

Convective transport of chemical compounds from the surface to the UT/LS: The western Pacific vs. North America

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&

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Motivation and Goals

Surface to the UT/LS transport of trace gases can exert significant impacts on atmospheric abundances of ozone (O_3) and hydroxyl radical (OH), which will have chemical and radiative impacts on the atmosphere.

- Vertical transport of CO and NMHCs
- Vertical transport of very-short-lived (VSL) halogen compounds, i.e. $CHBr_3$ and CH_2Br_2 .

Questions of interest

- Tropical and subtropical convective transport:
 - Frequency
 - The timescale it takes to ascent into the UT/LS
 - The regional differences in convective lofting: Western Pacific vs. North America
- What are transport efficiency of chemical compounds (the impact of tracer lifetime and solubility)?

What are our best choice of chemical tracers?

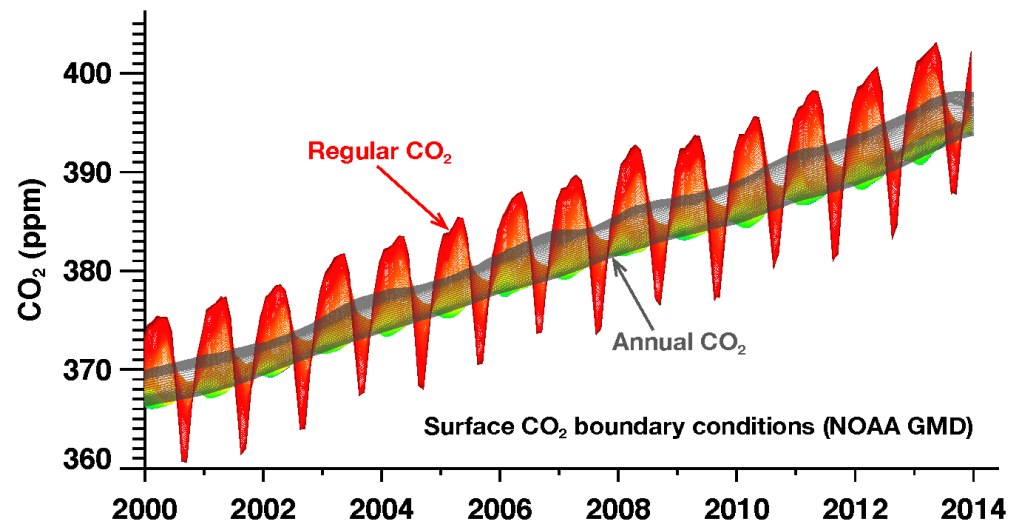
-- CO? VOCs and their ratios? Tracers of marine origin? CO_2 , SF_6 ...

- V2 - GEOS-5 with standard stratospheric chemistry and detailed VSL bromocarbons
- GCM simulations for the ATTREX/CONTRAST and SEAC⁴RS period
 - Horizontal resolution is 2° latitude x 2.5° longitude
 - 72 vertical layers from surface to 0.01 hPa, with ~1 km/layer resolution at the UT/LS region
(will conduct a future simulation with MERRA meteorology further comparison)

Two new CO₂ tracers

– transport clock as well as model transport diagnostics

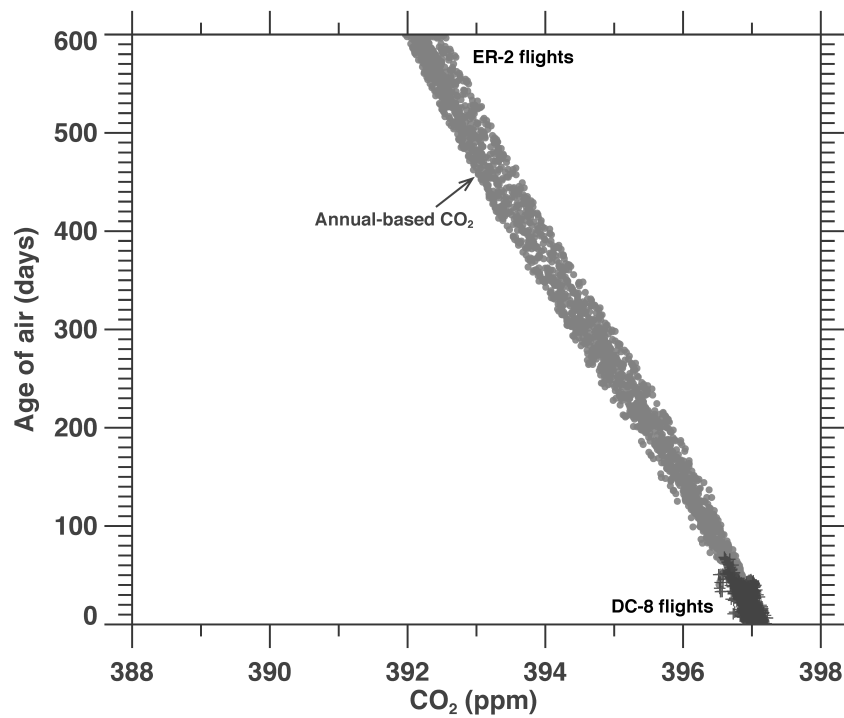
- 1) Regular CO₂ - with surface boundary conditions from the NOAA observations (latitude and monthly variable)
- 2) Annual CO₂ – same as regular CO₂ but with seasonal cycle removed at all latitude grids.



The model simulations are sampled at the same geographical and vertical location for the corresponding time as the observations for model-observation comparison.

SEAC⁴RS

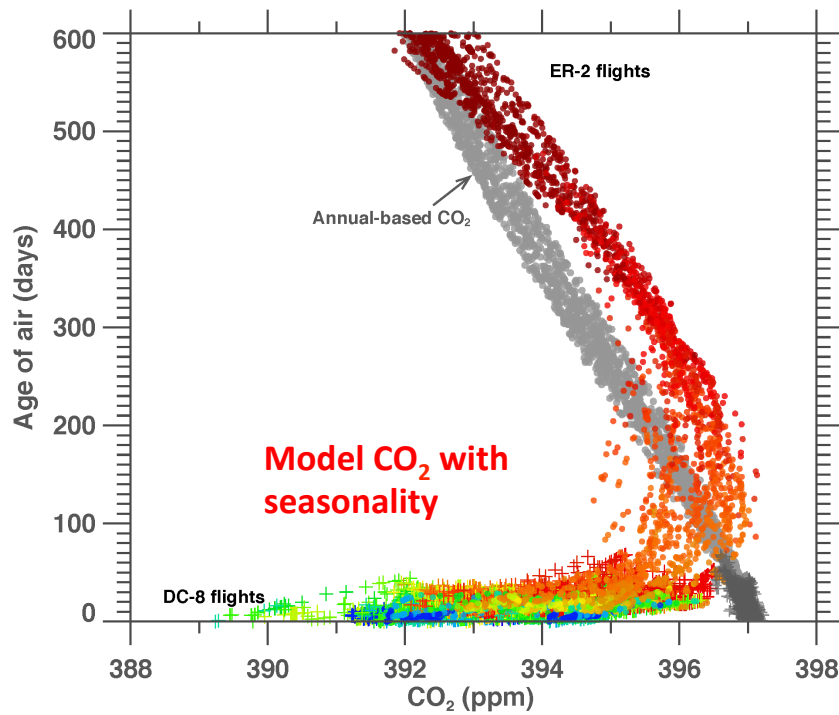
(N. America, Aug-Sep 2013)



Model annual-based CO₂ shows a linear correlation with age of air, thus an accurate clock tracer for transport

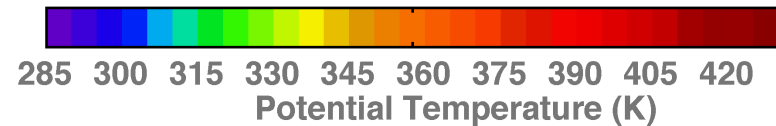
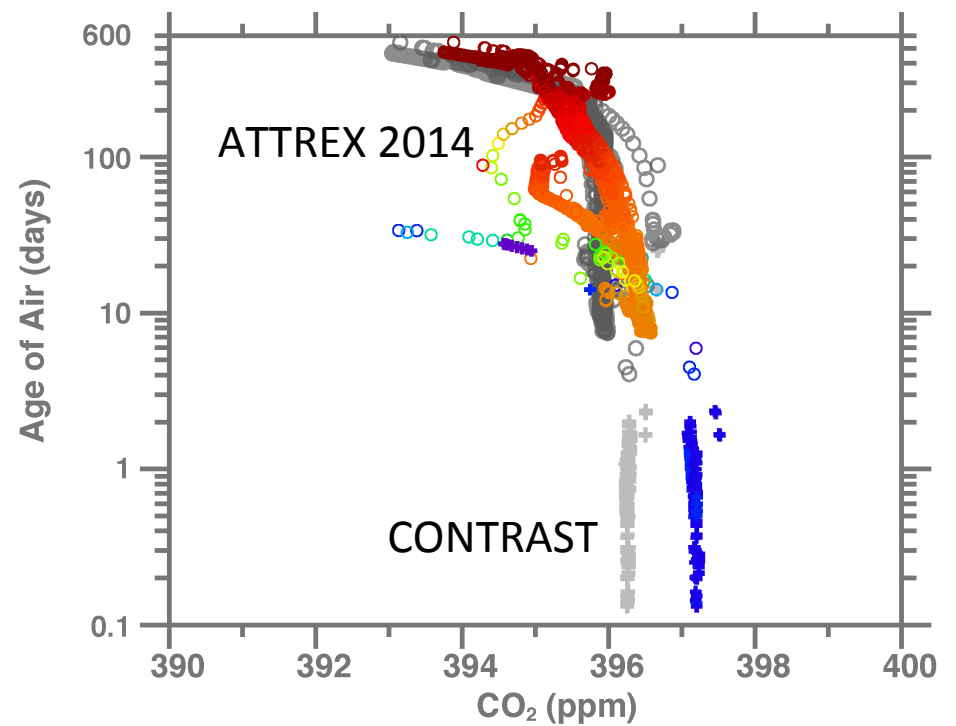
SEAC⁴RS

(N. America, Aug-Sep 2013)



ATTREX-3/CONTRAST

(W. Pacific, Jan-Feb 2014)



However, the actual application as a clock tracer is complicated due to the surface seasonal variations in CO₂



Two-step approach on transport analysis

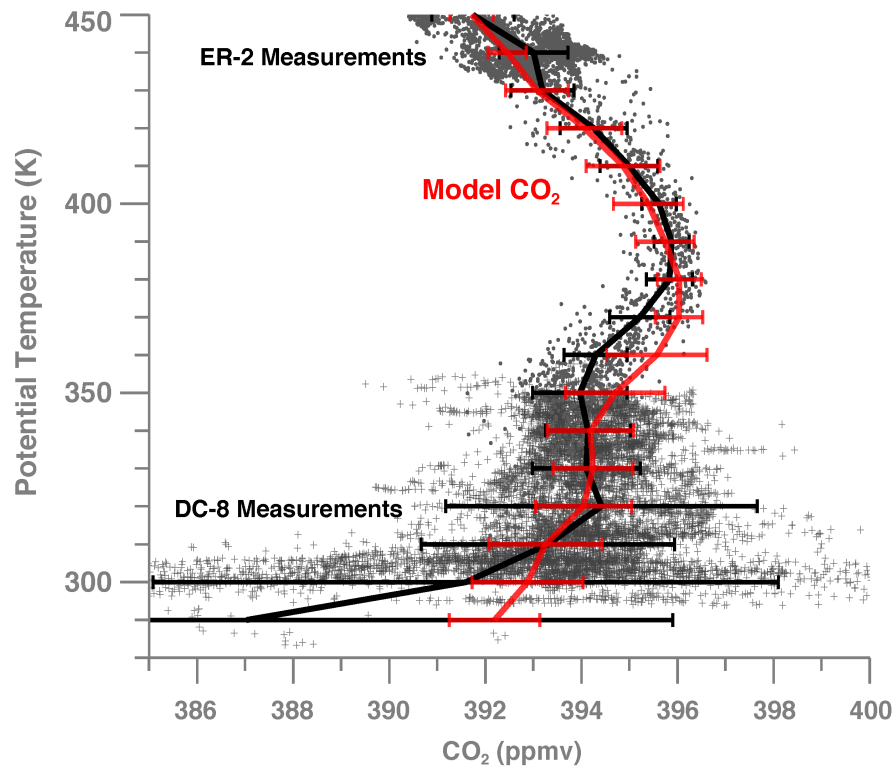
Our proposed 2-step approach:

Step 1. Use model CO_2 vs. observed CO_2 to evaluate model transport: convection and large scale ascent

Step 2. Use the model age of air to quantify transport timescale

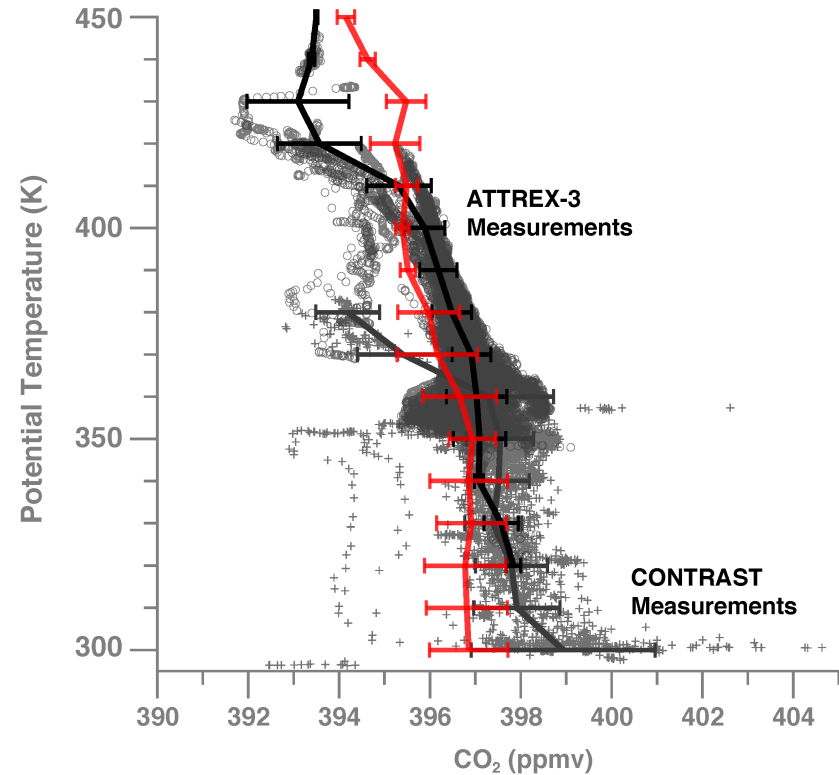
SEAC⁴RS

(N. America, Aug-Sep 2013)



ATTREX-3/CONTRAST

(W. Pacific, Jan-Feb 2014)

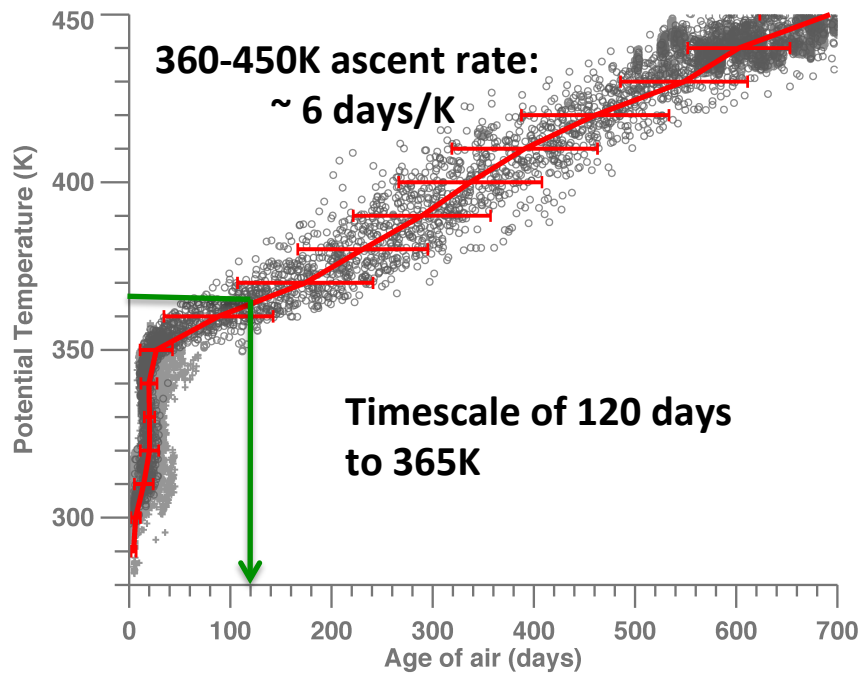


GEOSSCM modeled CO₂ agrees well with the SEAC⁴RS and ATTREX-3/CONTRAST observations, indicating credible model transport in the vertical.

Step 2: Quantify vertical transport Timescale using model age of air

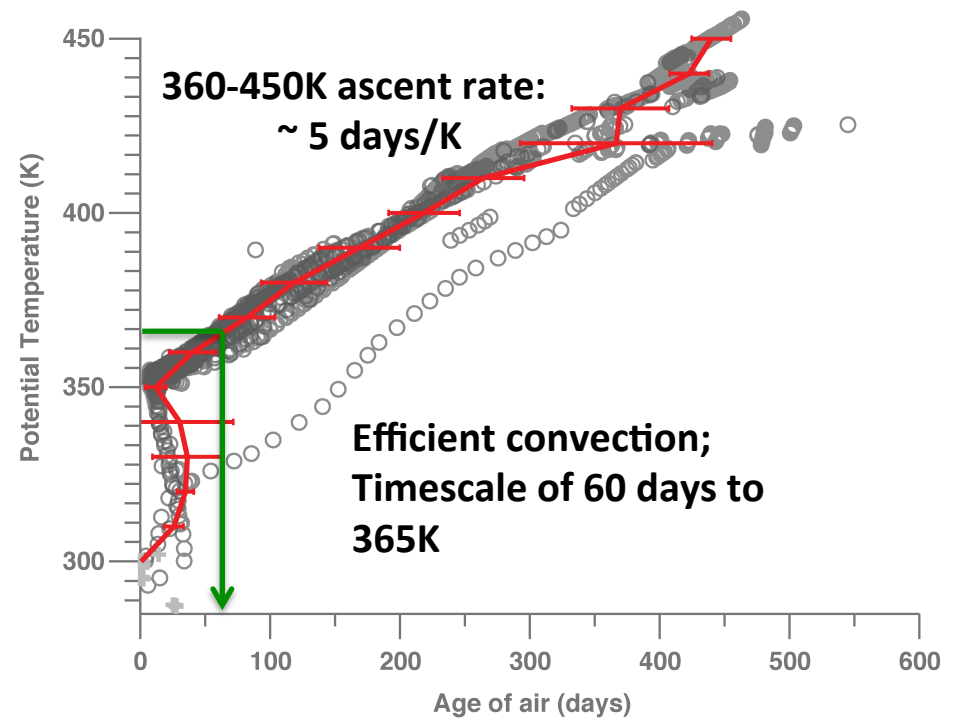
SEAC⁴RS

(N. America, Aug-Sep 2013)



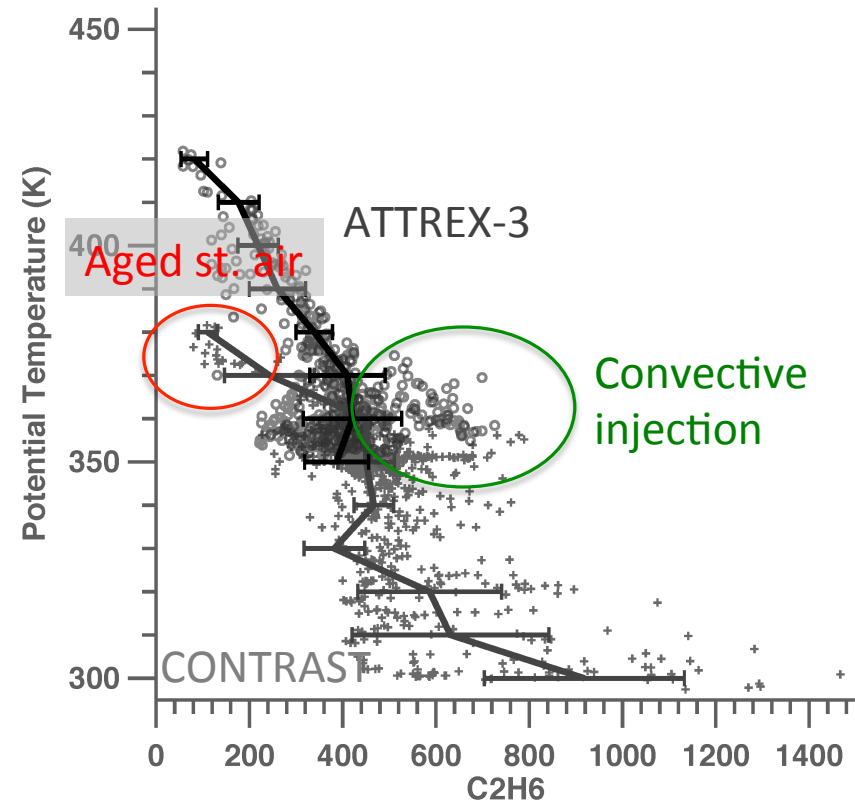
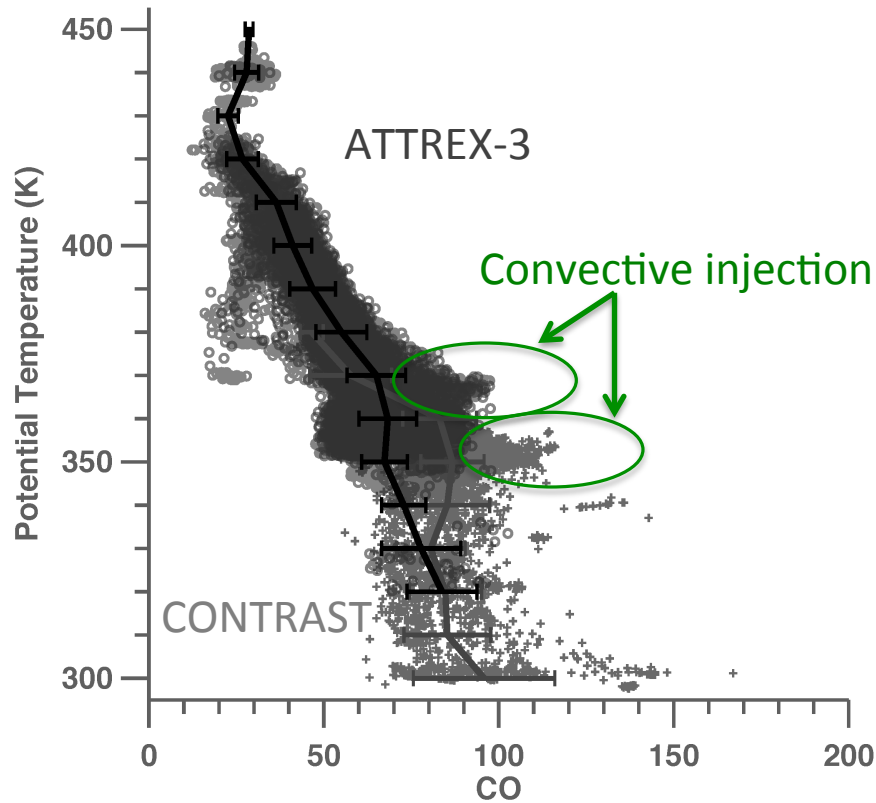
ATTREX-3/CONTRAST

(W. Pacific, Jan-Feb 2014)



Surface to UT/LS transport of pollution tracers:

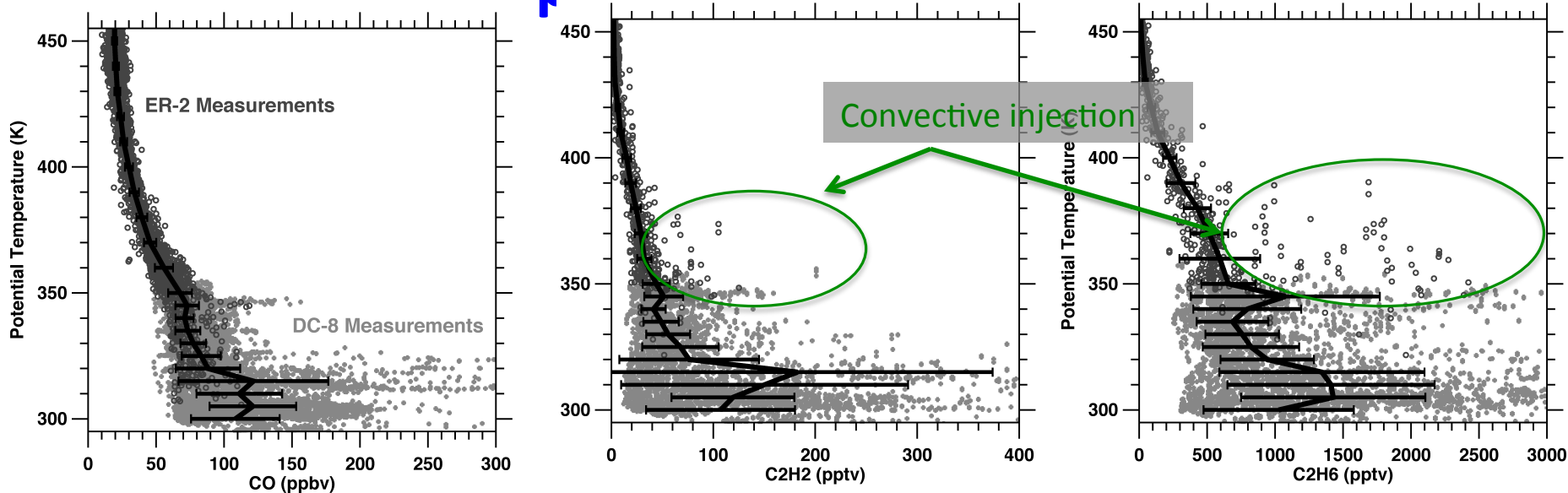
ATTREX-3 and CONTRAST



For the deep convection active western Pacific, CO and C₂H₆ work well as convection tracers.

Surface to UT/LS transport of pollution tracers:

Subtropical N. America: SEAC4DC



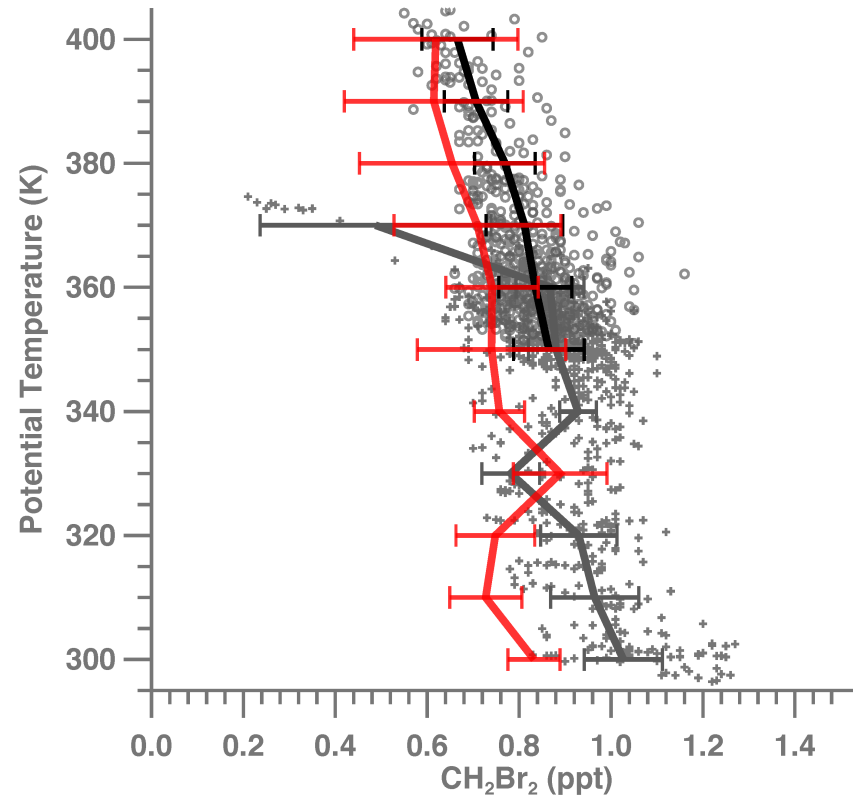
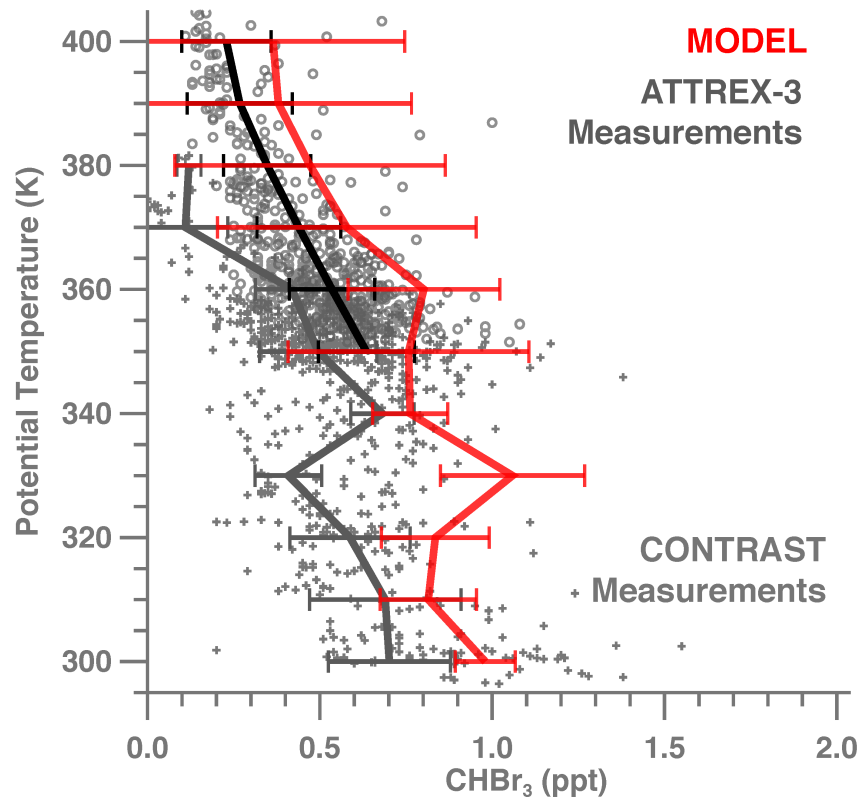
Tracer lifetime
At different
layers

	CO	C ₂ H ₂	C ₂ H ₆
Loss reactions	CO + OH ->	C ₂ H ₂ + OH -> C ₂ H ₂ + O ₃ ->	C ₂ H ₆ + OH -> C ₂ H ₅ --> CH ₃ CHO --> PAN C ₂ H ₆ + Cl -> C ₂ H ₅ --> CH ₃ CHO --> PAN
400-450K	~ 2 years	~ 6 months	~ 4 months
350-400K	~ 1 year	~ 10 months	~ 10 months
Troposphere	~ 1 month	~ 0.5 month	~ 3 months

Efficient C₂H₆ loss via the C₂H₆ + Cl reaction creates a large vertical as well as tropical vs. high-latitude gradients in C₂H₆ -> making C₂H₆ an excellent transport tracer for tropical convective lofting.

Surface to UT/LS transport of CHBr_3 and CH_2Br_2

Western Pacific: ATTREX-3 & CONTRAST



Model CHBr_3 are biased high wrt observations, but model CH_2Br_2 are biased low.

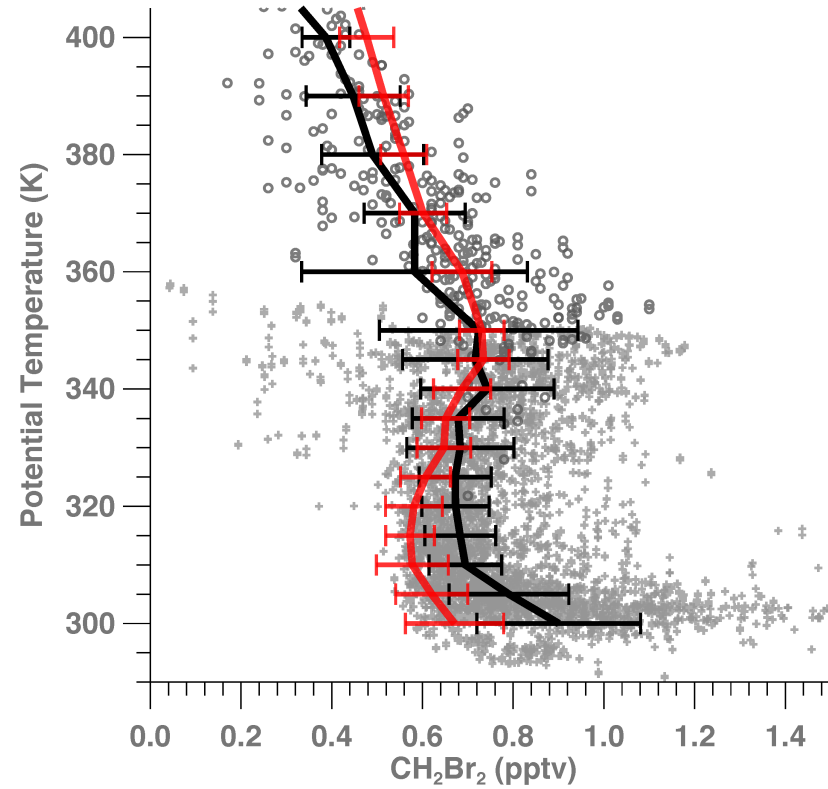
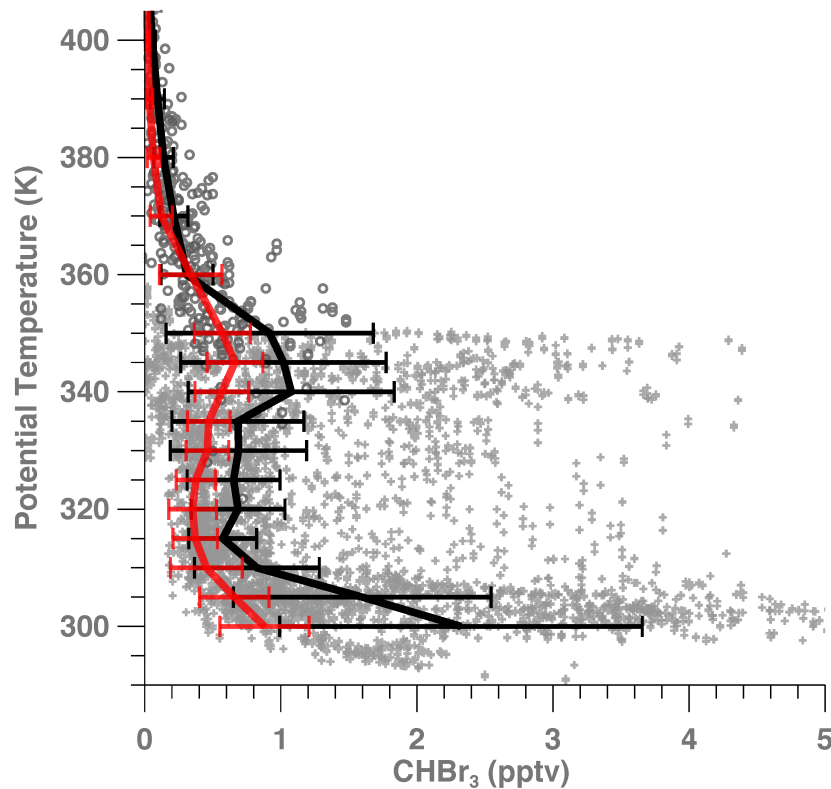
This is consistent with findings from the SHEVA measurements (Hossaini et al. 2013) (GEOSCCM emissions are fine, slightly high-biased). However, why?

Surface to UT/LS transport of CHBr_3 and CH_2Br_2

North America: SEAC4RS

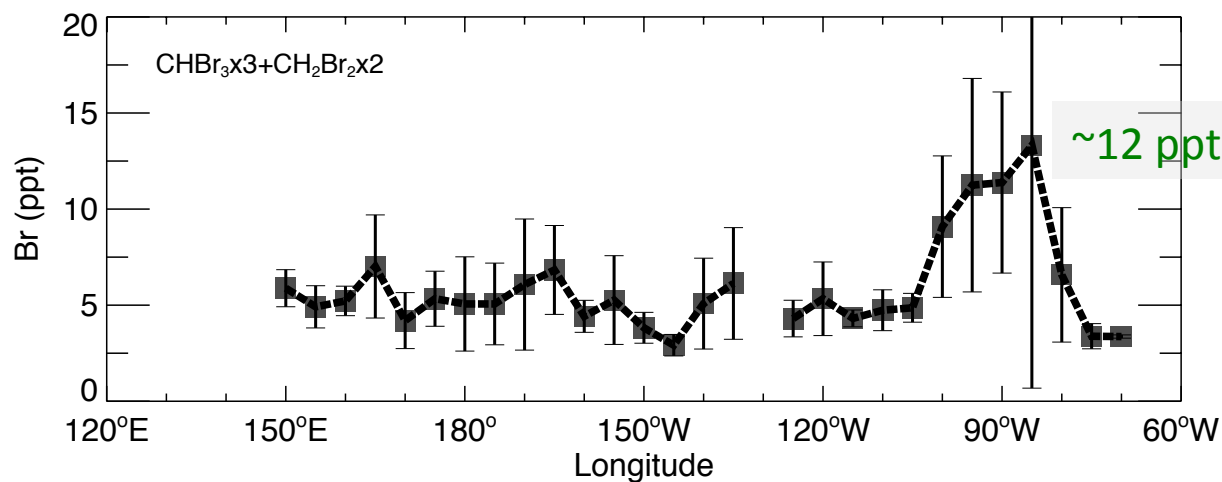
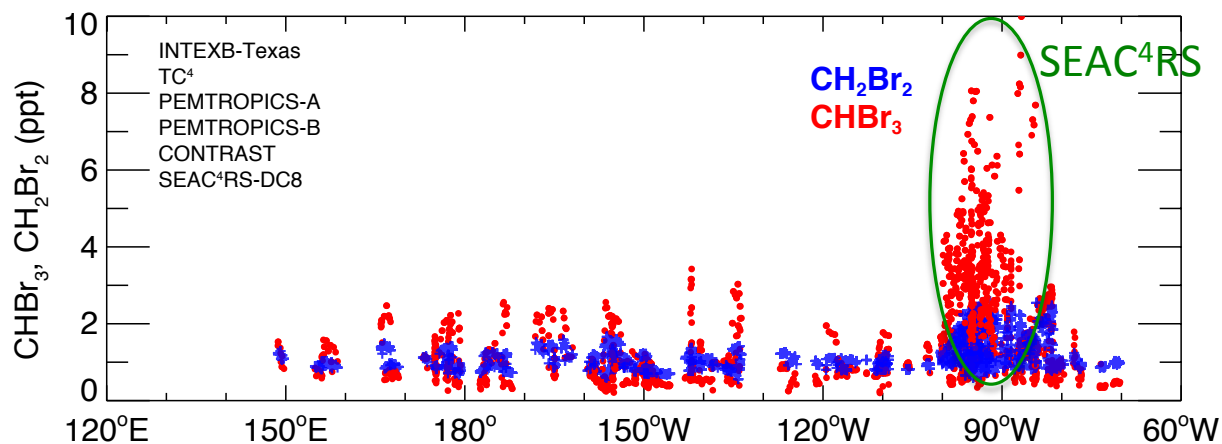
CHBr_3

CH_2Br_2

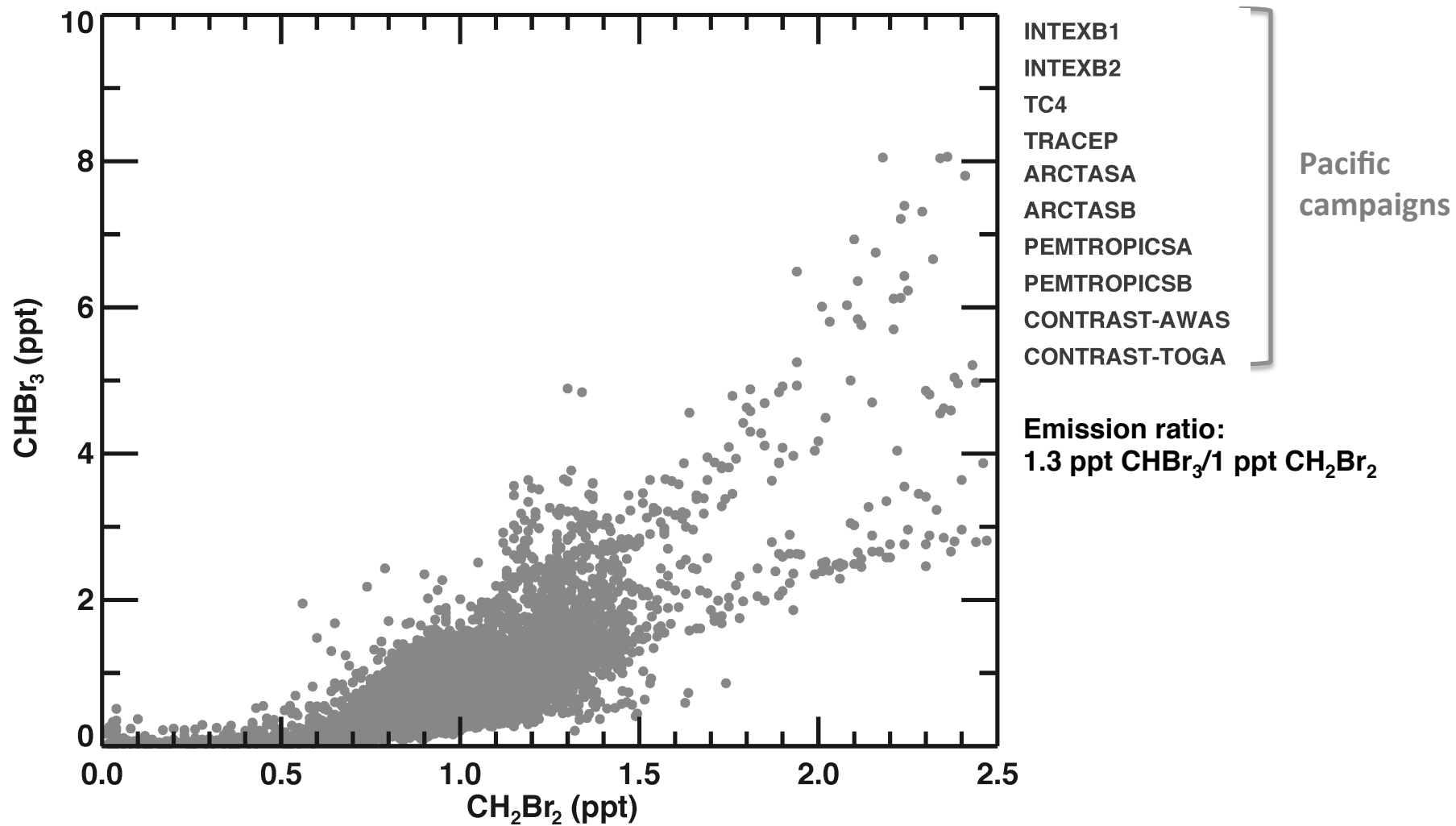


Tropospheric low biases -> Surface emissions of CHBr_3 and CH_2Br_2 are too small.
This contradicts the findings from CONTRAST and Hossaini et al. (2013).

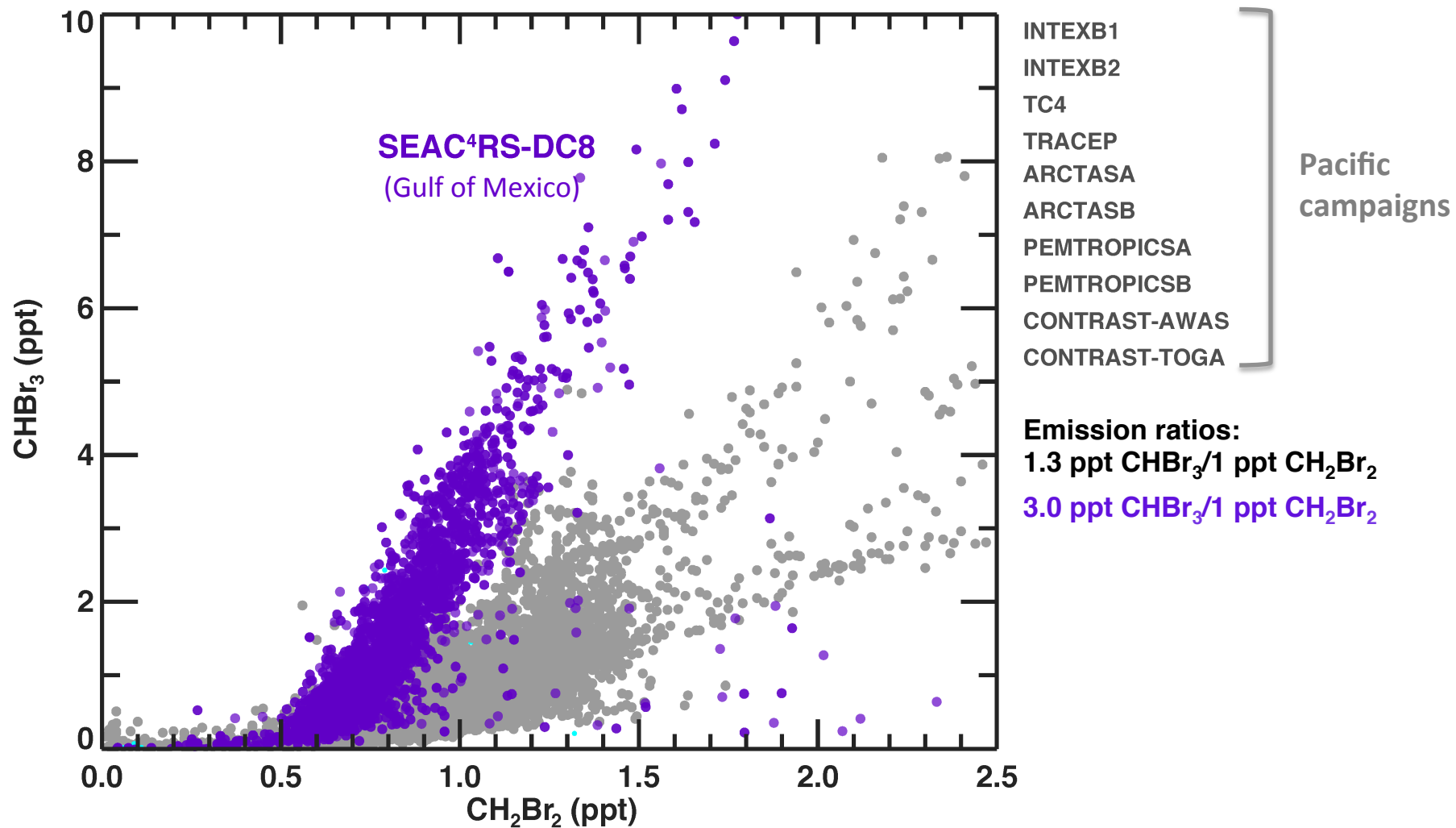
Tropical lower troposphere (Surface - 800 hPa, 30S-30N)



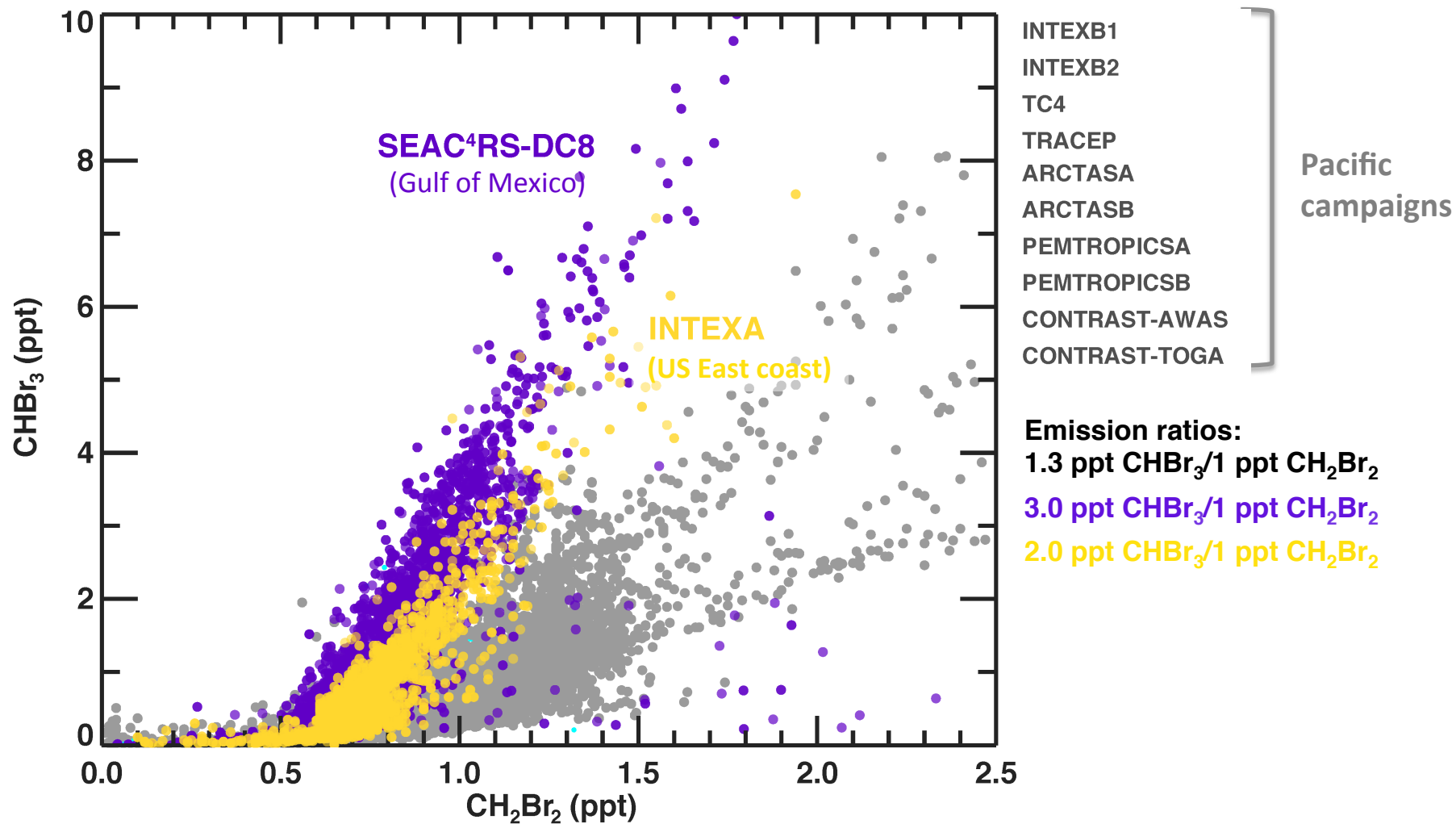
High $\text{CHBr}_3/\text{CH}_2\text{Br}_2$ ratios during SEAC⁴RS



High $\text{CHBr}_3/\text{CH}_2\text{Br}_2$ ratios during SEAC⁴RS

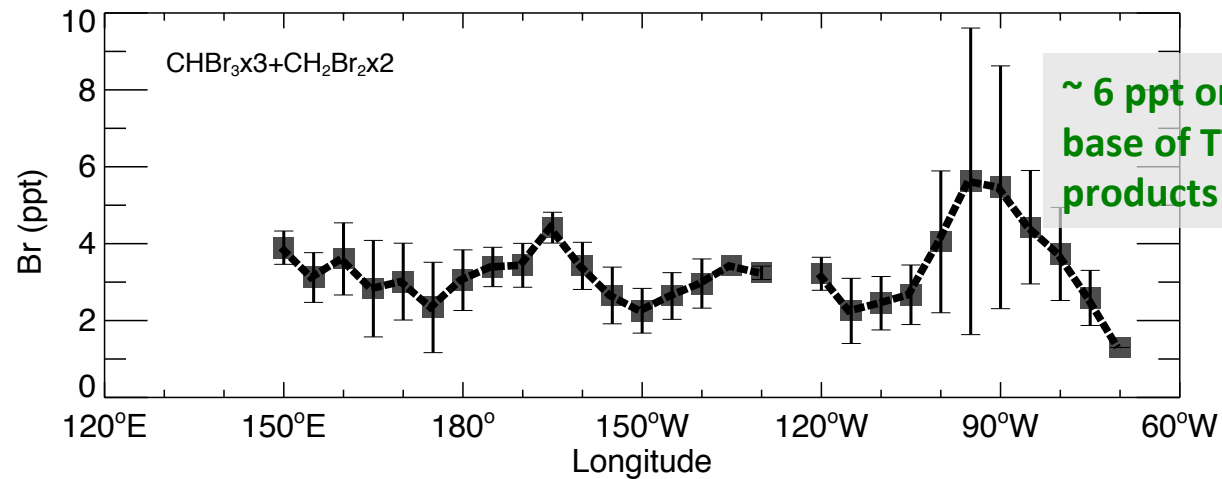
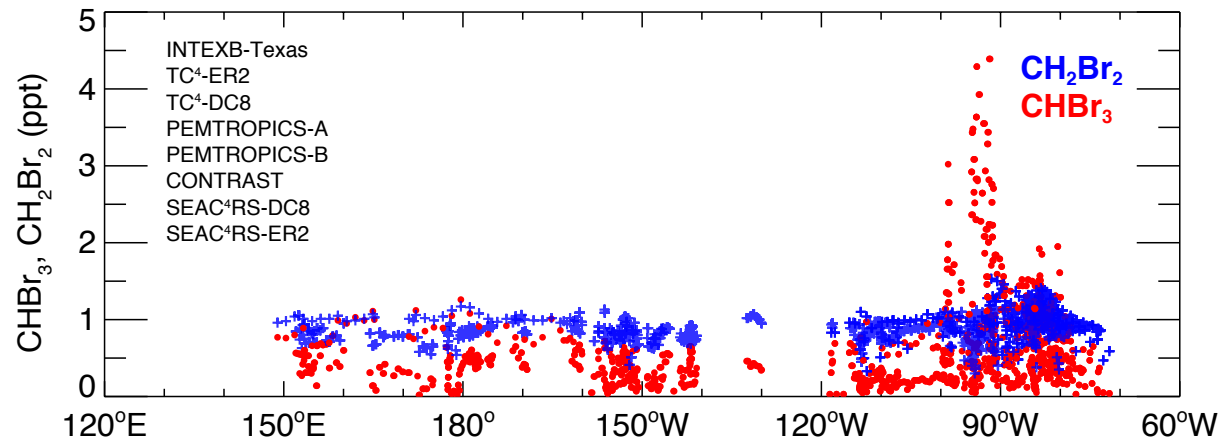


High $\text{CHBr}_3/\text{CH}_2\text{Br}_2$ ratios during SEAC⁴RS



The Atlantic Ocean features much higher biogenic bromocarbon emissions, in particular CHBr_3 , than the Pacific Ocean

Tropical upper troposphere (100-250 hPa, 30S-30N)





Conclusions and future work

- An optimal approach is to use a combination of chemical tracers to examine UT/LS transport
 - CO₂ and Age of Air:**
 - i) Rapid convective ventilation below 355K topped by slow ascent from 355K to 450K, ~5-6 days/K
 - ii) Western Pacific is convectively efficient: 60 days from surface to 365K
 - iii) The North American lower stratosphere is rather aged: 120 days from surface to 365K
 - CO and C₂H₆:**
 - i) Active convective transport from surface to UT/LS in the western Pacific (350-360K)
 - ii) Convective transport in continental N. America is sporadic, but can have injection as high as 400K
- The abundance of very-short-lived bromocarbons
 - About 5 ppt Br from CHBr₃ and CH₂Br₂ in the Western Pacific
 - The Gulf of Mexico features very high CHBr₃ and CH₂Br₂, ~ 12 ppt Br
 - The Gulf of Mexico (and the Atlantic Ocean) features very high CHBr₃/CH₂Br₂ ratios, implying a need to revisit the CHBr₃ and CH₂Br₂ emissions inventory

Future work

- Extend the data analysis to more chemical tracers and additional campaigns (different seasons and different regions)