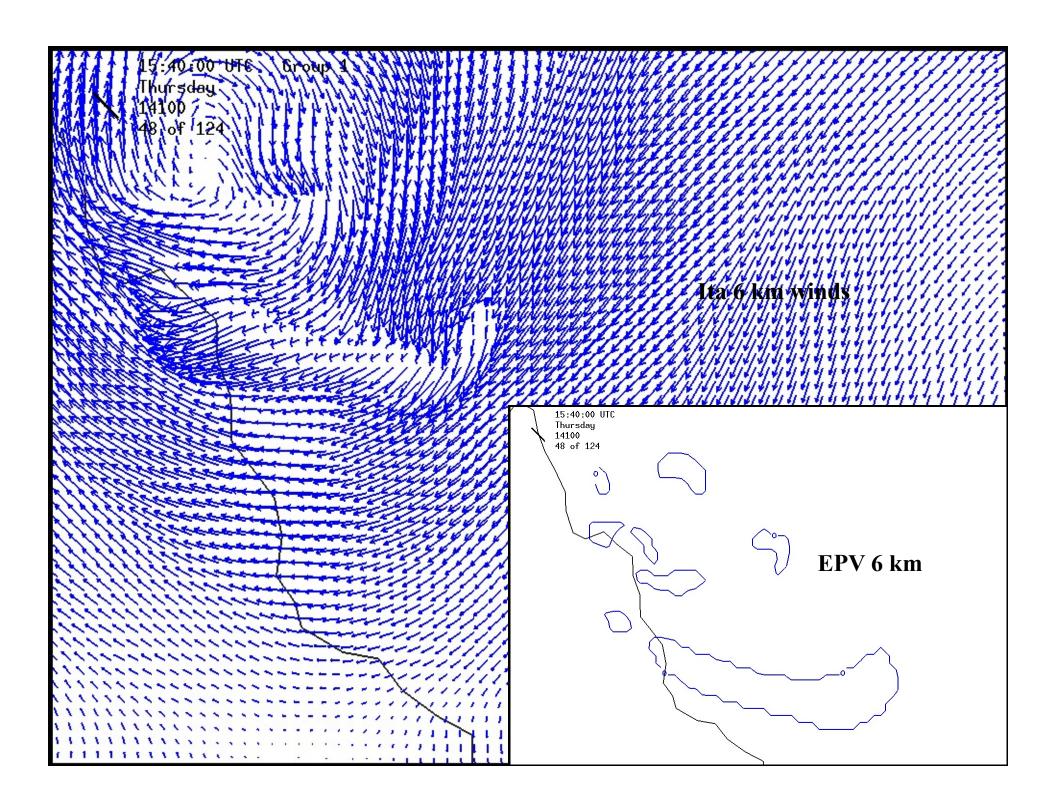


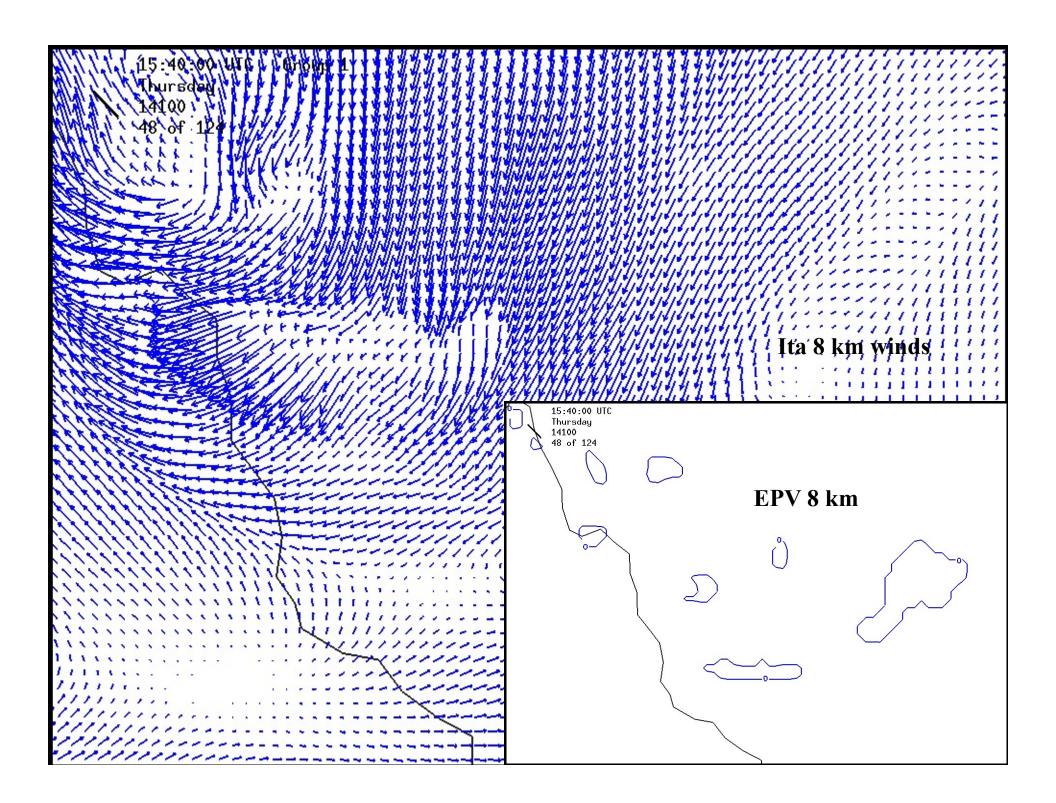
16.5 km winds; EPV

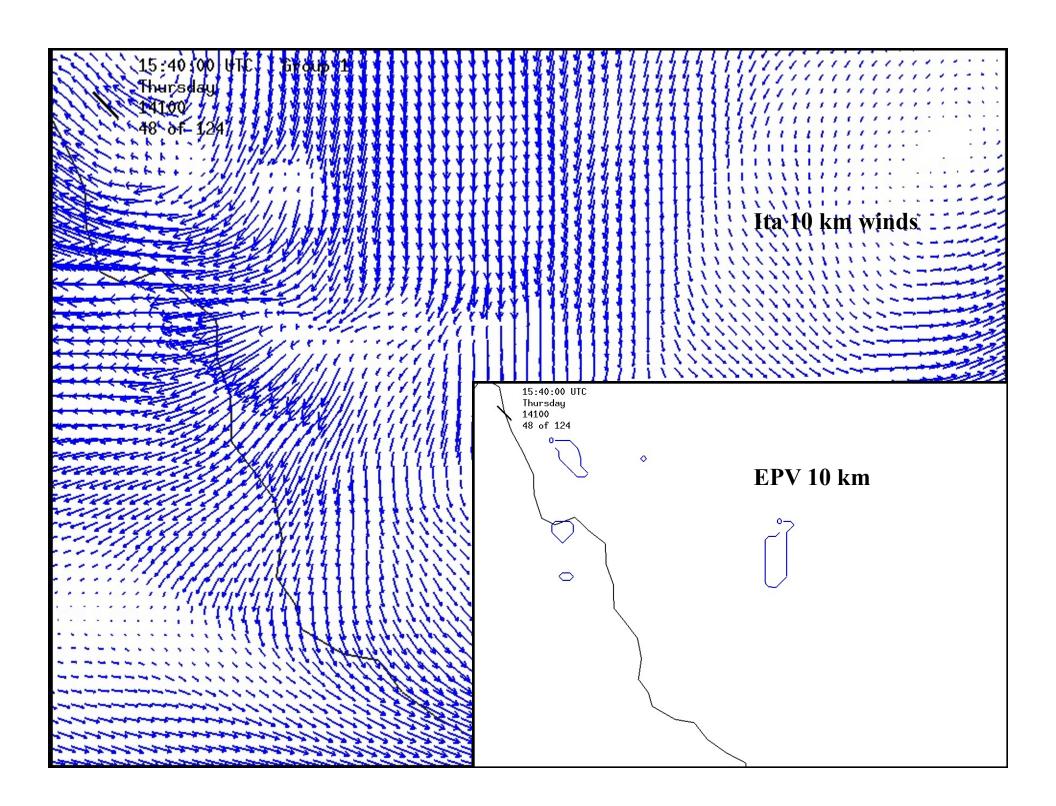


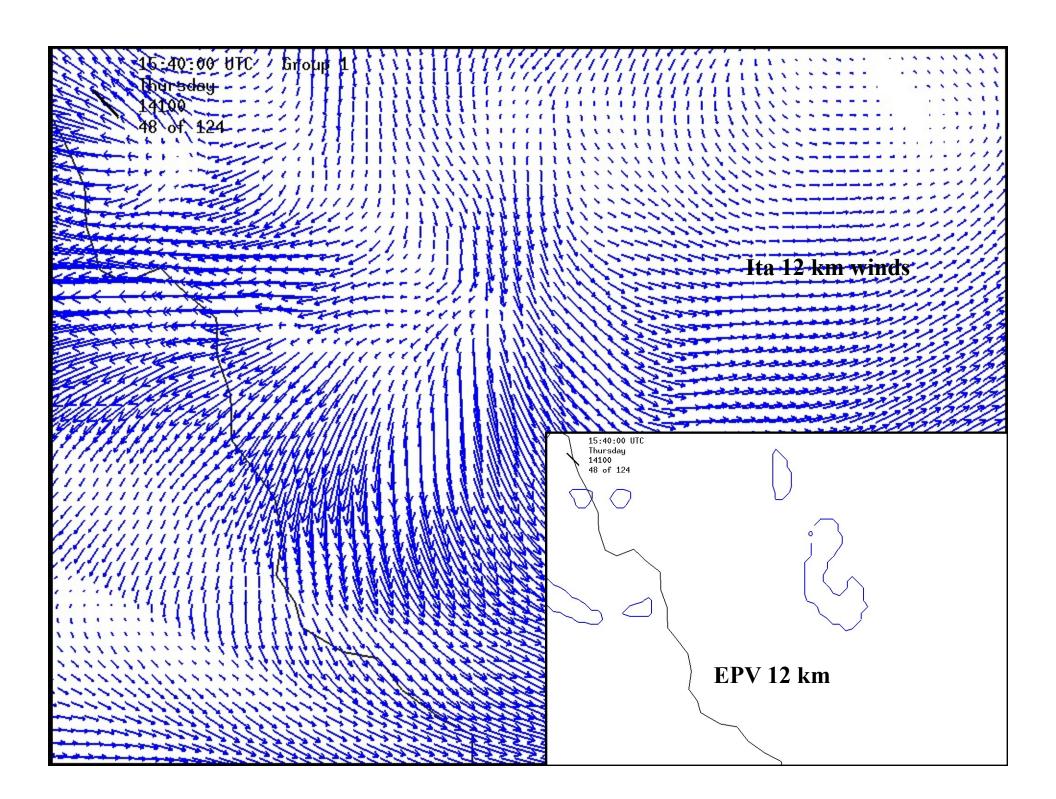


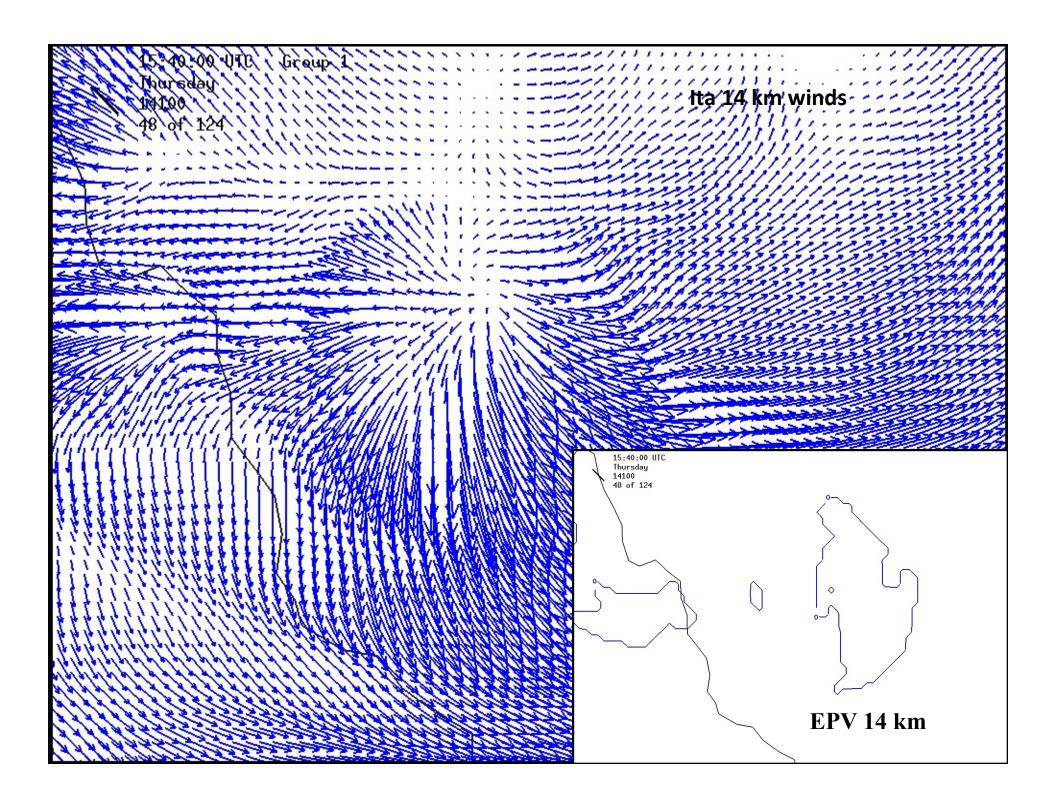


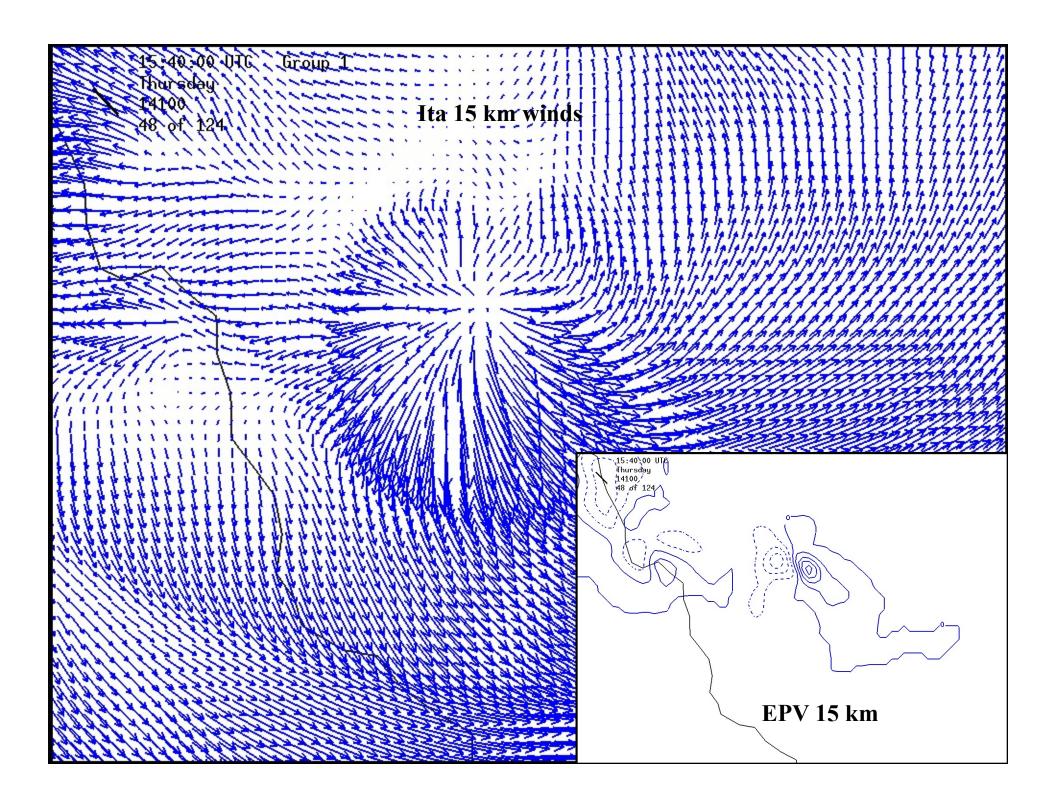


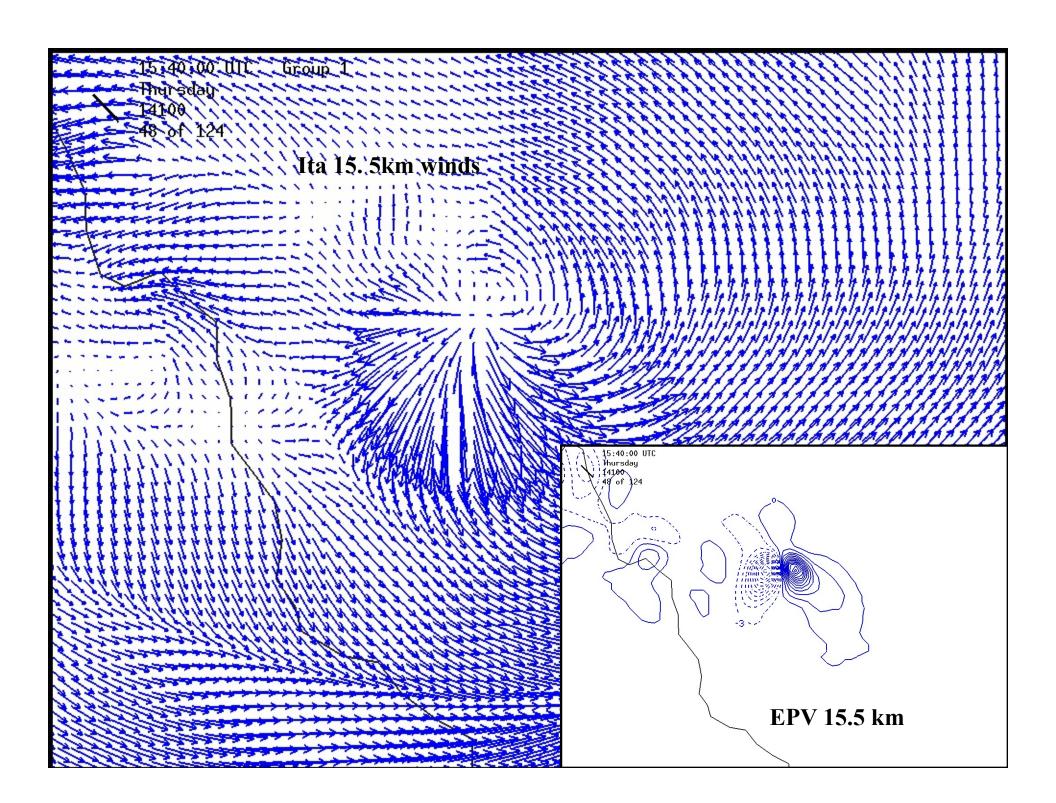


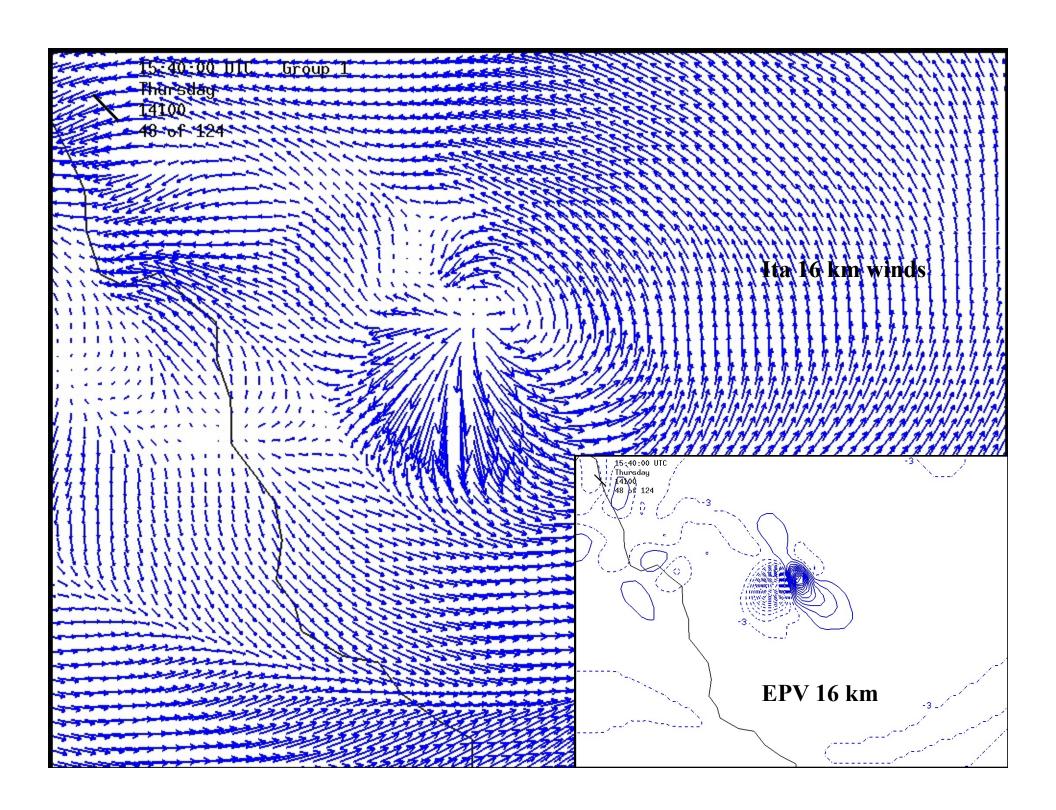


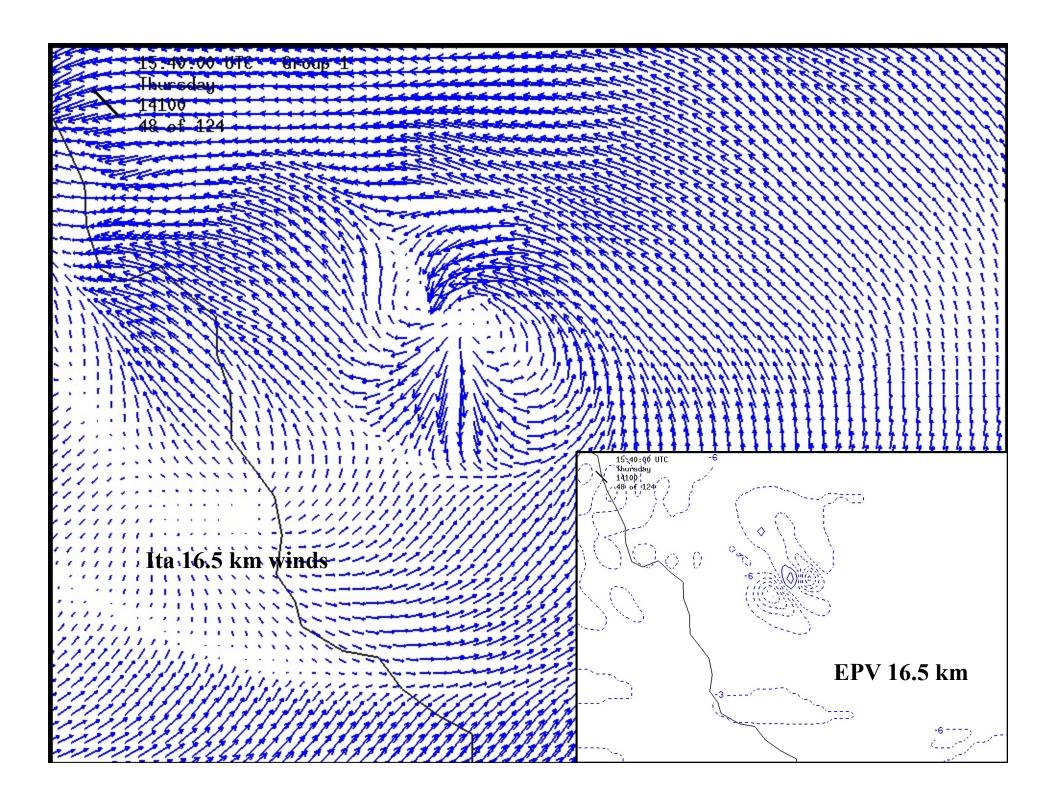












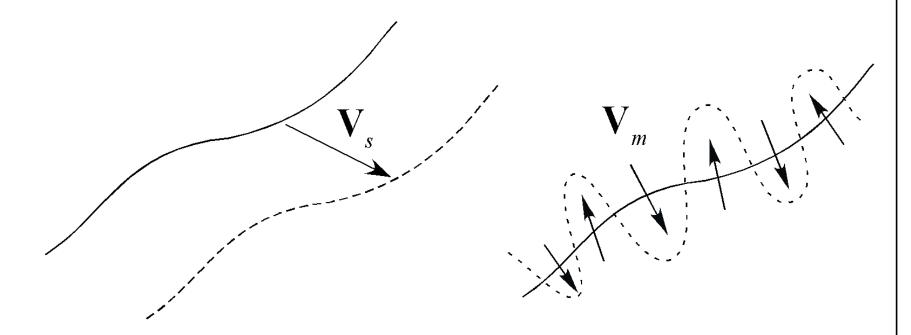
16.5 km winds – convection stirs the base of the stratosphere

- STE occurs on the periphery of cyclones and above convective complexes
- PV dipoles produced by convection, amplified in upper TS azimuthal shear
- Inertial instability *is* half of PV dipoles
- Inertial instability facilitates accelerations, outflow, STE
- Extratropical transition: creation of negative PV filament from convective scale injections

Future work:

- Quantify STE with 2-scale and annulus methods
- Explore drying effect
- Sort out relationship among w, θ_e dipoles and PV dipoles
- Simulate more convective complexes

The two-scale method



 V_s = smoothed winds advect material surface

$$V_m = V - V_s = resolved mixing winds$$

$$Flux = n_{\chi} \cdot V_{m,\perp}$$

