Role of saturation in the water vapor diurnal cycle in the South American Tropical Tropopause Layer CT3LS 2015



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Motivation

How does overshooting convection affect the water vapor variability in the tropical tropopause layer and lower stratosphere?

TRO-pico project

Coordinator E. Rivière Co-coordinator J.-P. Pommereau www.univ-reims.fr/TRO-pico

Involves field campaigns in Bauru (southern Brazil), ground-based, space and balloon-borne observations, and modelling.



Content

- 1. Baseline and global scale overview
- 2. Regional scale case study: South America
 - Impact of overshooting convection
 - Role of the saturation in H₂O diurnal cycle
- 3. Conclusion

Section 1

Baseline studies & Overview of large scale impact of overshooting convection on the TTL/LS H₂O

Baseline study 1: Liu and Zipser, JGR, 2005



The Overshooting Precipitation Features (OPF) diurnal cycle above land, which peaks at 1600 LT with large amplitude, is different from that above ocean, which peaks at 0400 LT with small amplitude.

Baseline study 2: Liu and Zipser, JGR, 2009



The UT Microwave Limb Sounder (MLS) H_2O daytime sampling (1330 LT) is on average smaller than the night-time sampling (0130 LT) over land (D-N<0), and the opposite over ocean (D-N>0).

Baseline study 3: Carminati et al., ACP, 2014

MLS relative D-N H₂O (%) 2005-2012 solstitial seasons



 \rightarrow UT **negative** D-N in continental most convective regions. \rightarrow TTL/LS **positive** D-N in continental most convective regions.

Section 2

Regional scale case study: South America Impact of overshooting convection & Role of TTL saturation





D-N H₂O, temperature, RHi, and IWC ± 16

 Daytime moistening in Amazonia vs night-time moistening in Bauru at 100 hPa in convective season.



D-N H₂O, temperature, RHi, and IWC ± 16 days running mean

- Daytime moistening in Amazonia vs night-time moistening in Bauru at 100 hPa in convective season.
- Opposite sign of D-N in T and RHi linked to MLS sampling times in phase with the convective cooling in Amazonia but not in Bauru.



D-N H₂O, temperature, RHi, and IWC ± 16

- Daytime moistening in Amazonia vs night-time moistening in Bauru at 100 hPa in convective season.
- Opposite sign of D-N in T and RHi linked to MLS sampling times in phase with the convective cooling in Amazonia but not in Bauru.
- Daytime increase of IWC almost year-round near the equator, but only during the convective season in the tropics.



Differences in the TTL

Relative difference between Bauru and Amazonia of MLS 10-year (2004-2014) daily averaged H_2O , temperature, RHi, and IWC ± 16 days running mean.



 \rightarrow Bauru is warmer and has a lower RHi than Amazonia year-round.

- \rightarrow H₂O and IWC present a seasonal cycle of large amplitude.
- \rightarrow Wet and ice-charged summer TTL in Bauru is the sign of a convective moistening with ice injection/sublimation.

Differences in the TTL

Relative difference between Bauru and Amazonia of MLS 10-year (2004-2014) daily averaged H_2O , temperature, RHi, and IWC ± 16 days running mean.



But, it does not explain the difference in D-N between Bauru and Amazonia!

Hypothesis: The residence time of ice crystals



In a **super-saturated** TTL, ice crystals must sediment to subsaturated levels or perdure until the saturation is not longer sustained to sublimate, which should **require more time** than if they were injected in a sub-saturated TTL.

How to verify this hypothesis?

MOCAGE-VALENTINA assimilation tool

ightarrow mocage ctm

- ^O Horizontal resolution: global $(2^{\circ} \times 2^{\circ})$
- Vertical resolution: 47 levels from surface to 5 hPa in σ coordinates
- Dynamical/Met. forcing provided by ARPEGE analyses
- ^O Time step: 15 min for chem/phy processes & 60 min for dynamics
- O 140 hPa: Limit between ARPEGE constraints and chemical species

\rightarrow VALENTINA (PALM coupler)

- O 3D-Var method
- $^{\rm O}$ $\,$ Model error has been fixed to be 45% $\,$
- Assimilation with hourly frequency
- O Covered period: Nov 2011 to Mar 2012 (to be extended to Mar 2013)
- MLS water vapor is used from 316 to 5 hPa
- ^O Validated with MIPAS data sets

(Payra et al., AMTD, in prep.)

A 24-h resolved cycle for H_2O

MLS observations assimilated in MOCAGE-VALENTINA

Amazonia_{100hPa}

Daytime moistening, consistent with delayed sublimation of ice crystals in super-saturated TTL. Not seen in the CTM!

Bauru_{100hPa}

Night-time moistening, consistent with rapid sublimation of ice crystals in sub-saturated TTL.



Conclusion

At large scale

- Impact of convection in function of the season and of the surface.
- Dehydrating Cold Trap processes in oceanic regions.
- Hydrating overshooting injection of ice crystals in continental regions.

At regional scale in South America

- Overshooting injection of ice crystals both in tropical and equatorial regions, but the time scale for sublimation (and thus moistening) differs.
 - ⇒ In sub-saturated tropical TTL, rapid sublimation and late afternoon/nighttime moistening.
 - ⇒ In super-saturated equatorial TTL, delayed sublimation and late night-time/morning moistening (not seen in the CTM!).



Supplementary material

MLS characteristics

4 operational products relevant for this study:

- H_2O (190 GHz) (Read et al., 2007)
- Temperature (118 GHz) (Schwartz et al., 2008)
- RHi (from H₂O and Temperature products) (Read et al., 2007)
- Cloud Ice Water Content (IWC) (240 GHz) (Wu et al., 2008)



3 layers of interest 56 hPa \rightarrow LS 100 hPa \rightarrow LL 177 hPa \rightarrow UT 3 independent pieces of information Resolution (VxH) ~3x220 km Precision 6-25% Accuracy 4-20% (in 177-56 hPa pressure range)

Supplementary material

Assimilation

