

Trajectory and Microphysical Modeling of H₂O and Clouds in the Tropical Tropopause Layer

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Introduction

Stratospheric humidity is mainly controlled by the freeze-drying of tropospheric air as it ascends across the tropical cold-point tropopause (Fig. 1a). The details of the Tropical Tropopause Layer (TTL) dehydration mechanism, including the roles of deep convection, waves and cloud microphysical processes (Fig. 1b), are not well understood. The goal of this research is to better understand the processes that control H₂O and cirrus clouds in the TTL.

- ★ How well do our trajectory and cloud microphysical models simulate the observed H₂O and cloud fields in the TTL?
- ★ What are the impacts of convection, subgrid-scale waves, and cloud microphysics on TTL humidity and cirrus cloud abundance?
- ★ How sensitive are the simulated results to the radiative heating rates (which determine the vertical motions of parcels along their diabatic trajectories)?

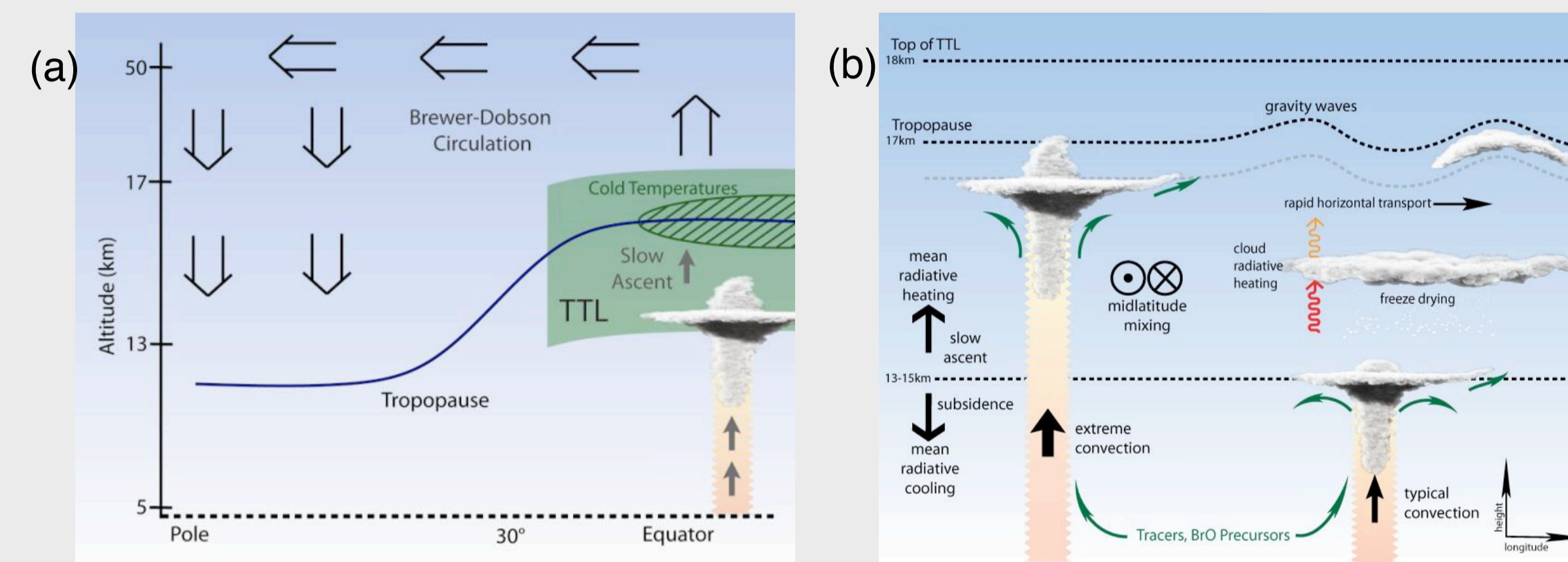


Fig. 1: (a) Schematic of the zonal mean cross-section of the TTL where air is dehydrated ("freeze-dried") as it ascends through the cold-point tropopause. (b) Longitude-height cross-section of the TTL and the many processes that influence TTL humidity and clouds.

Conclusions

• Waves dehydrate (by 0.5 ppmv), whereas convection and microphysics moisten (by 0.6 and 0.7 ppmv, respectively) the 100 hPa level (Fig. 10, Table 1).

• Waves and convection both increase the tropical mean cloud occurrence by ~25% near the 100 hPa (~16.5 km) level (Fig. 9). More than half of the cloud occurrence at and above the cold point is wave driven, while in situ formation of clouds downstream of convection dominates cloud occurrence below ~100 hPa.

• TTL humidity and clouds are sensitive to variations in the radiative heating rates. Temporal variability of the heating rates dehydrate the 100 hPa level and increase cloud occurrence frequencies throughout the TTL (not shown).

• Heterogeneous nucleation and convective ice injection have relatively minor impacts on TTL humidity and cloud frequency (not shown).

Table 1: Effects of waves, convection, and microphysics on the tropical mean 100 hPa H₂O mixing ratio and cloud occurrence frequency in the mid to upper TTL (16 - 18 km). Percent change relative to the base simulation are shown in parenthesis.

	H ₂ O, ppmv	cloud occurrence, %
waves (sub-grid scale)	-0.46 (-19%)	+3.6 (+29%)
convection	+0.56 (+23%)	+3.2 (+26%)
microphysics	+0.71 (+29%)	N/A

Ueyama et al. (2015, JGR, in review)

Methodology

Trajectory Model

- Calculate 60-day backward (diabatic) trajectories on 1 Feb 2007 every 2° latitude x 2° longitude in the tropics (20°S - 20°N) from the 372 K potential temperature (~100 hPa) level.
 - Heating rates (vertical motions) from offline radiative transfer calculations (Yang et al., 2010) and ERA-interim wind and temperature fields
 - ERA-Interim temperatures and winds with enhanced wave-driven variability [Kim and Alexander, 2014] (Fig. 4)

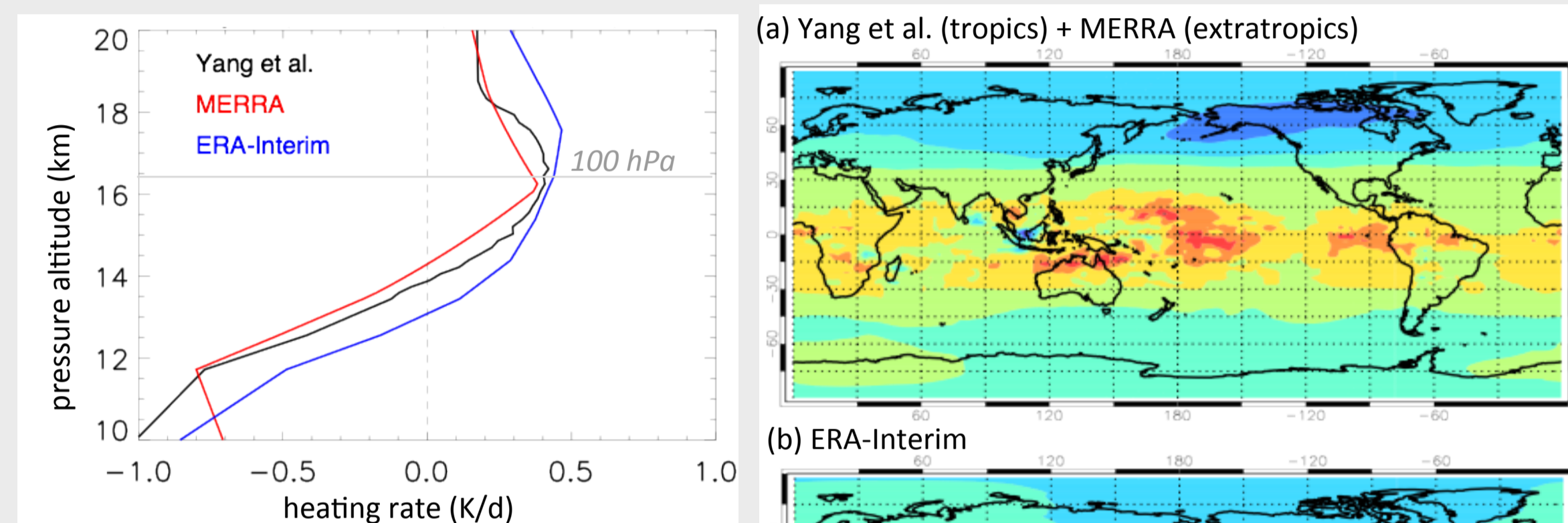
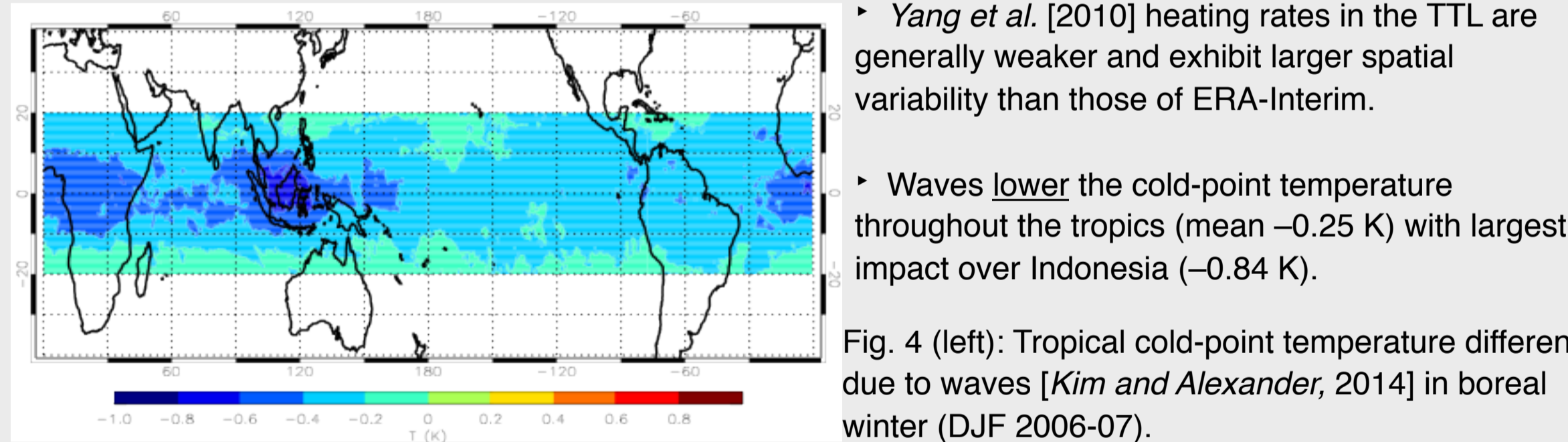


Fig. 2 (top): Vertical profiles of tropical (20°S-20°N) mean radiative heating rate in boreal winter (DJF 2006-07) based on (black) Yang et al. [2010], (red) MERRA and (blue) ERA-Interim reanalysis data.

Fig. 3 (right): Boreal winter (DJF 2006-07) radiative heating rates at ~100 hPa level: (a) Yang et al. + MERRA, (b) ERA-Interim.



- Yang et al. [2010] heating rates in the TTL are generally weaker and exhibit larger spatial variability than those of ERA-Interim.
- Waves lower the cold-point temperature throughout the tropics (mean -0.25 K) with largest impact over Indonesia (-0.84 K).

Cloud Microphysical Model

◆ Time-dependent, one-dimensional (vertical) model that tracks the growth, sedimentation, sublimation of individual ice crystals

- Simulate ice clouds and H₂O along parcel trajectories
- Initialize the model with MLS H₂O profile at Day -60

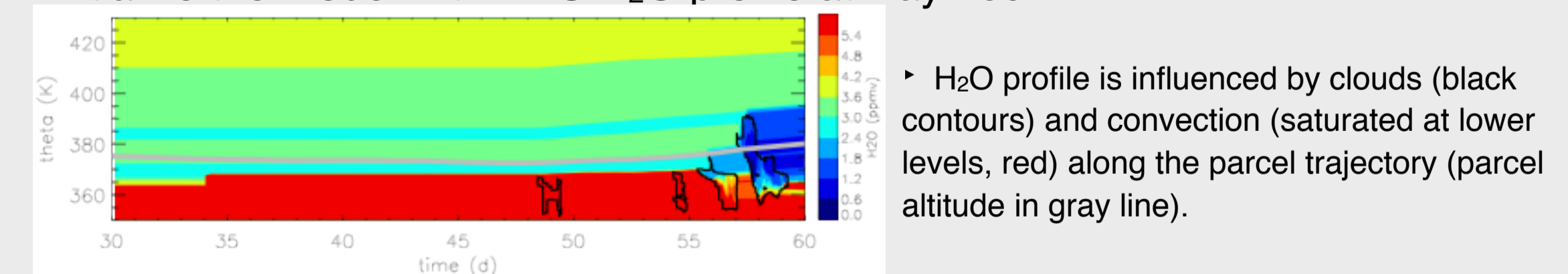


Fig. 5: A sample time-height "curtain" of H₂O mixing ratio of a given parcel trajectory.

◆ Convection scheme

Trace trajectories through geostationary satellite convective cloud-top height fields, and saturate the column up to the cloud-top altitude

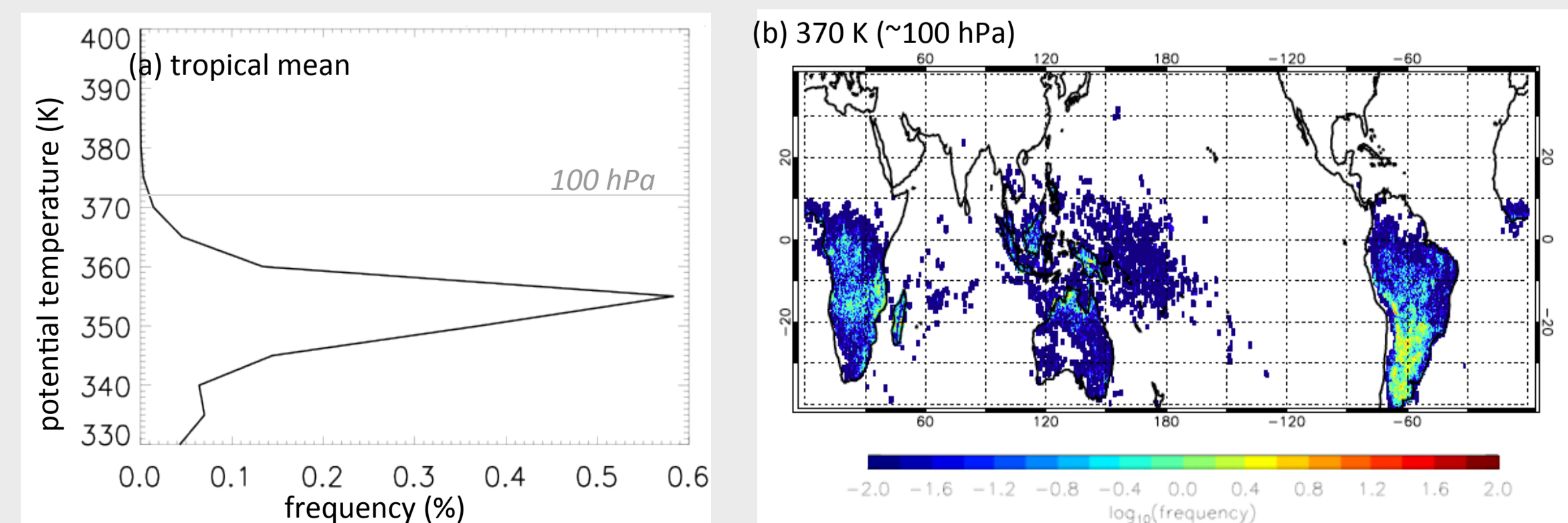


Fig. 6: Convective cloud-top height distribution in boreal winter (DJF 2006-07): (a) tropical (20°S-20°N) mean profile, (b) spatial distribution of convective clouds with tops reaching the 370 K level.

Results

Model vs. Observations

- The simulated 100 hPa H₂O field agrees reasonably well with MLS observations ($r = 0.54$, RMSE = 0.5 ppmv), but is ~20% too dry.
- Spatial distribution of cloud occurrence frequencies in the mid to upper TTL is well correlated with that of CALIPSO

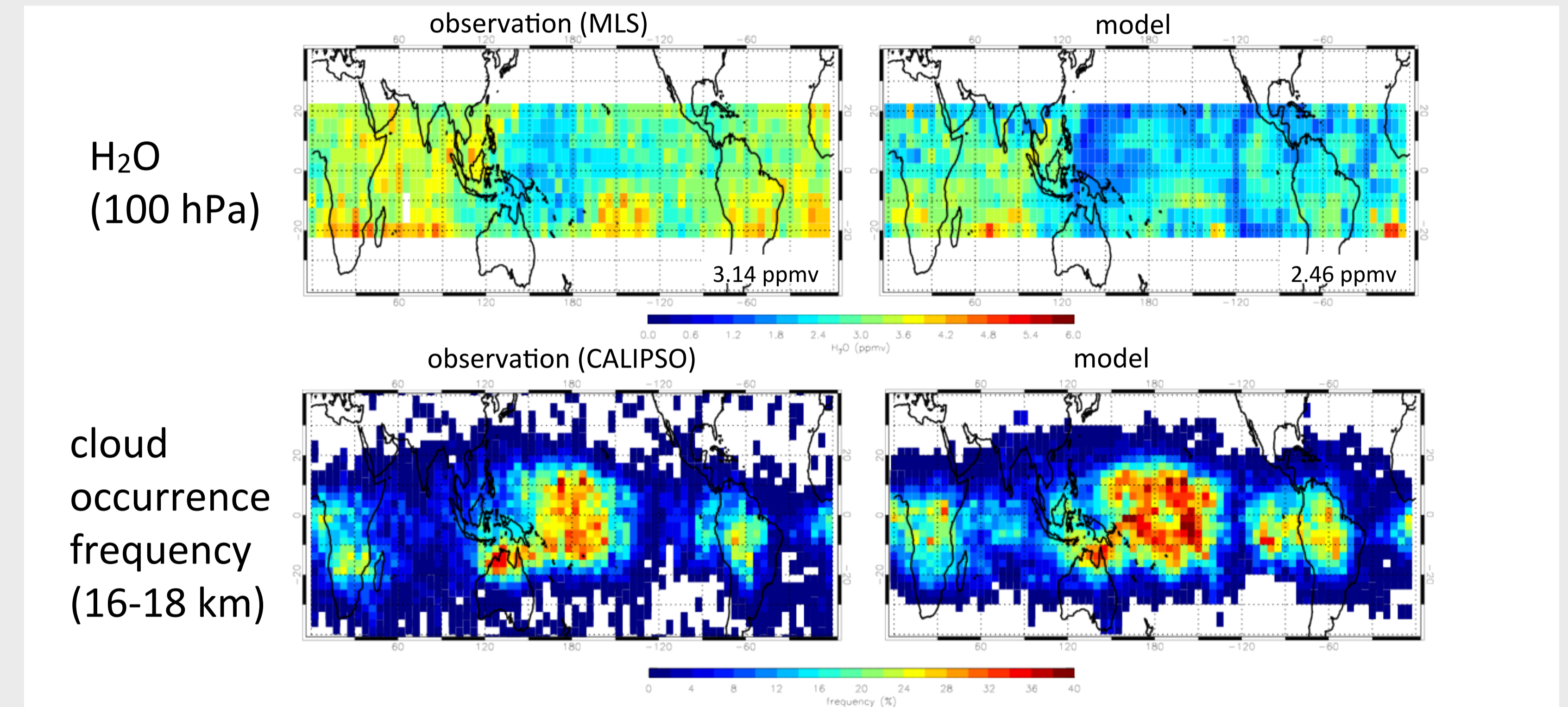


Fig. 7: (top) The 100 hPa H₂O mixing ratios on 1 Feb 2007 of the centered 7-day mean MLS observation and model. (bottom) Cloud occurrence frequencies in the mid to upper TTL (16-18 km) during Dec 2006 - Jan 2007 based on CALIPSO observation and model data.

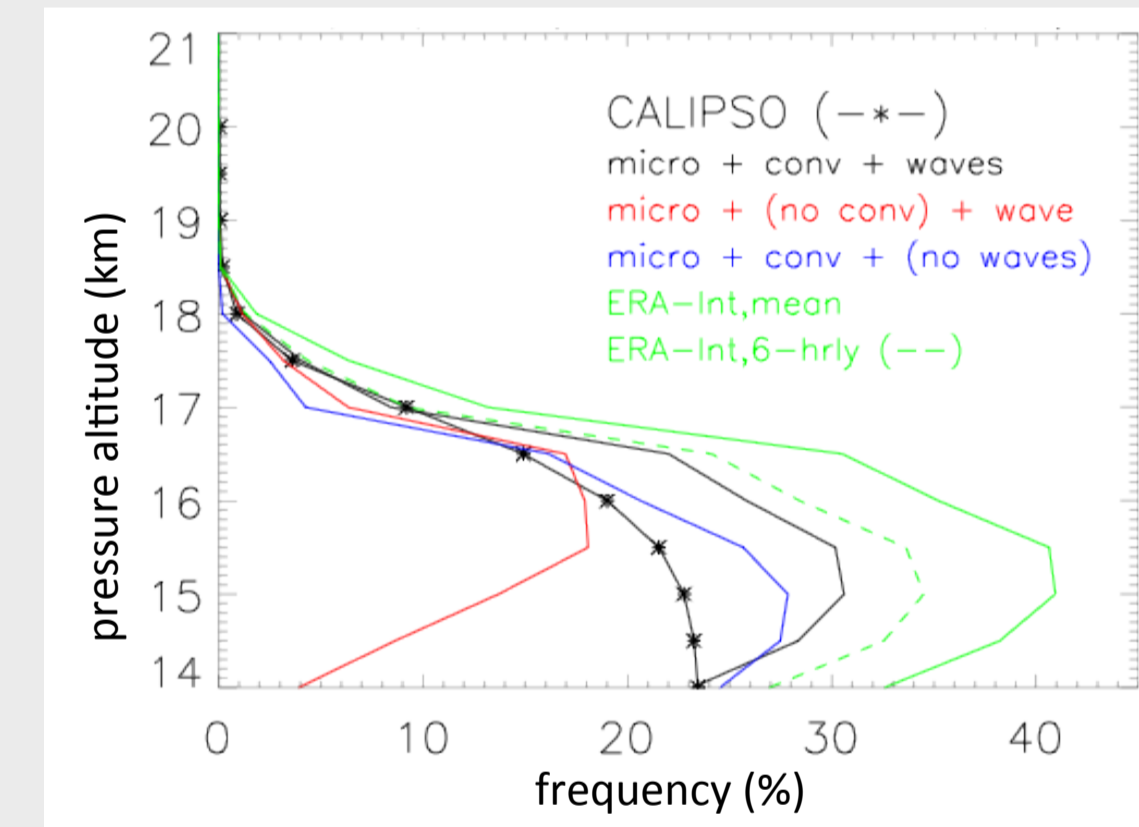


Fig. 8: Vertical profiles of tropical (20°S-20°N) mean cloud occurrence frequency in boreal winter (DJF 2006-07) based on CALIOP and model.

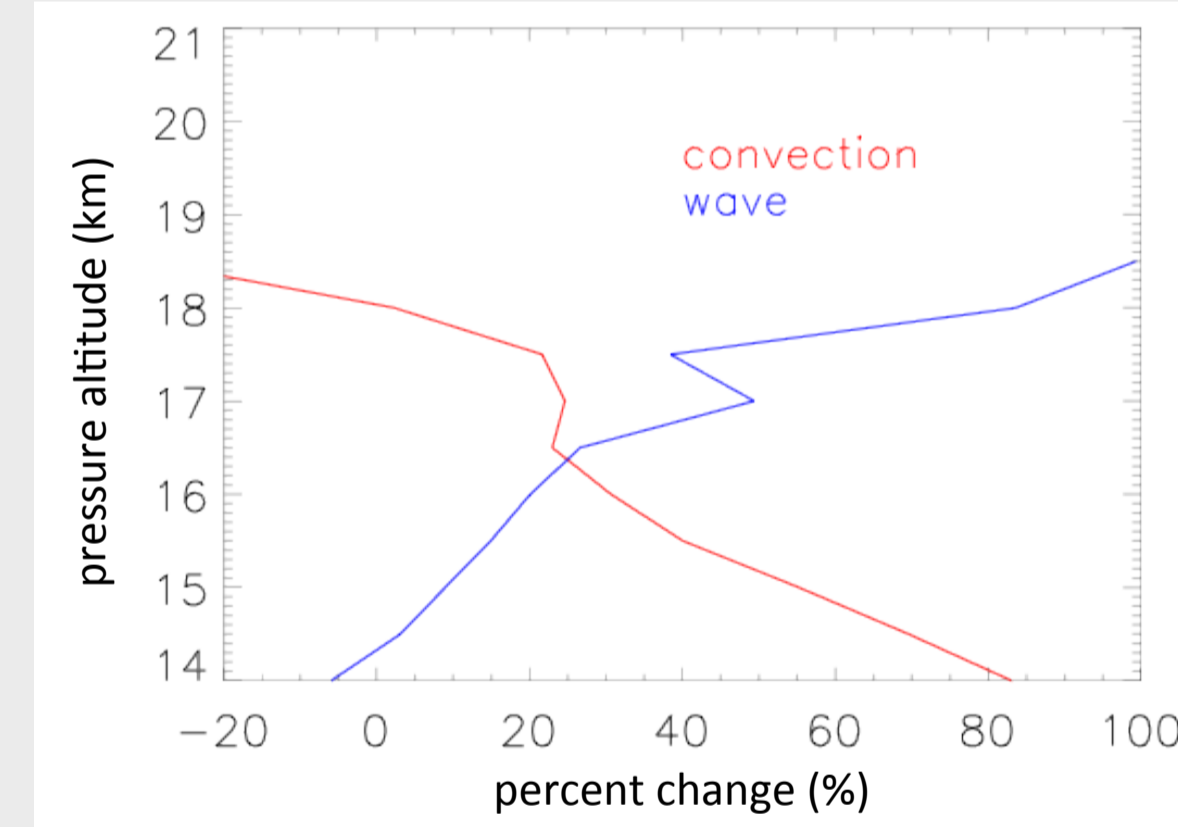


Fig. 9: Vertical profiles of the fractional change in the tropical (20°S-20°N) mean cloud occurrence frequency due to convection and waves.

- Cloud frequencies in the simulation with microphysics, convection, and waves agree well with those of CALIPSO at and above the cold point (~17 km): 8.4% vs. 9.2% at 17 km.

• Waves and convection increase cloud occurrence equally by ~25% near the 100 hPa (~16.5 km) level.

• Cloud occurrence at and above the cold point is mainly wave driven, while clouds in the lower TTL are primarily in situ clouds that form downstream of convection.

Sensitivity Tests

- Waves dehydrate (by 0.46 ppmv), whereas convection and microphysics moisten (by 0.56 and 0.71 ppmv, respectively)
- Waves and convection both increase cloud occurrence in the mid to upper TTL.

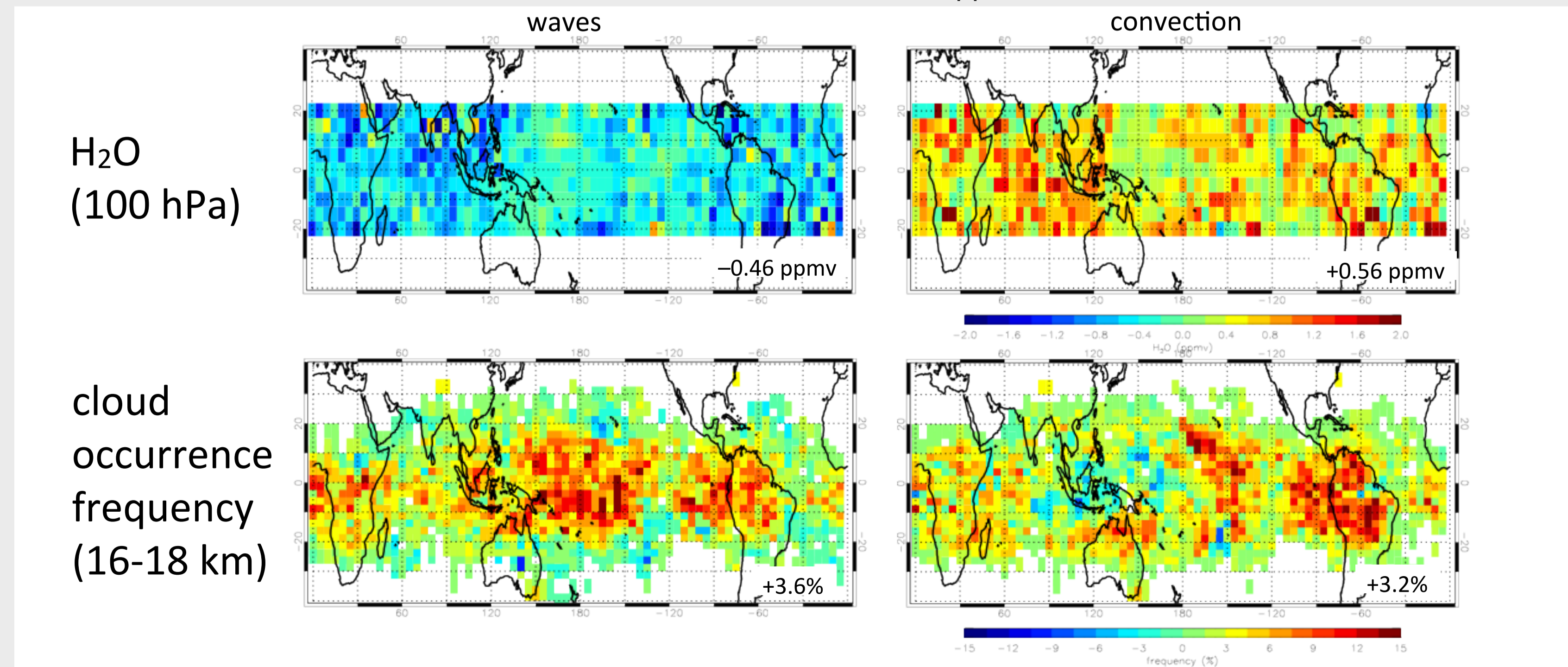


Fig. 10: The impacts of (left) waves, (middle) convection, and (right) microphysics on the (top) 100 hPa H₂O field and (bottom) cloud occurrence frequency in the mid to upper TTL.