

# Impact of Overshooting Deep Convection on the Stratospheric Water Vapor: an A-Train Satellite View

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# Objectives

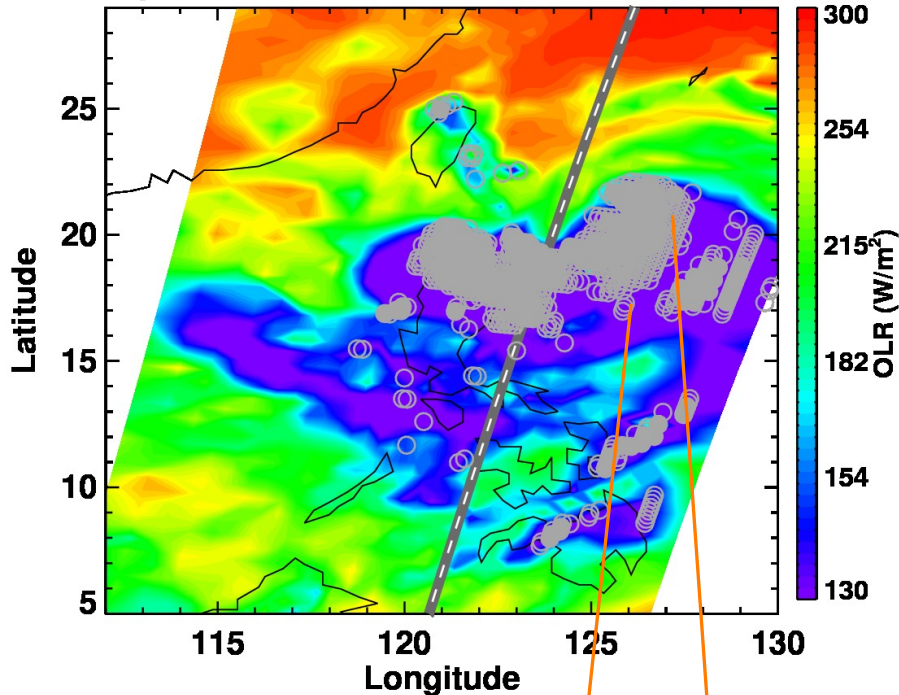
- Does overshooting deep convection (ODC) hydrate or dehydrate the lower stratosphere and by how much?
- How does the convective influence on the stratosphere water vapor (SWV) evolve with climate change ?



- ❑ Combine deep convective clouds (DCC) measurements from CloudSat and Atmospheric Infrared Sounder (AIRS) on Aqua with the UTLS data from the Microwave Limb Sounder (MLS) on Aura.

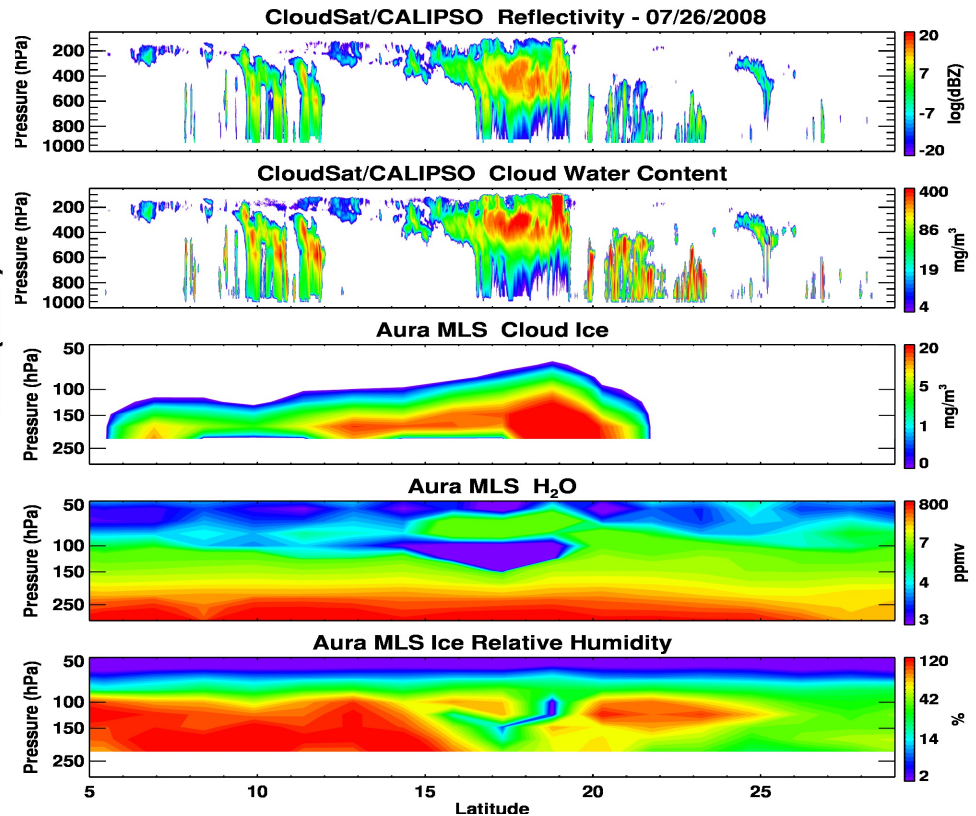
# A-Train Advantage

## Aqua-AIRS OLR and DCC Locations



$$DTW = BT_{1231} - BT_{1419} < 0 \text{ or } < -4$$

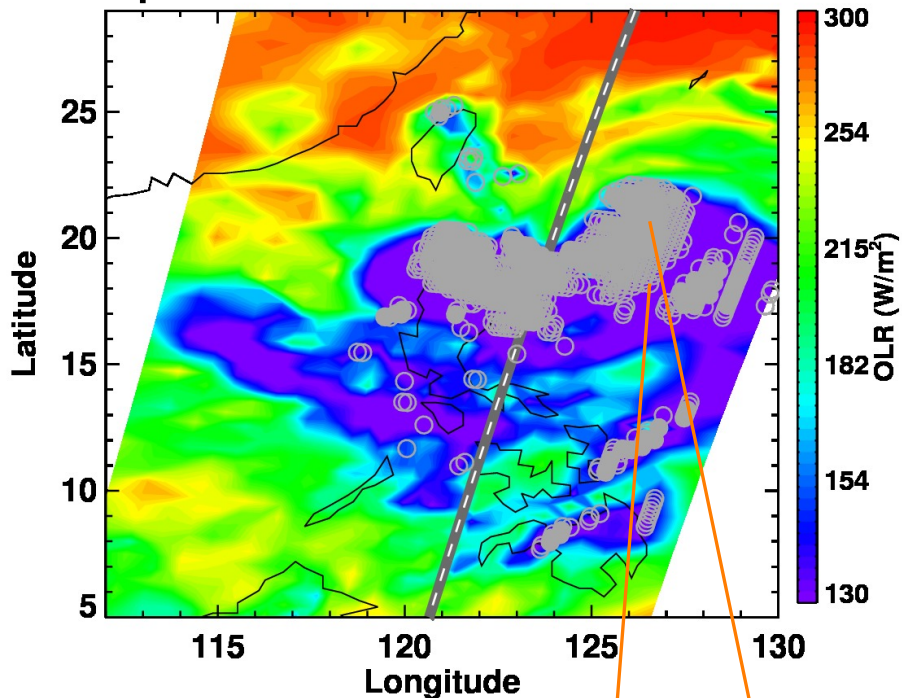
Aumann and Ruzmaikin (2013, ACP)



- AIRS: broad spatial coverage
- CloudSat: vertically resolved cloud structure
- MLS: vertical resolved water vapor and temperature profiles

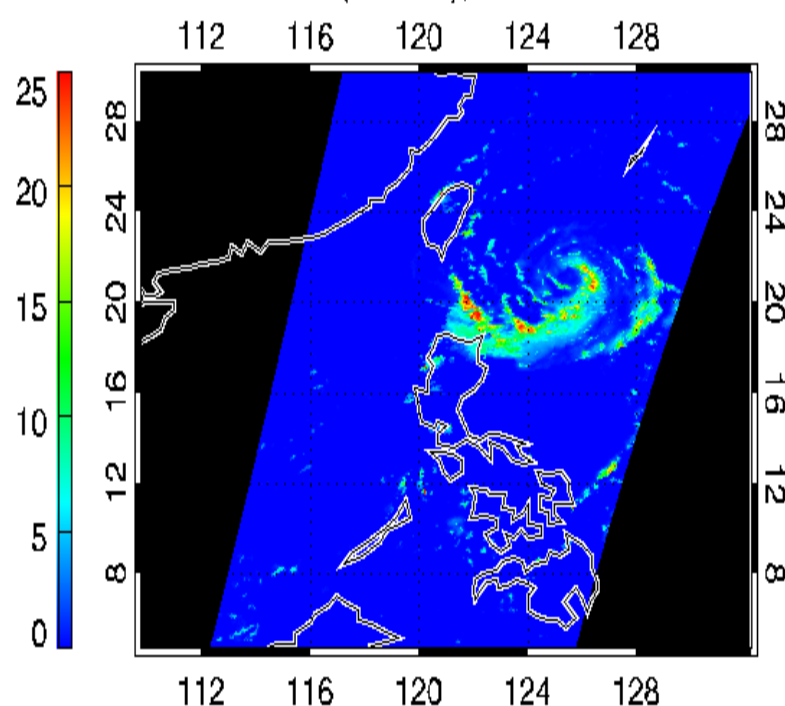
# AIRS DCC

**Aqua-AIRS OLR and DCC Locations**



$$DTW = BT_{1231} - BT_{1419} < 0 \text{ or } < -4$$

**AMSRE rain rate (mm/hr), 2008/07/26 17:18**

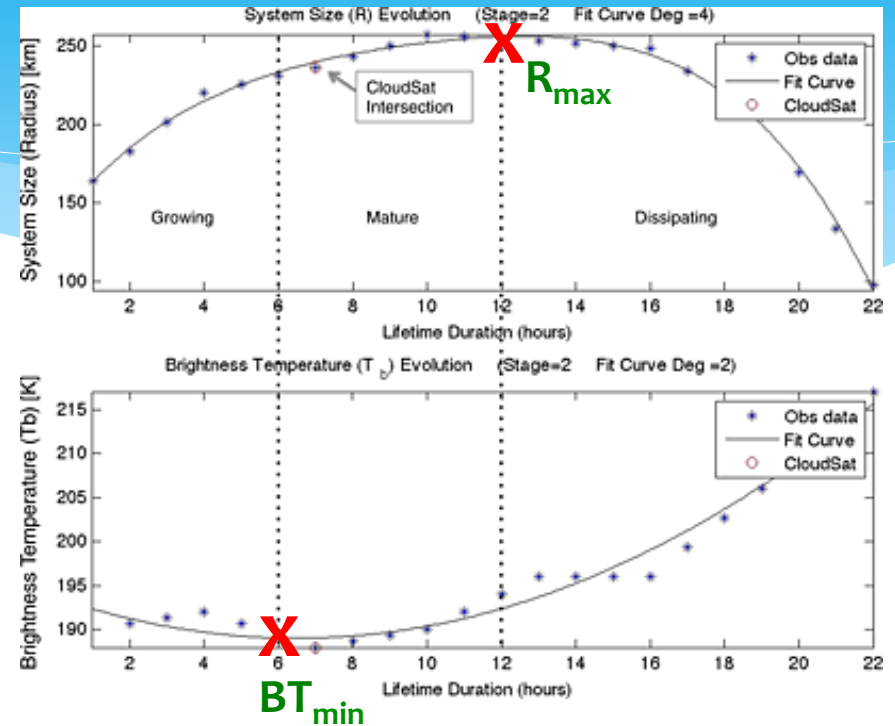
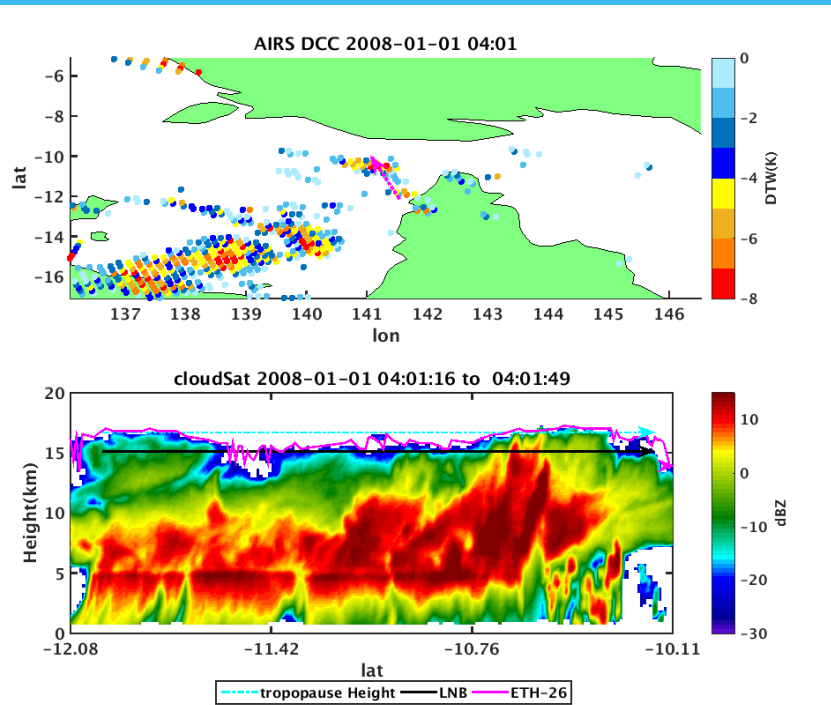


Aumann and Ruzmaikin (2013, ACP)

**Table 1.** Percent of DCCs identified with various IR brightness temperature thresholds and differences, segregated by surface type and time of day, from AIRS data in 2002–2012.

	Threshold	Tropical zone	Ocean day	Ocean night	Land day	Land night
DCC210	bt900 < 210K	0.800	0.557	0.650	0.932	1.09
DCCw0	bt1231 - bt1419 < 0	0.644	0.501	0.629	0.847	1.01
DCCw4	bt1231 - bt1419 < -4	0.158	0.107	0.184	0.235	0.165
DCC200	bt900 < 200K	0.140	0.092	0.161	0.197	0.139
DCCt2	bt900 - T <sub>Trop</sub> < 2	0.061	0.031	0.070	0.072	0.130

# CloudSat ODC



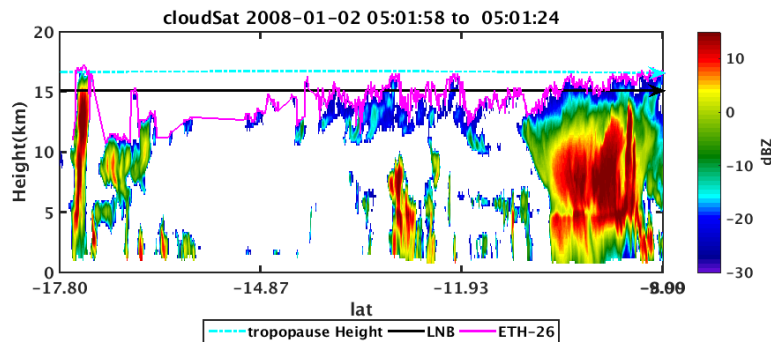
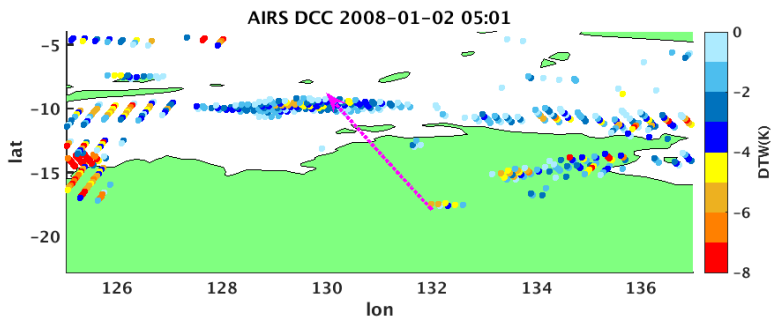
## CloudSat ODC selection criteria:

1. Cloud class: deep convection clouds
2. Cloud Top Height (CTH)  $\geq 10$  km
3. Cloud Base Height (CBH)  $\leq 2$  km
4. Continuous radar echo
5.  $CTH = LNB + \delta$  ( $\delta > 0$ )

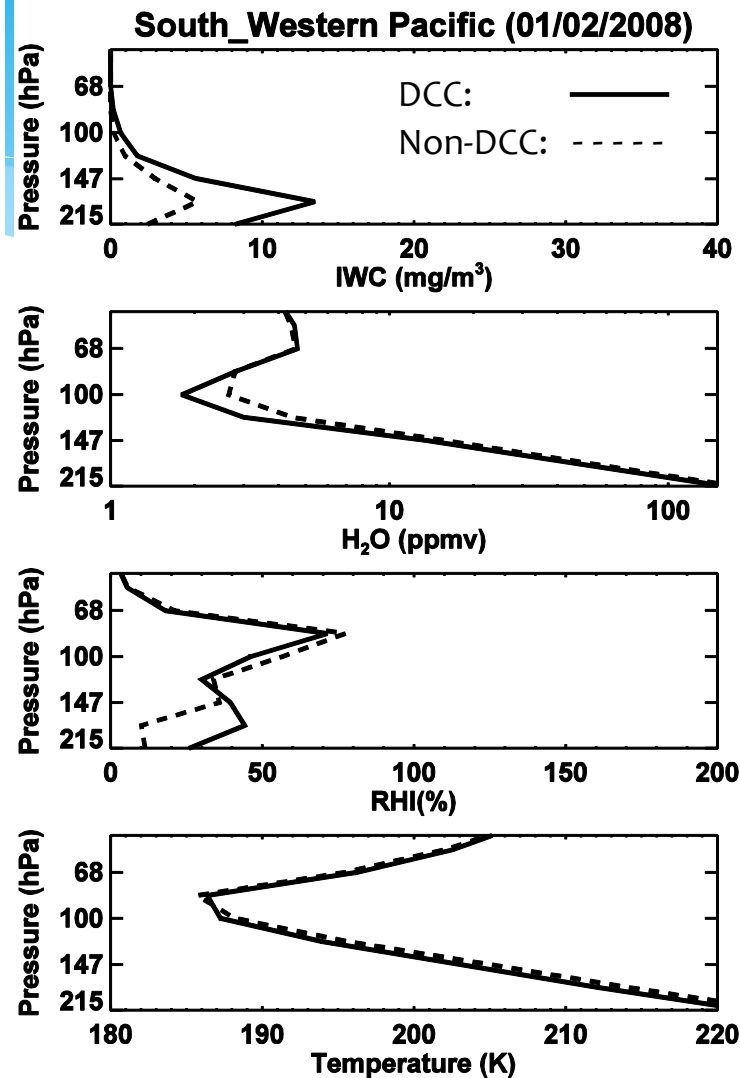
## ISCCP-CT:

- **Growing stage:** before reaching  $BT_{min}$
- **Dissipating Stage:** after reaching  $R_{max}$
- **Mature Stage:** in-between growing and dissipating stages

# A DCC in Western Pacific (winter)



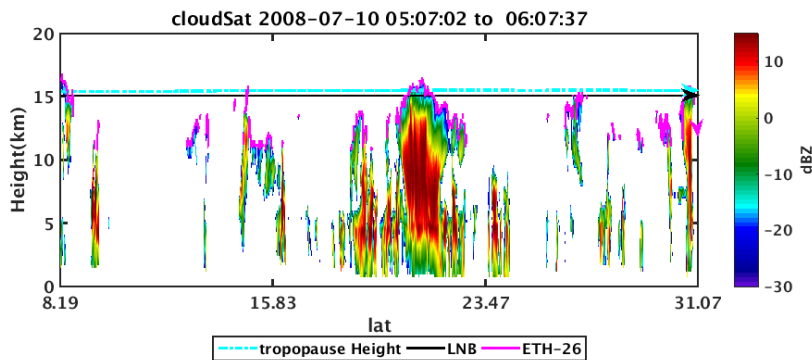
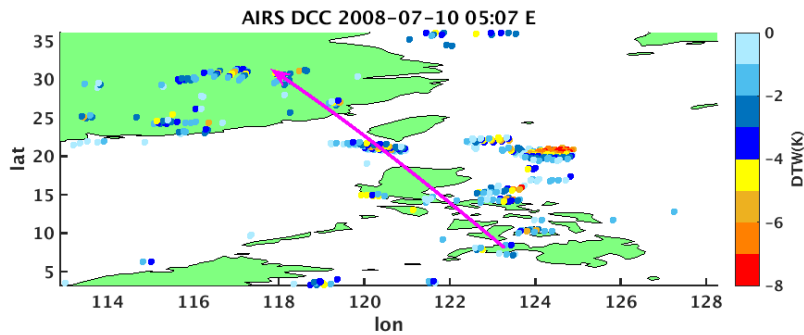
DTW =  $-5.2$  K  
 CTH =  $17.3$  km  
 TH =  $16.7$  km  
 TH: tropopause height



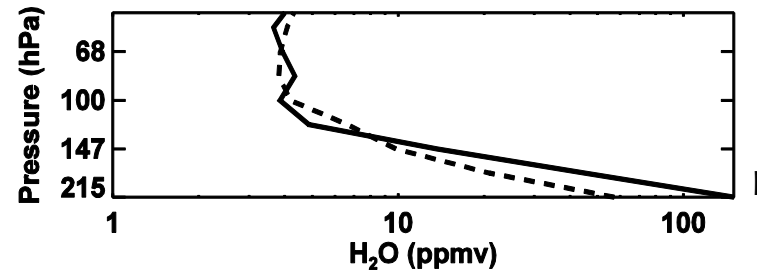
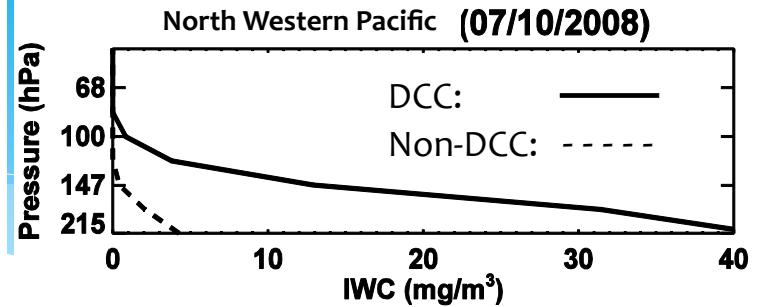
83 hPa  $\text{H}_2\text{O}$   
 DCC: **2.8**  
 Non-DCC: **2.8**

- Dehydration in the TTL, no difference in the LS

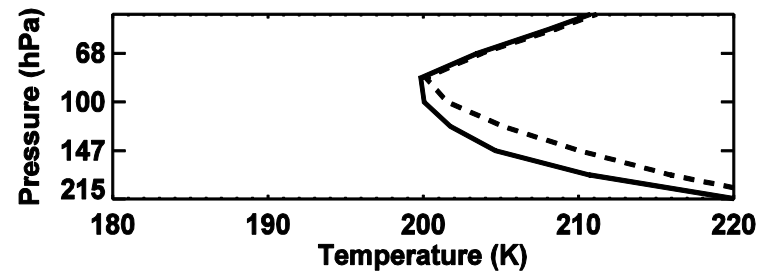
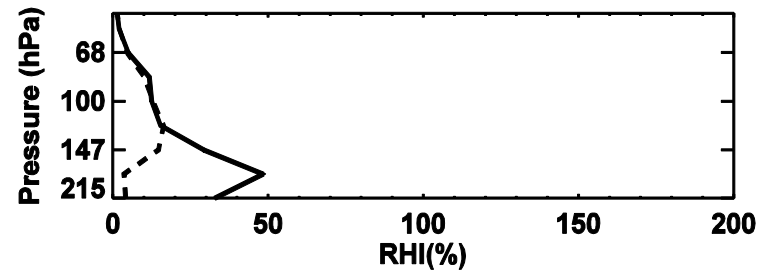
# A DCC in Western Pacific (summer)



DTW = -5.2 K  
 CTH = 16.5 km  
 TH = 15.5 km

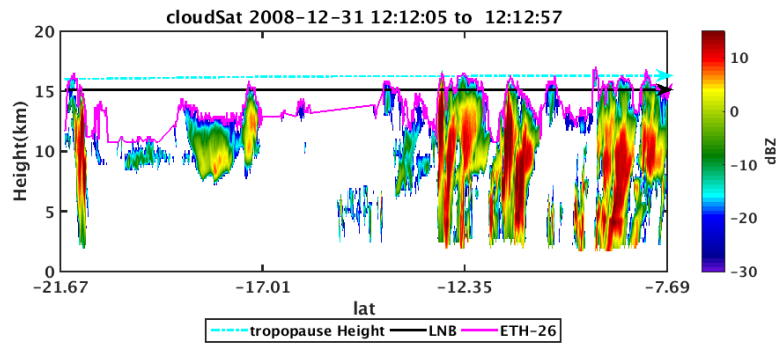
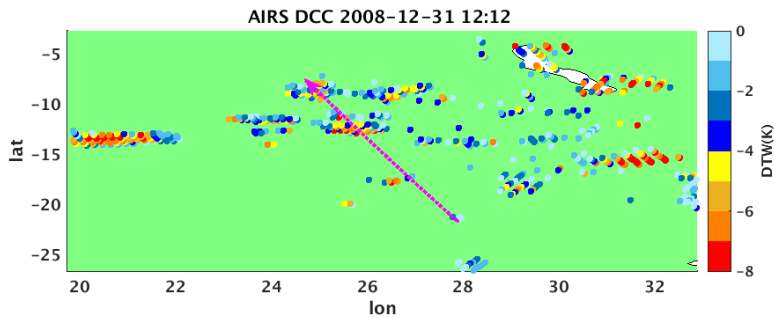


83 hPa H<sub>2</sub>O  
 DCC: 4.3  
 Non-DCC: 3.8

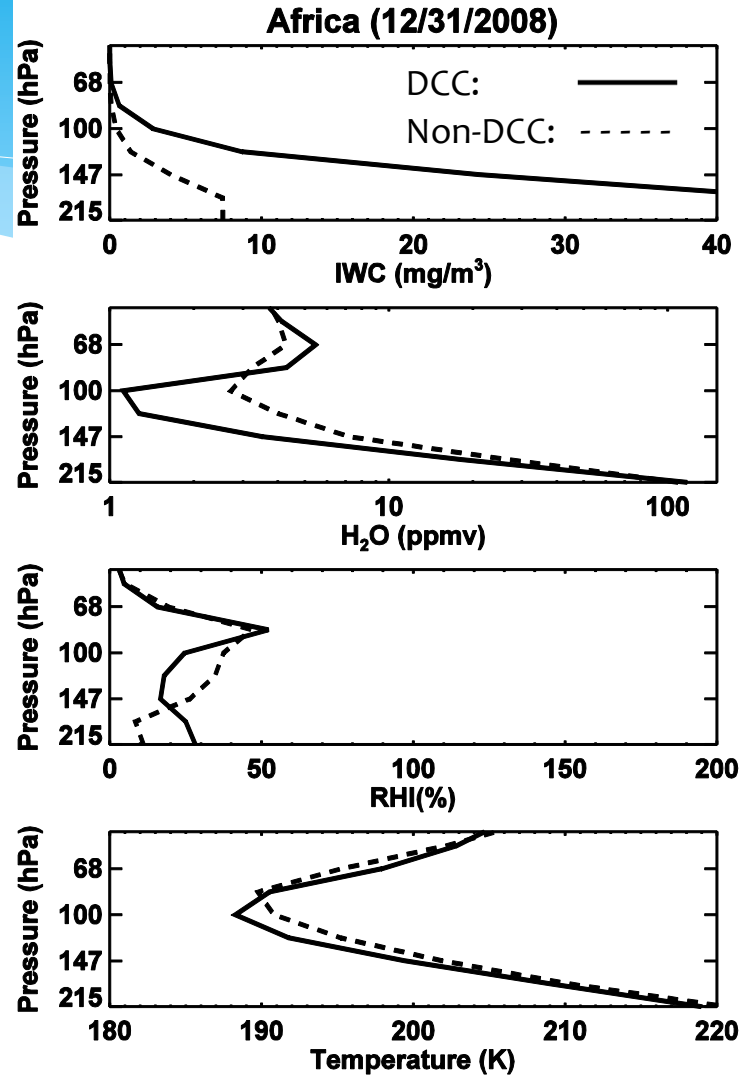


- Dehydration in the TTL, hydration in the LS by 0.5 ppmv

# A DCC in Central Africa (austral summer)



DTW = -6.0 K  
 CTH = 16.5 km  
 TH = 16.3 km

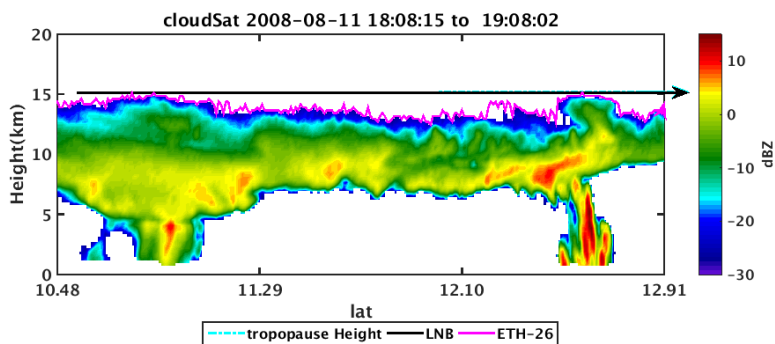
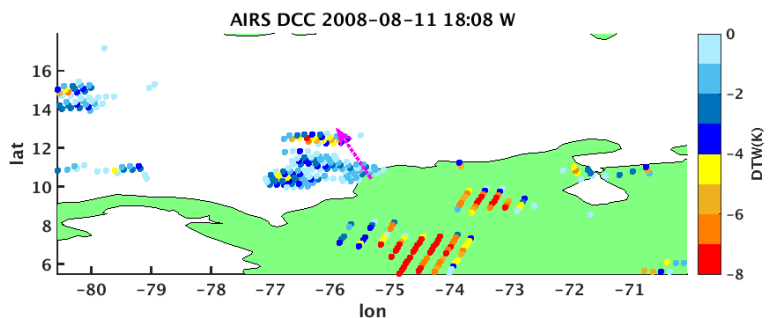


83 hPa H<sub>2</sub>O  
 DCC: 4.3  
 Non-DCC: 3.2

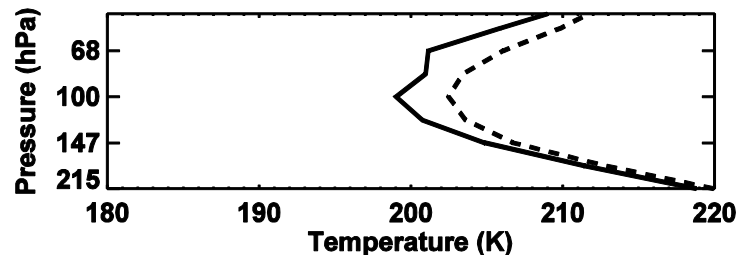
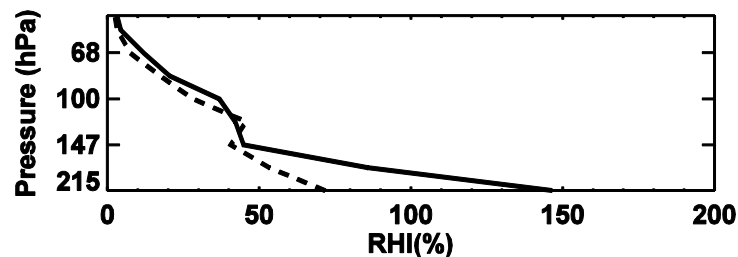
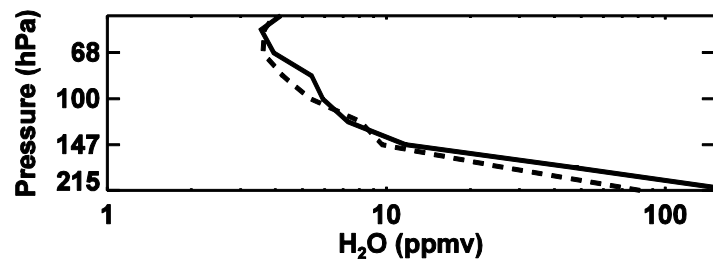
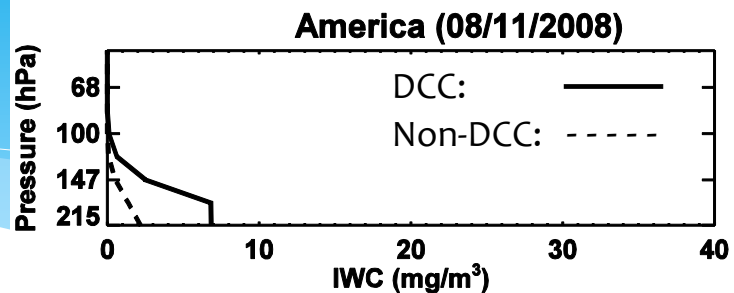
- Dehydration in the TTL, hydration in the LS by 1.1 ppmv



# A DCC in Central America (summer)



DTW = -2.0 K  
 CTH = 15.1 km  
 TH = 15.2 km

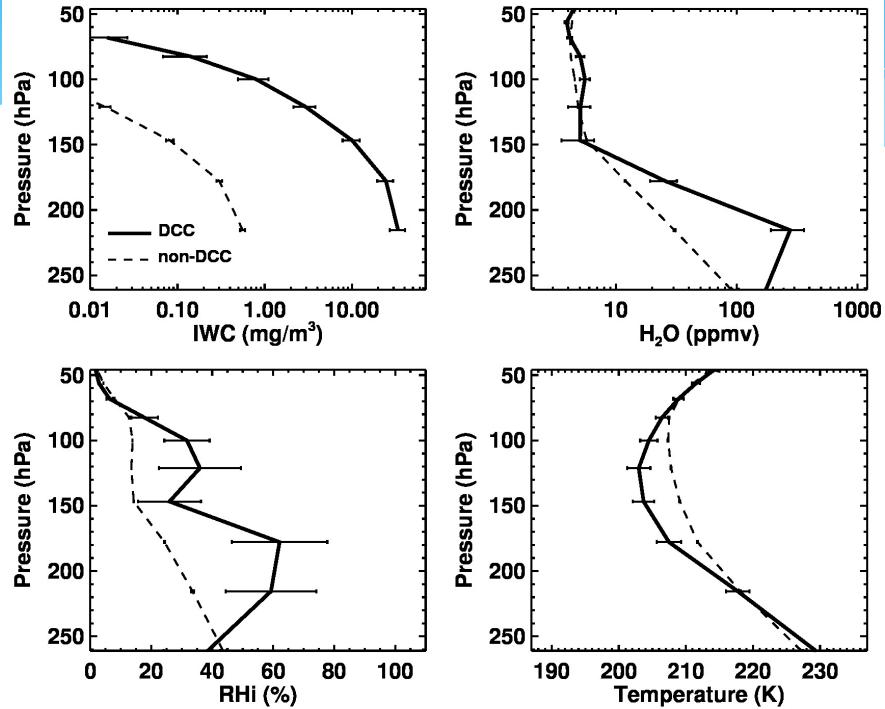


83 hPa H<sub>2</sub>O  
 DCC: 5.4  
 Non-DCC: 4.3

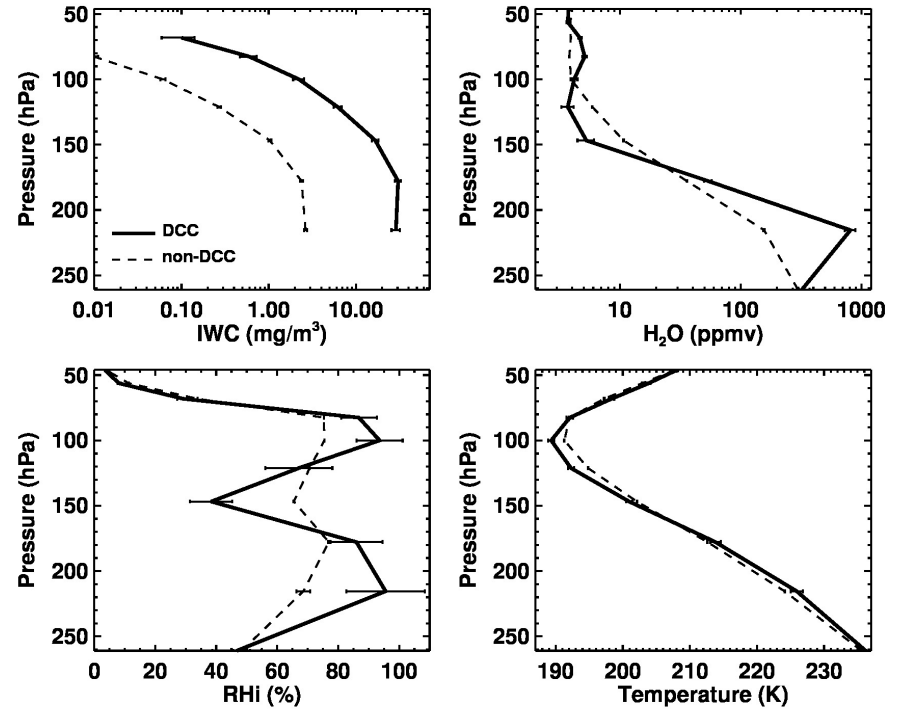
- No dehydration in the TTL, hydration in the LS by 1.1 ppmv

# Two Summer Monsoon Regions

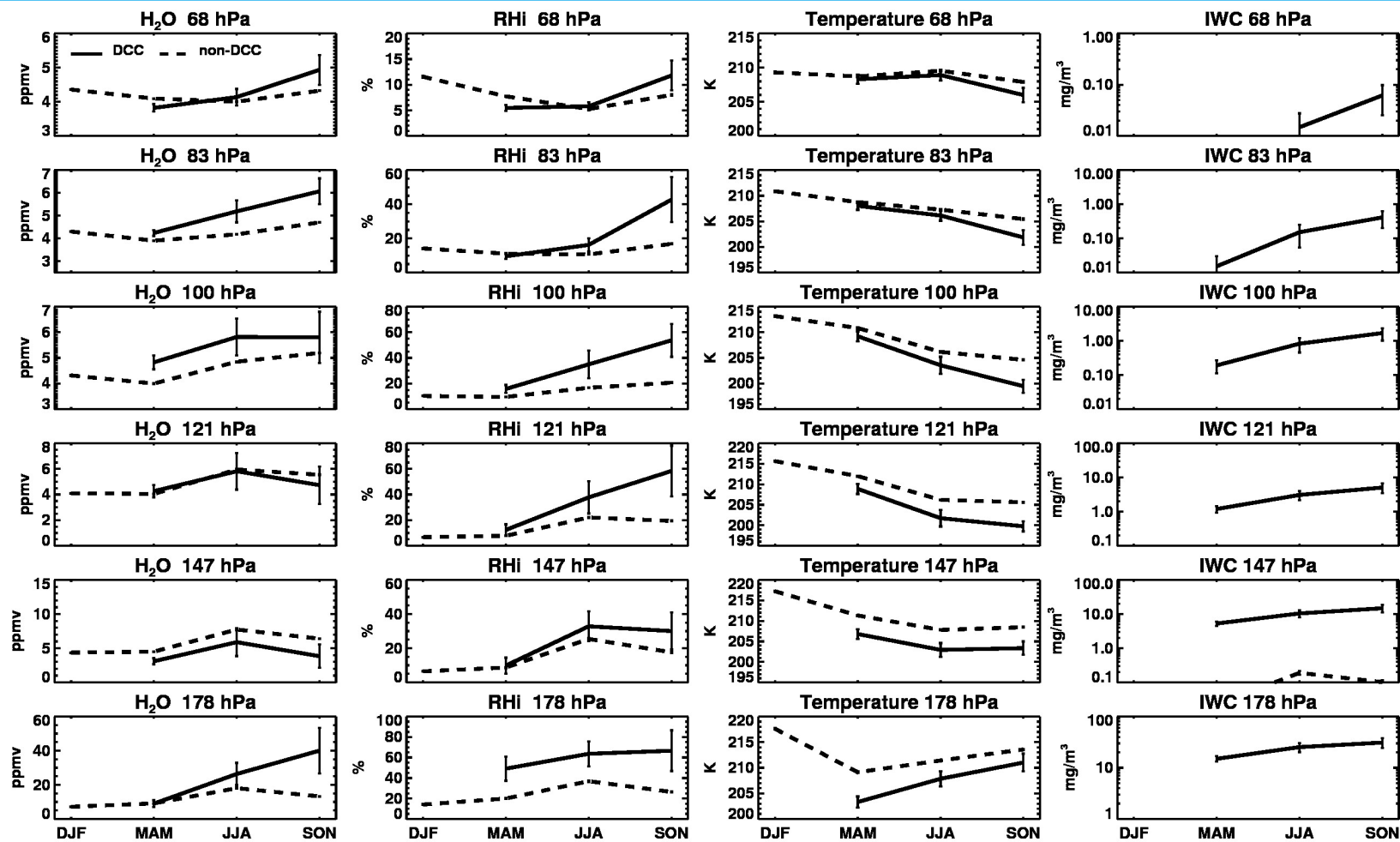
North America (Lon:250-290°; Lat:20°N-50°N)



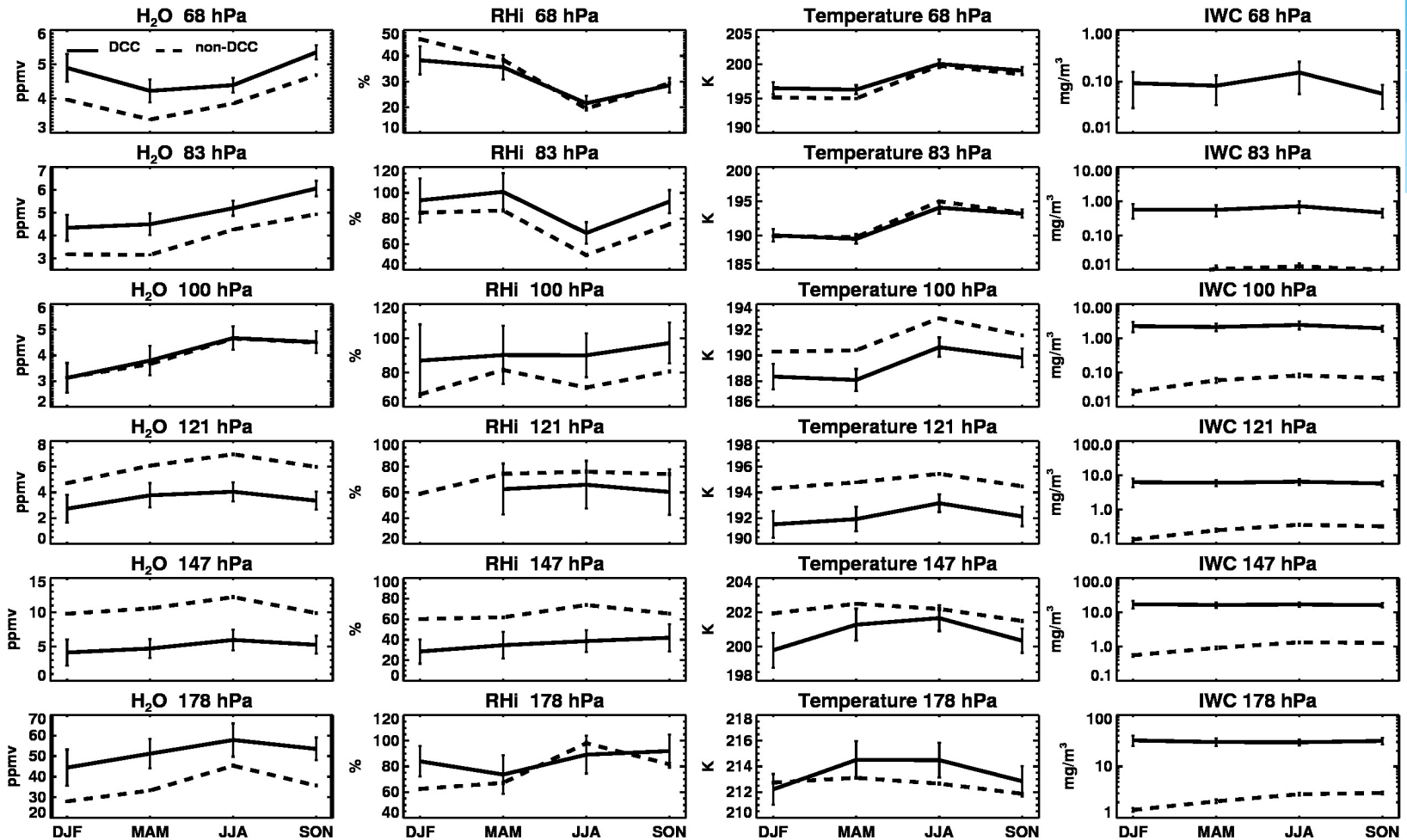
South Asia (Lon:70-110°; Lat:0°-20°N)



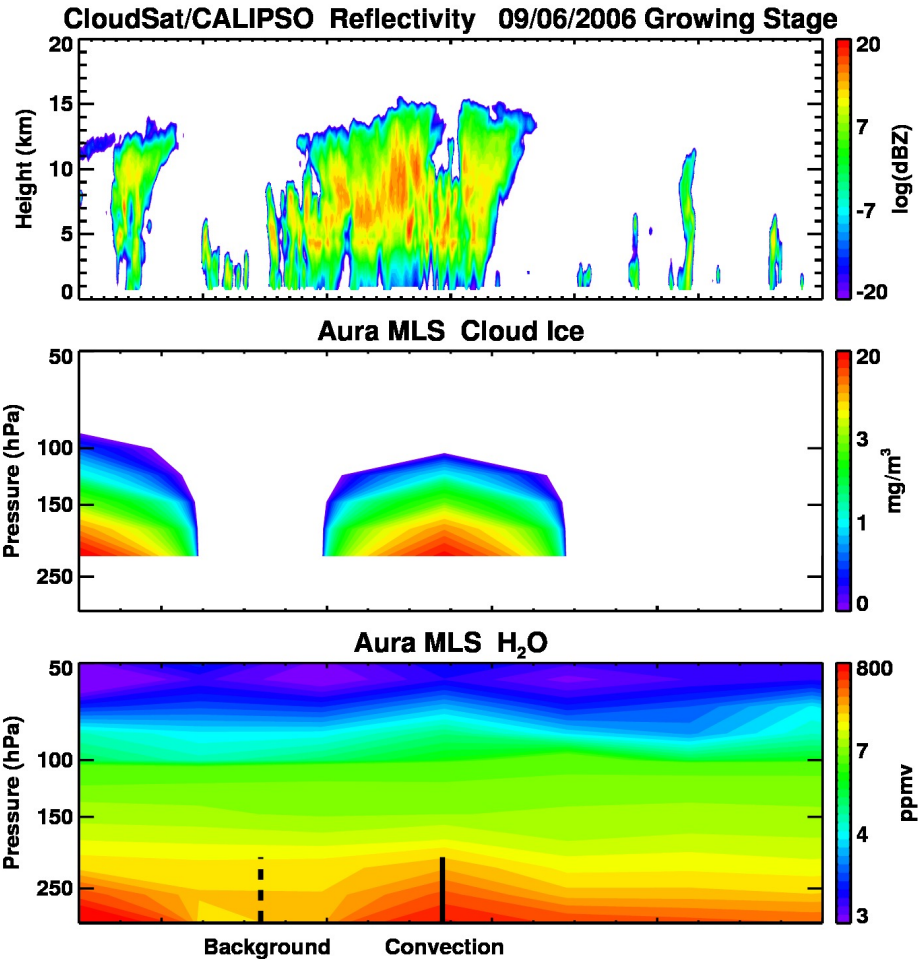
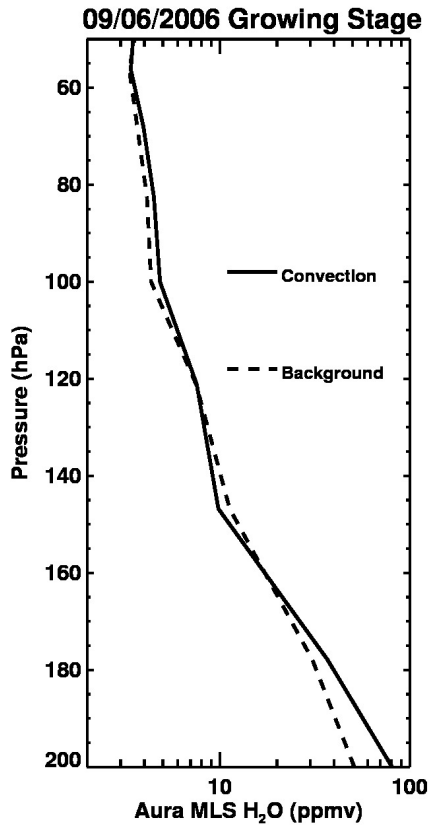
# Seasonal Variation (North America)



# Seasonal Variation (South Asia)

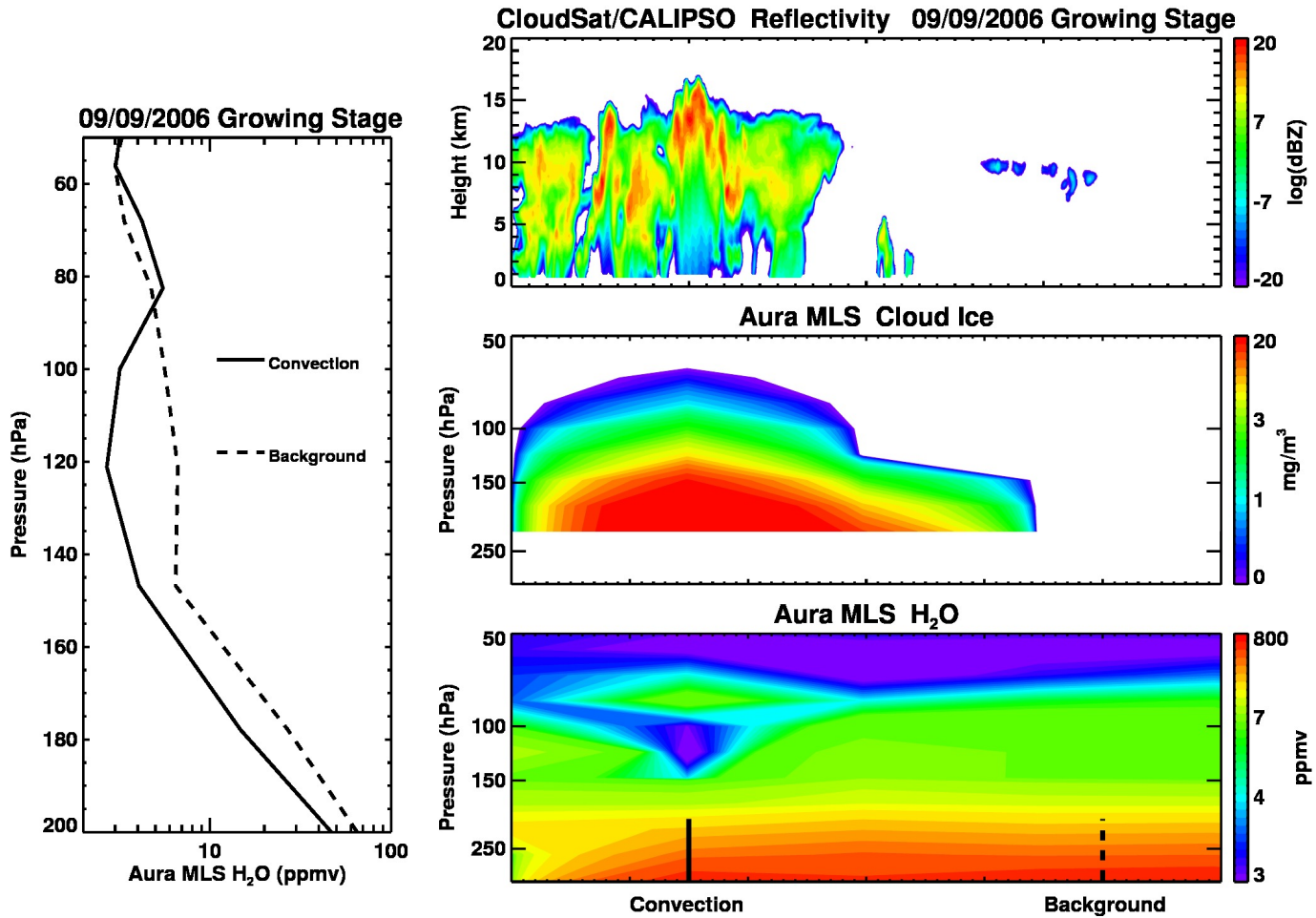


# A Growing Stage ODC



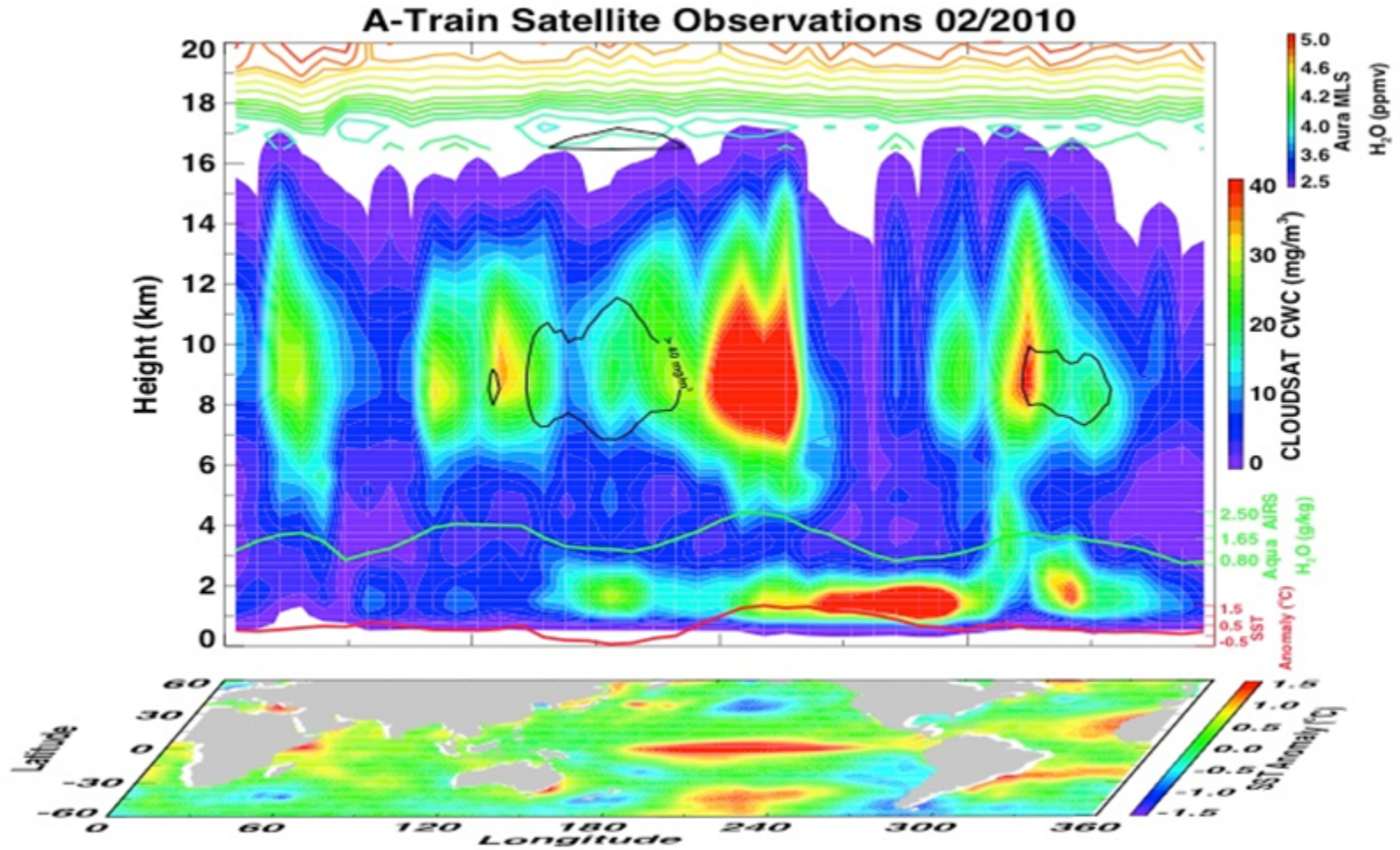
Overshooting Deep Convection at (4.33°N, 156.06°E)

# A Mature Stage DCC

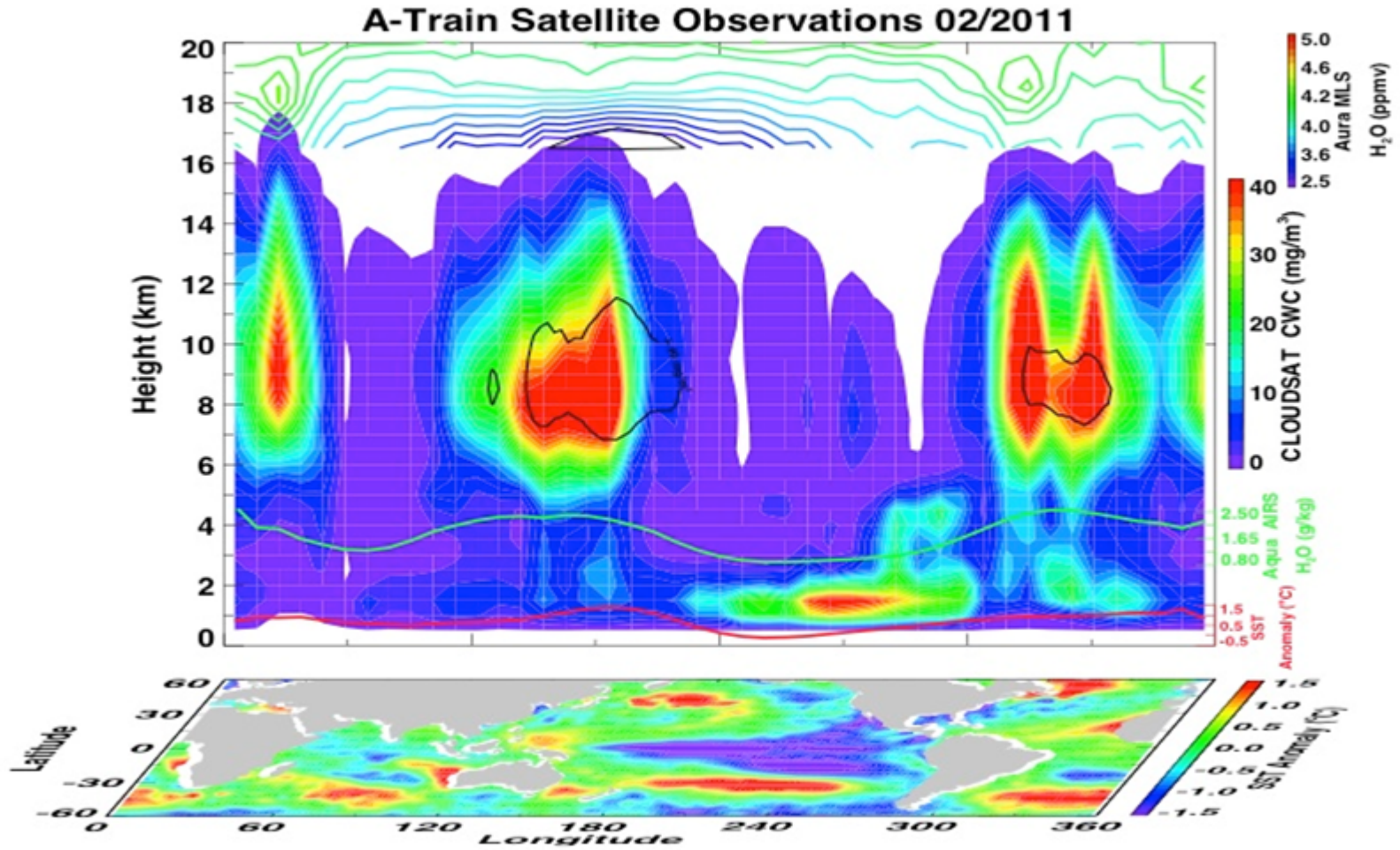


Overshooting Deep Convection at (5.0°N, 173.20°E)

# Interannual Variations

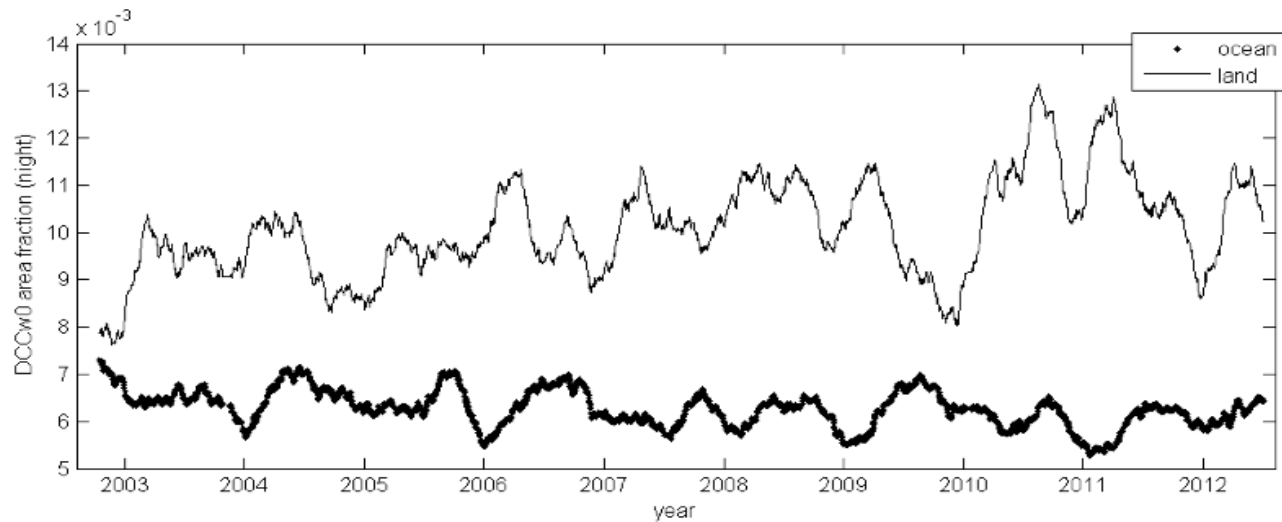
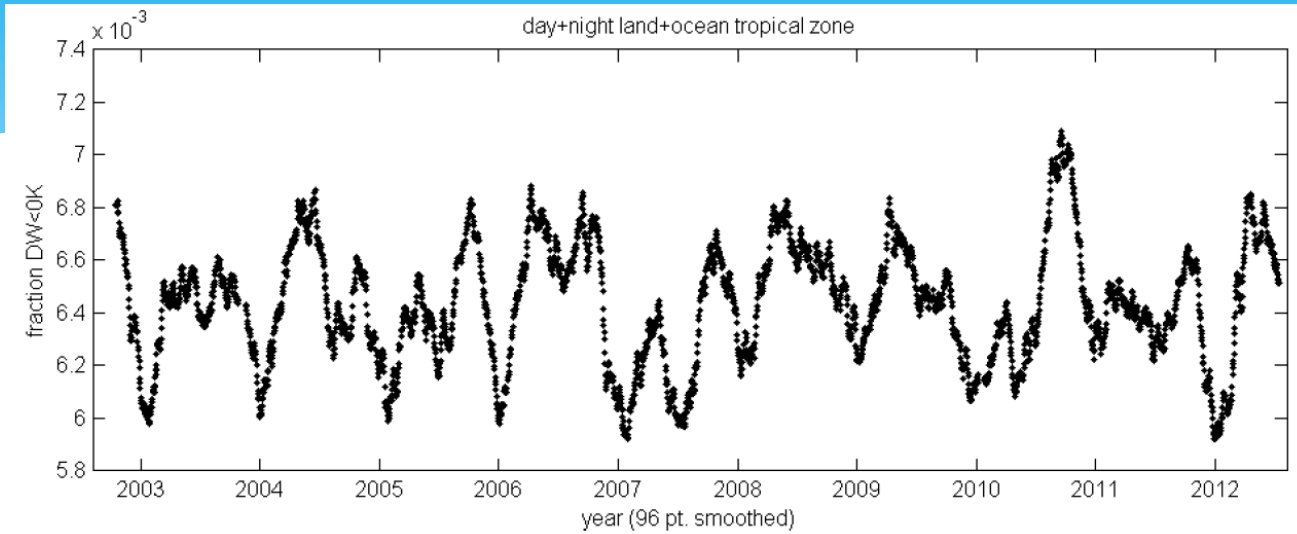


# Interannual Variations





# Ten-Year Trends



# Summary

- We find that ODC tends to cool and dehydrate the TTL, but can moisten the lower stratosphere up to 30%. However, large regional and seasonal differences exist for the deep convective impact on the SWV.
- The convective influence on the SWV also depends on the life cycles of ODC.
- Interannual variations of the SWV are influenced by the coupling between deep convection and tropopause temperature.
- Although ODC could moisten the lower stratosphere, its rarity makes the overall convective moistening effect quite small on the annual mean water vapor budget ( $< 0.2\%$ ), but it may play a significant role in the SWV variabilities.
- Whether the ODC influence on the SWV would strengthen under global warming is still an open question.