

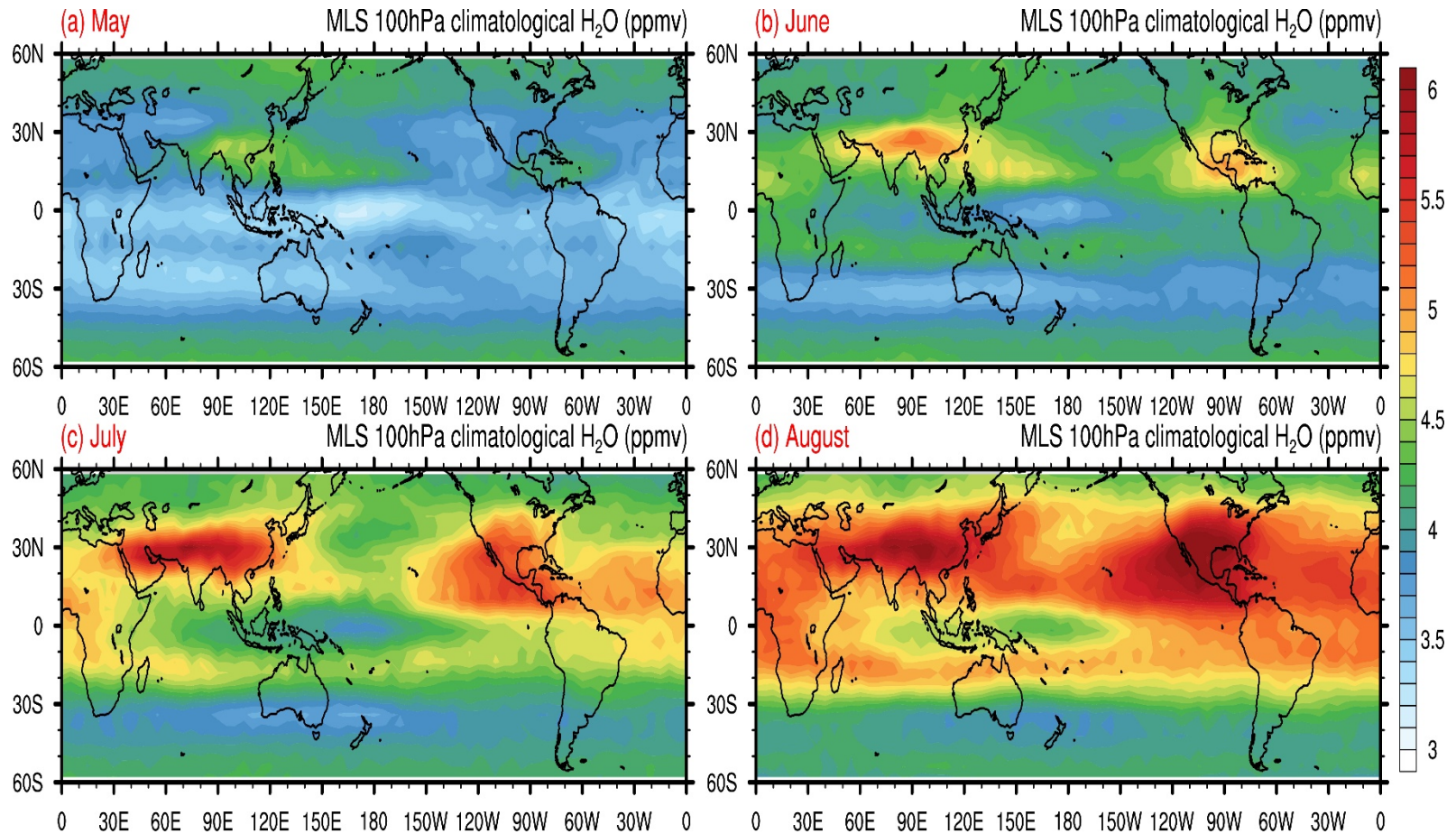
Influence of convective regime  
migration on water vapor transport in  
the Asian monsoon lower stratosphere

**Kai Zhang, Rong Fu**, The University of Texas at Austin

**William Randel**, NCAR

**Tao Wang**, JPL/Caltech

## Climatological 100 hPa H<sub>2</sub>O

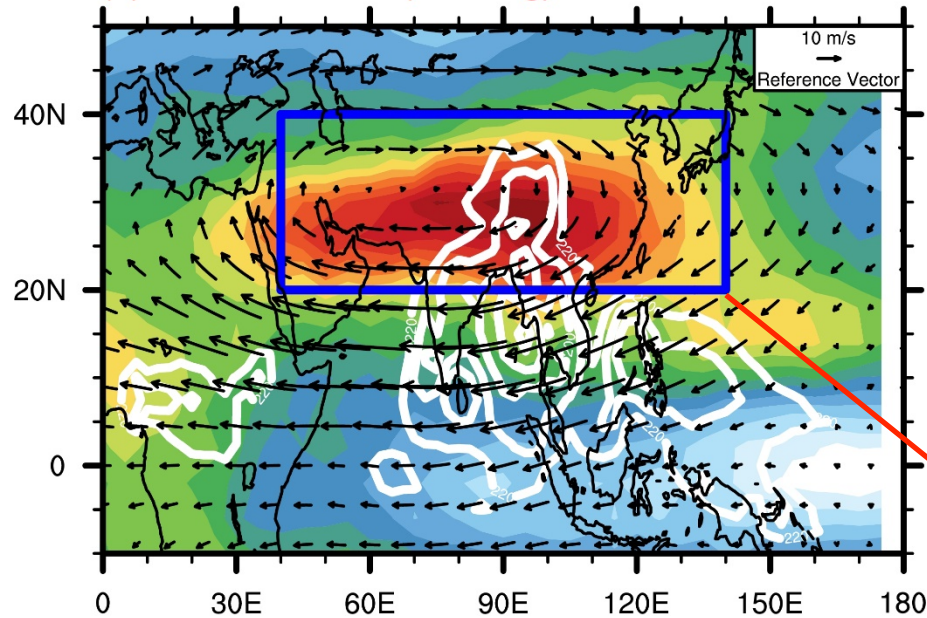


Q: What causes such a seasonal enhancement?

More convective injection?

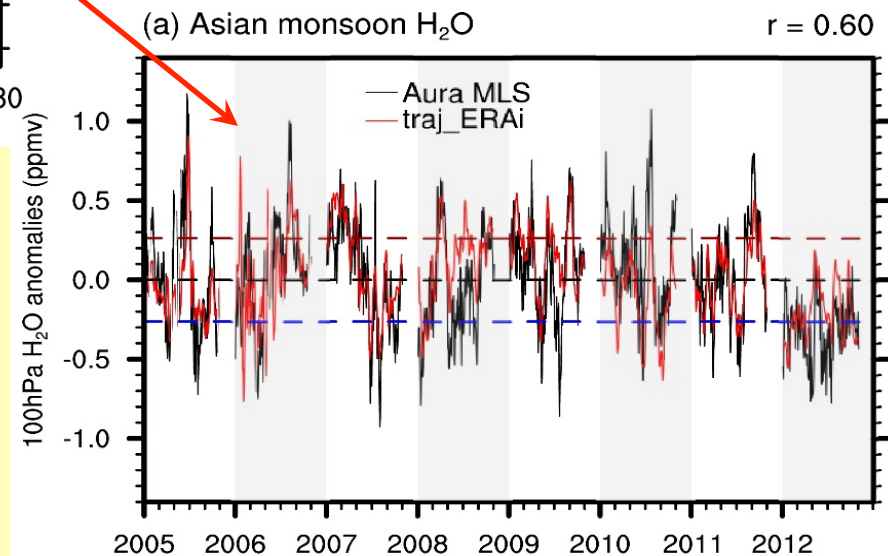
## Averaged 100 hPa H<sub>2</sub>O over the Asian monsoon region

(a) Asian monsoon (Jun-Aug)



- Aura MLS daily data (May-Sep, 2005-2013)  
- Domain-filling forward trajectory model

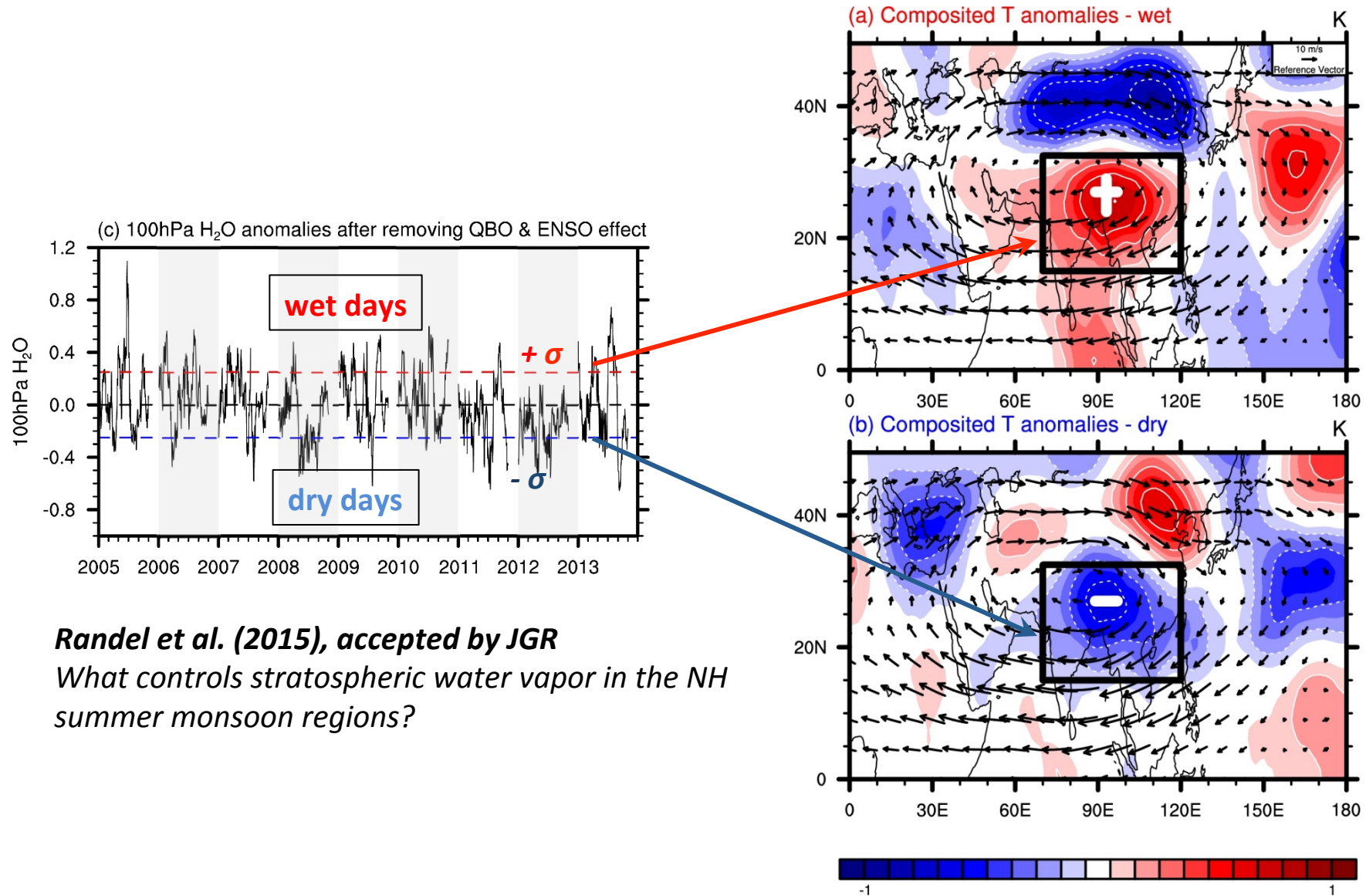
1. The good fit indicates that the variations of 100 hPa H<sub>2</sub>O are primarily controlled by temperature and large-scale circulation, and convective injection plays a minor role.
2. The traj can also well simulate the seasonal enhancement.



Q: Why is there such a seasonal enhancement?

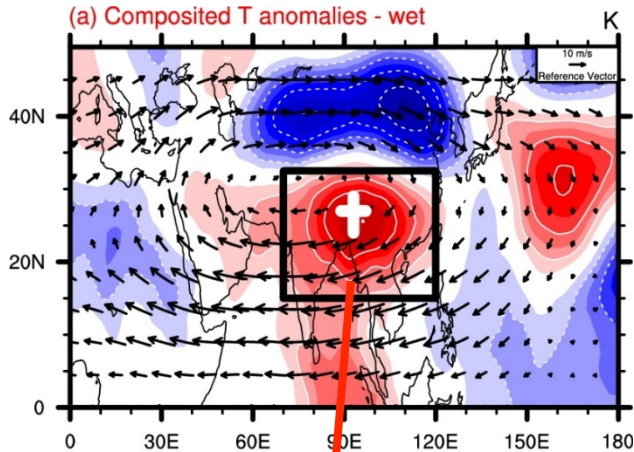
More convective injection 

# Variation of H<sub>2</sub>O and its correlation with 100 hPa temp (*Randel et al., 2015*)

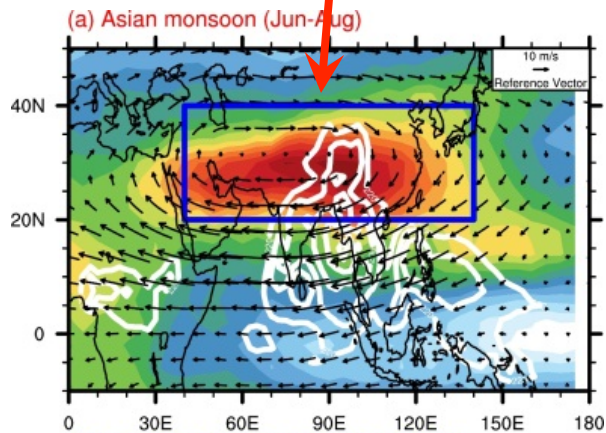


*Randel et al. (2015), accepted by JGR*

*What controls stratospheric water vapor in the NH summer monsoon regions?*



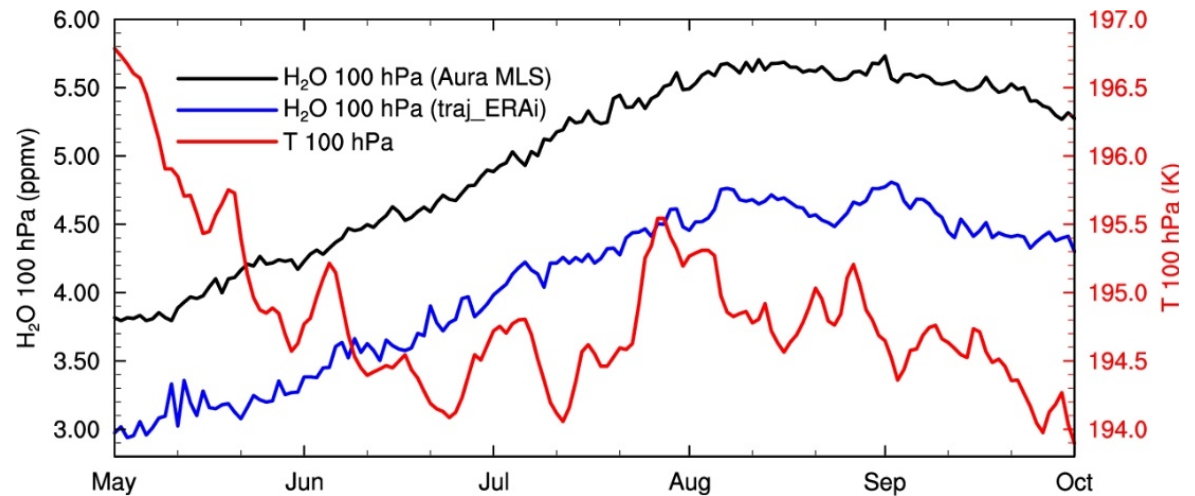
However, the temperature over this region cannot well explain the seasonal increase of LS water vapor in the Asian monsoon region.



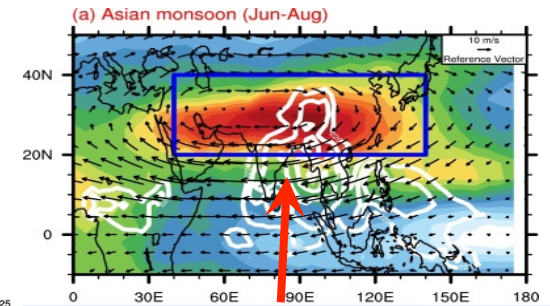
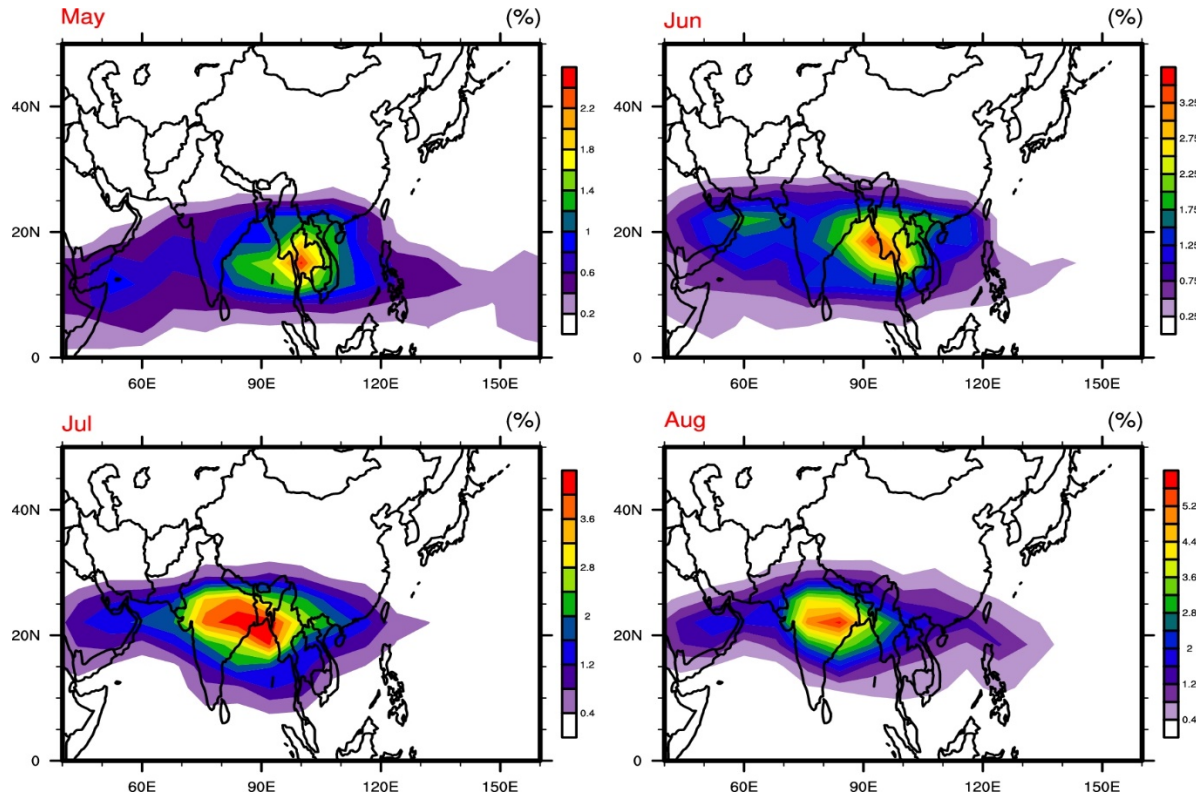
But the traj model can also well capture the increase, and the water vapor in the model is determined by temperature.

**How to explain the conflict?**

>> One possibility is that dehydration occurs over different regions.



## Westward migration of dehydration pattern



All the air parcels were selected at 100hPa over the Asian monsoon region during May-Sep 2005-2013.

In trajectory model, water vapor content is determined by the final dehydration temperature.

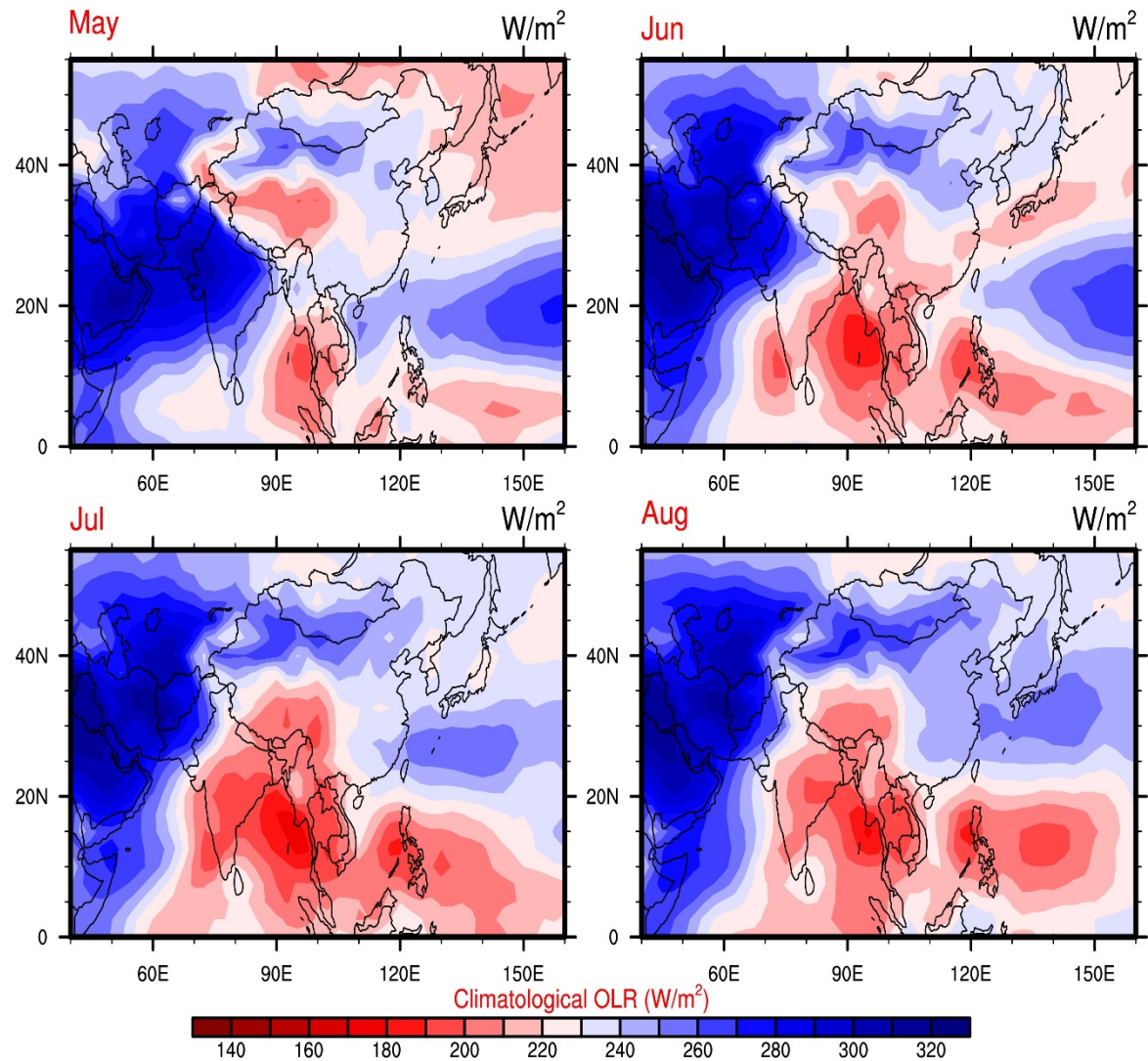
>> The dehydration pattern shows a westward migration, which could well explain the increase during the monsoon season.



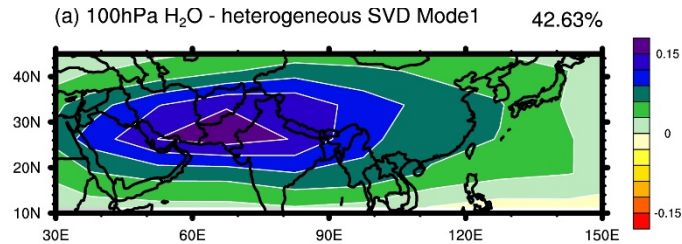
## Climatological OLR

Red shadings  
( $OLR < 220 \text{ W/m}^2$ )

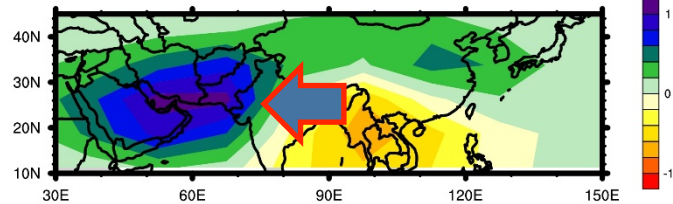
The convection pattern shows a westward migration during the monsoon season from May-July and retreat after August.



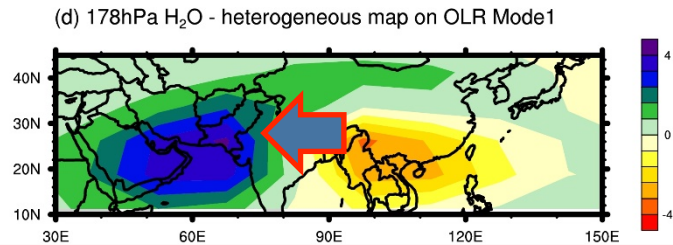
100hPa H<sub>2</sub>O



147hPa H<sub>2</sub>O

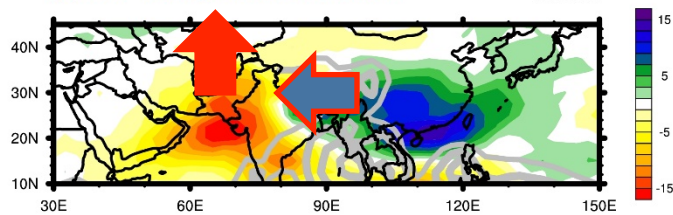


178hPa H<sub>2</sub>O

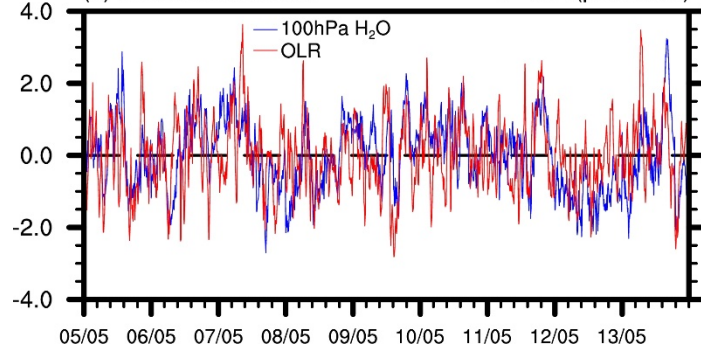


OLR

(d) OLR - homogeneous SVD Mode1 42.63%



(e) Time series Mode1  $r = 0.41 (p < 0.001)$

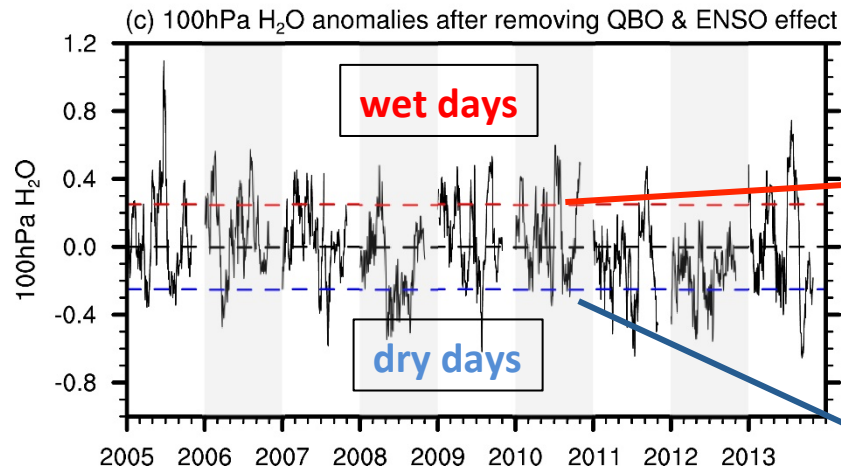


This figure shows the **SVD analysis** of 100hPa H<sub>2</sub>O anomalies and OLR using Aura MLS daily observations. There is a significant correlation between H<sub>2</sub>O and OLR.

>> Due to the **westward shift of convection**, humidity regime also shifts westward in the entire upper troposphere. As a result, more water vapor gets dehydrated and enters the LS over west side.

>> On the other hand, this convection pattern will lead to a warmer tropopause temperature on the east side, and thus more water vapor entering the LS over this region. (see *Randel et al., 2015, JGR*)

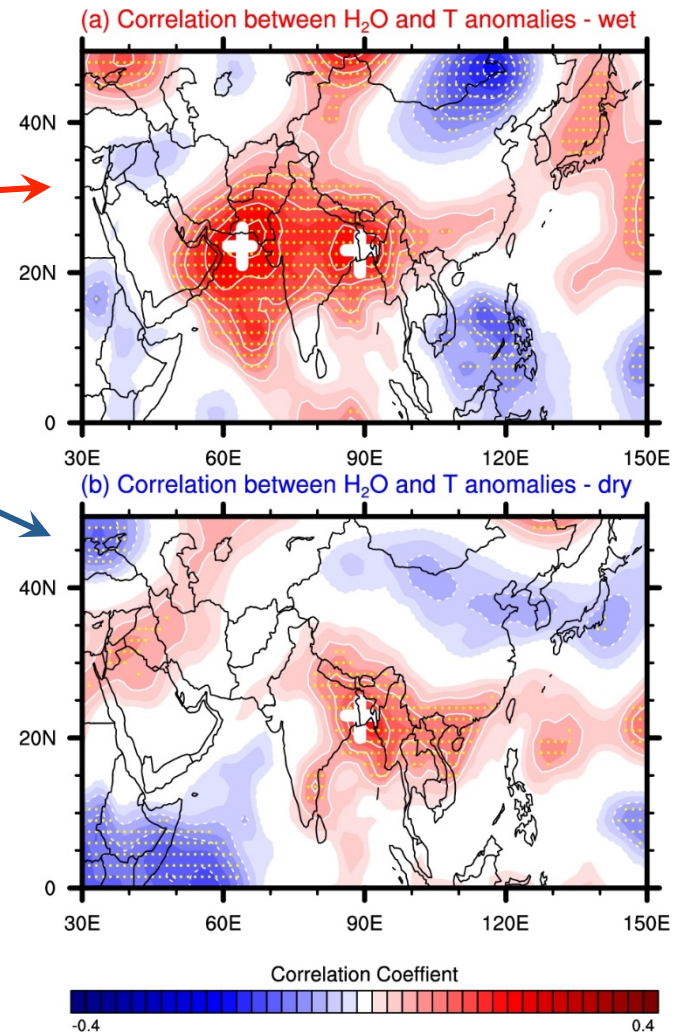
## Different dehydration pattern during wet and dry days: *Observations*



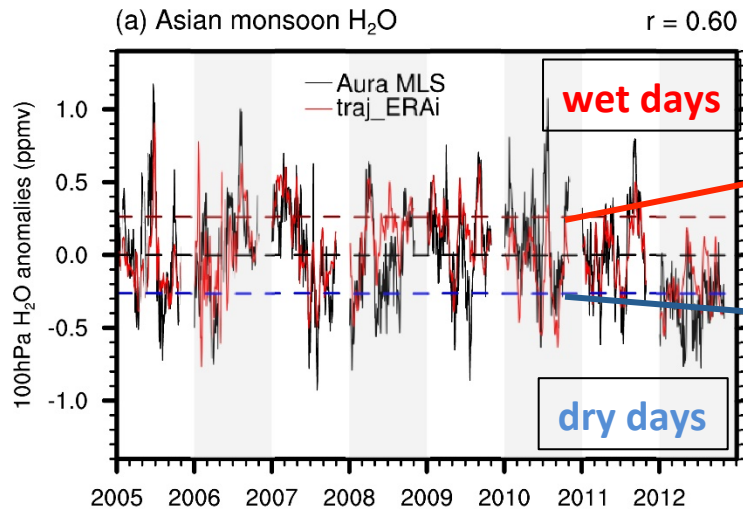
### *Hint:*

Positive correlation → temperature  
controlling region → dehydration region

During wet days, H<sub>2</sub>O variations are more correlated with temperature on the west side, which means that more H<sub>2</sub>O parcels are dehydrated by this region.



## Different dehydration pattern during wet and dry days: *Traj Model*

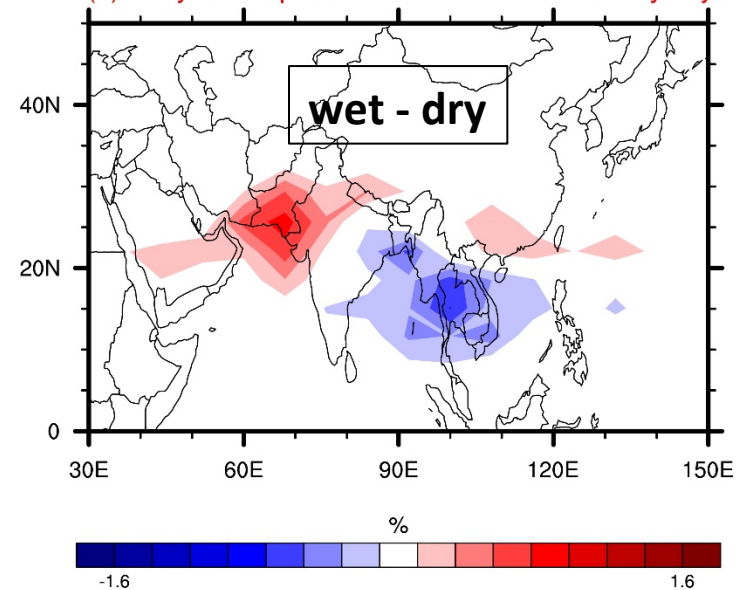


We trace back all the parcels in the Asian monsoon region at 100hPa from May-Sep 2005-2013.

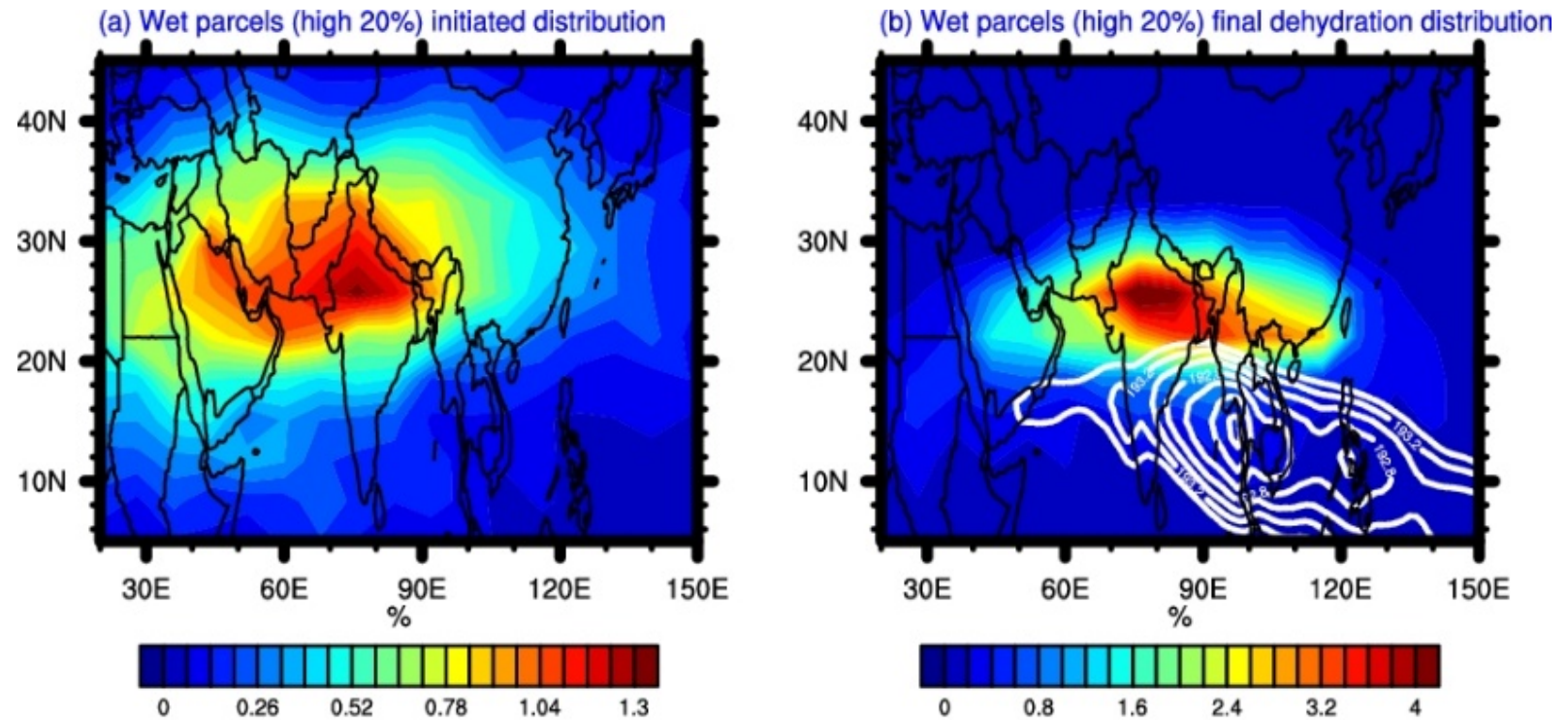
We can get the final dehydration locations for wet and dry days, and check their differences.

In trajectory model, during wet days, more air parcels are finally dehydrated over the west side, which is consistent with the observational results.

(b) Dehydration preference between wet and dry days

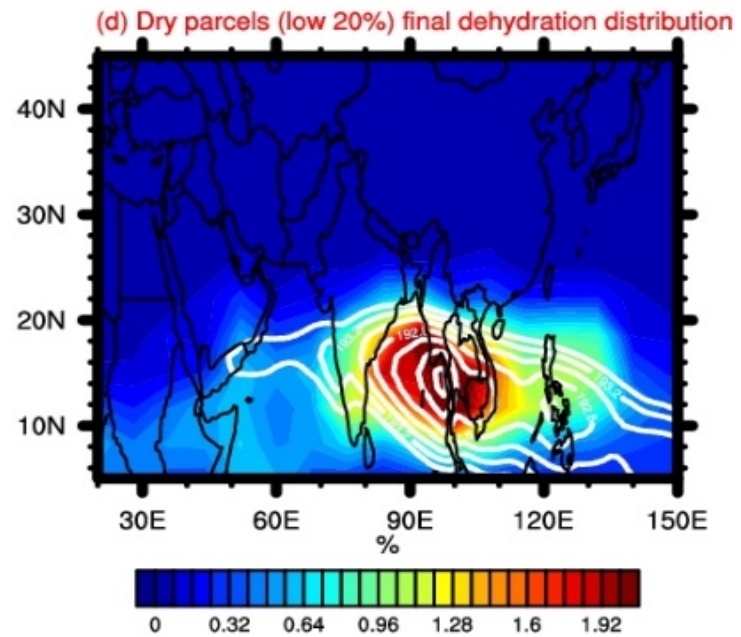
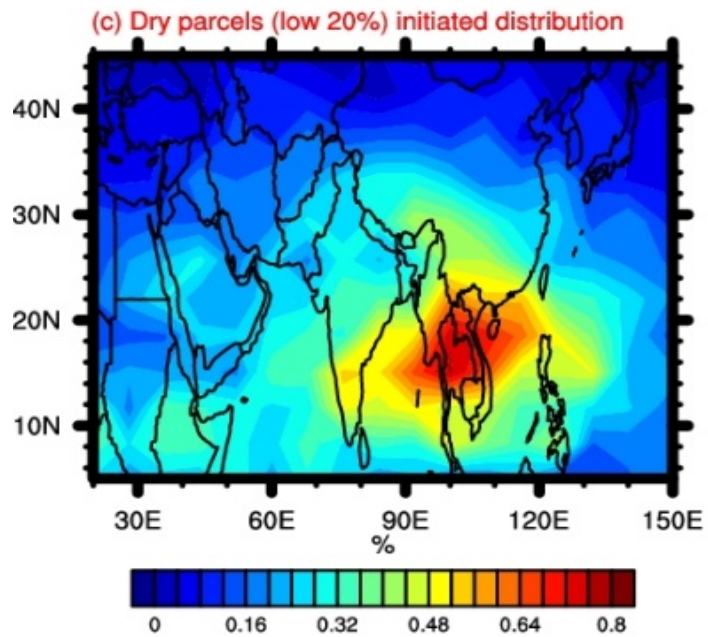


## Wet parcels (*high 20% percentile*) in traj model



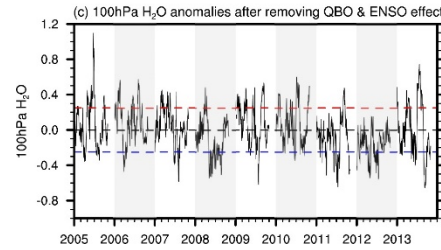
All the air parcels were selected at 100hPa over the Asian monsoon region during May-Sep 2005-2013.

Dry parcels (*low 20% percentile*) in traj model



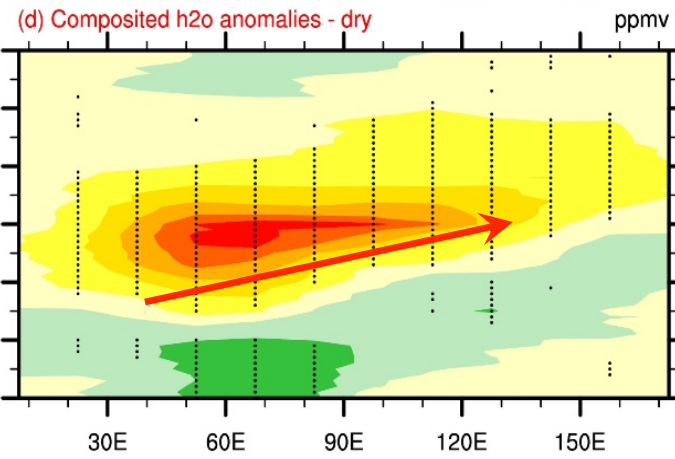
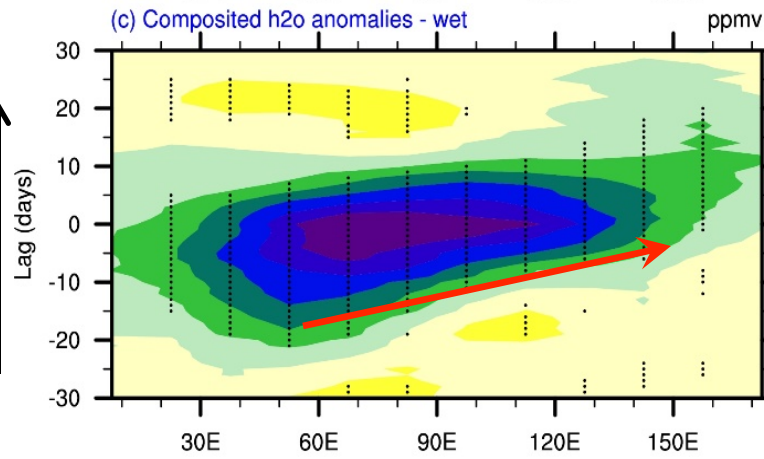
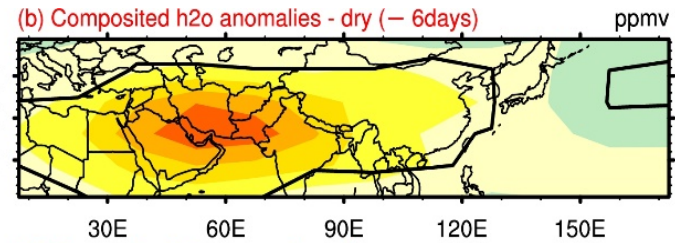
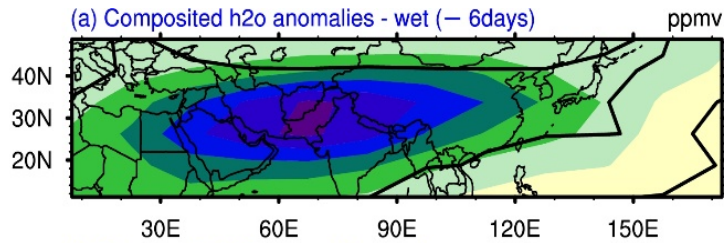
# How do water vapor anomalies transport in the LS over the Asian monsoon region?

## Composite H<sub>2</sub>O anomalies using Aura MLS observations

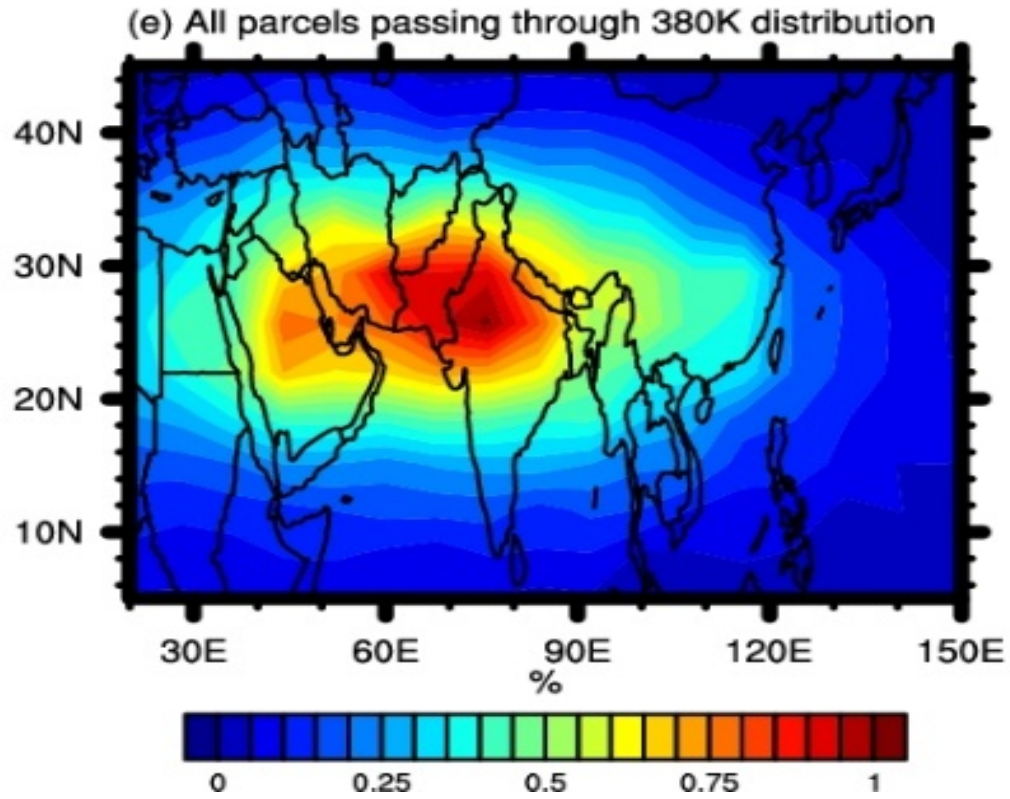


**Wet days**

**Dry days**

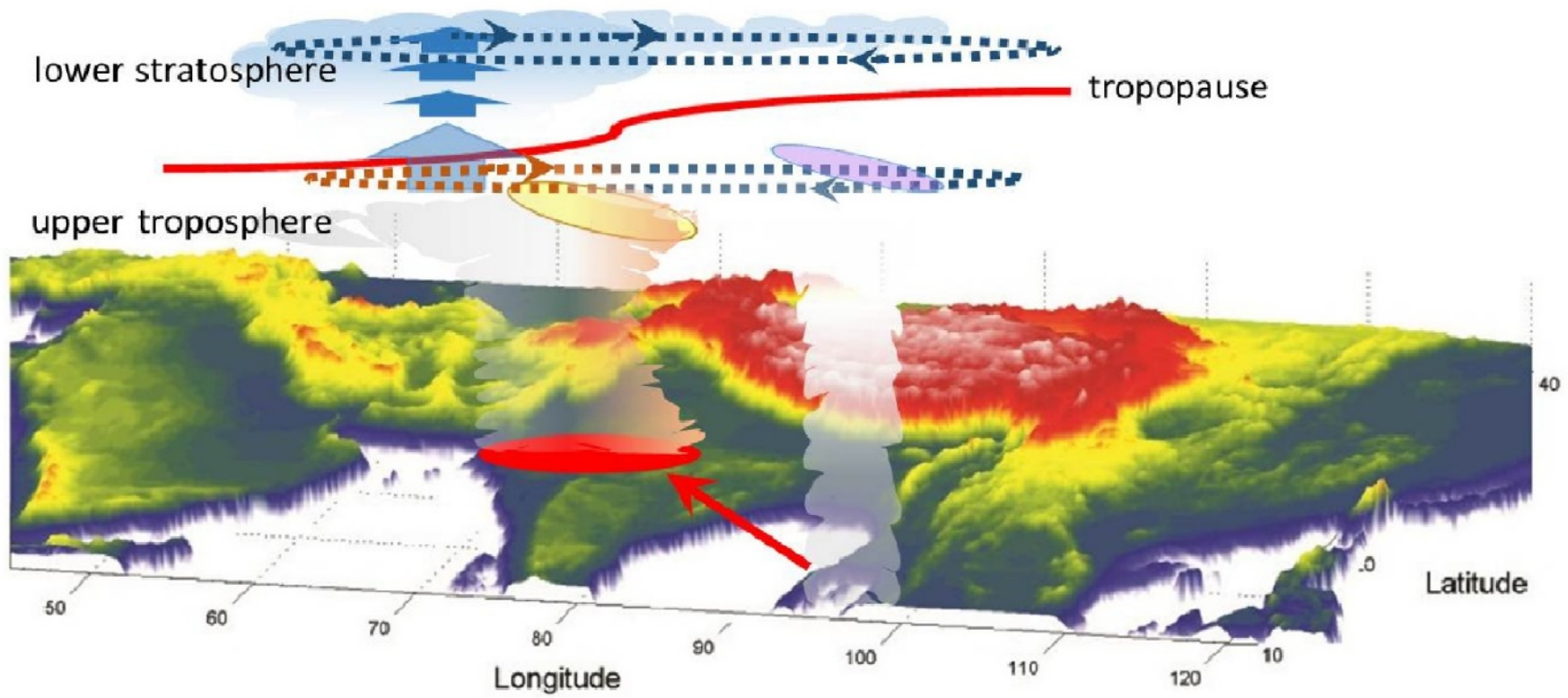


**Where did all parcels pass through tropopause?**

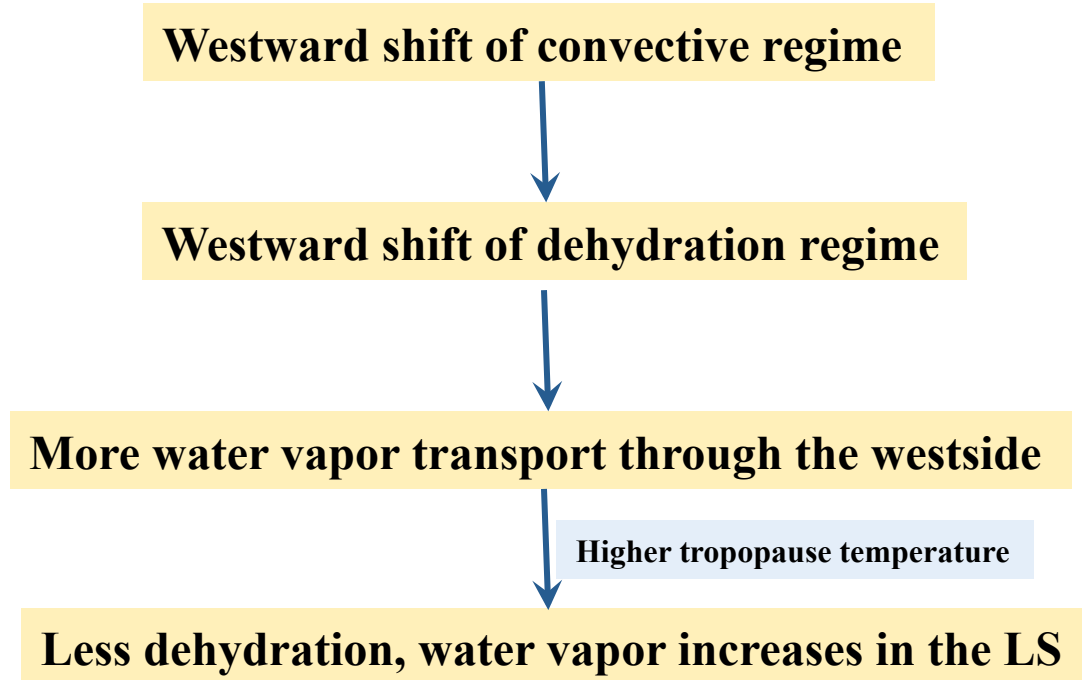




# Conclusion



## Conclusion



1. Water vapor variations over the Asian monsoon LS are primarily controlled by temperatures;
2. A westward migration of convection will moist the LS.

Thanks!

An example of 2011

