Cold Pool Analysis from the Cold Pool Model Intercomparison Project (CP-MIP)

J. Kazil, U. Colorado/NOAA – R. Vogel, U. Hamburg – L. Bariteau, U. Colorado/NOAA – C. Bayley, MPI-M – P. Alinaghi, TU Delft – N. Antary, U. Hamburg – L. Bariteau, U. Colorado/NOAA – C. Bayley, MPI-M – P. Alinaghi, TU Delft – N. Antary, U. Leeds – A. Ekman, Stockholm Stockholm Stockholm U. – N. Falk, Colorado State U. – A. Fridlind, NASA GISS – S. Ghazaye, U. Cologne – T. Heus, Cleveland State U. – F. Hoffmann, LMU Munich – D. Mechem, U. Kansas – R. Neggers, U. Cologne – G. Raghunathan, Cleveland State U. – N. Robbins, MPI-M – J. Savre, LMU Munich – H. Schulz, Danish Meteorological Institute – S.-I. Shima, U. Hawaii – N. Tobias, U. Hawaii – S. van den Heever, Colorado State U. – T. Yamaguchi, U. Colorado/NOAA – T. Yanase, U. Hyogo – P. Zuidema, U. Miami

The Project

- CP-MIP identifies, characterizes, and quantifies:
- Differences between models in the representation of cold pools
- Biases relative to observations
- This poster shows simulations of a cold pool that forms underneath a "Flower" type shallow cumulus cloud

•The Flower cloud and its cold pool formed on February 2 2020, near Barbados, in the region sampled by the EUREC⁴A and ATOMIC field campaigns



Simulations

- Domain moves along an ERA5 500 m ASL wind speed trajectory
- •The trajectory runs through 12.75 °N, -58.5 °E on 2020-02-02, 17h00 UT • Domain: 144 × 144 × 7 km³
- Duration: 2020-02-01, 16h00m00s 2020-02-02, 18h00m00s UT
- dt = 2 s, dx = dy = 150 m, dz = 40 m 0-5 km, 19 geometric levels 5-7 km Periodic boundary conditions
- ⊢orcing:
- ERA5 geostrophic wind and subsidence
- ERA5 horizontal advective tendencies of temperature and moisture
- ERA5 sea surface temperature
- Interactive surface fluxes of sensible heat, latent heat, and momentum

Initialization

- Aerosol number: 400 mg⁻¹ in the boundary layer
- Cylindrical water vapor anomaly in the boundary layer, amplitude = 1 g kg⁻¹. • Mimics mesoscale moisture accumulation that would occur had the
- simulations been initialized 1–2 days earlier.
- Mesoscale moisture anomalies give rise to the formation of Flower-type trade cumulus clouds and thus to cold pools
- Advantages:
- The simulations need not to be initialized 1–2 days earlier,
- All simulations start out with the same mesoscale moisture anomaly



SAM (bin)

- System for Atmospheric Modeling (Khairoutdinov and Randall, 2003)
- 2-moment bin microphysics (Feingold et al., 1996, Tzivion et al., 1987)
- Prognostic aerosol/cloud/rain drop numbers

SAM (bulk)

- System for Atmospheric Modeling (Khairoutdinov and Randall, 2003)
- 2-moment bin-emulating bulk microphysics (Feingold et al., 1998; Wang and Feingold, 2009a, 2009b; Yamaguchi et al., 2017)
- Prognostic aerosol/cloud/rain drop numbers

MicroHH

- •GPU-enabled LES model (Van Heerwaarden et al., 2017)
- 2-moment bulk microphysics (Seifert and Beheng, 2006)
- Fixed cloud drop number

DALES

- Dutch Atmosphere Large Eddy Simulation (Heus et al., 2010)
- 2-moment bulk microphysics (Seifert and Beheng, 2006)
- Fixed cloud drop number

SCALE

- Scalable Computing for Advanced Library and Environment (SCALE) library LES (Sato et al., 2015)
- •1-moment bulk scheme (Tomita, 2008)
- Fixed cloud drop number



25 50 75 100 125

25 50 75 100 125

25 50 75 100 125

25 50 75 100 125

Evolution of Cold Pool Properties

- of rapid sub-cloud evaporation of rain in the bulk scheme.







Next Steps

- Compare
- Evaluate

• MicroHH and DALES, with the same cloud microphysics scheme, produce similar cloud water path, rain water path, and surface precipitation, as well as similar timing in these quantities. •SAM produces a nearly identical cloud water path with the bin and bulk microphysics. Yet, the bulk microphysics gives a high rain water path with low surface precipitation. This is the result

•SCALE's single-moment microphysics produces the lowest values of cloud water path, rain water path, but with still appreciable surface precipitation.

•The results separate into three groups based on the model cloud microphysics when considering the cold pool fraction and cold pool inversion height:

SAM (bin and bulk 2-moment microphysics) – with the caveat that the bulk microphysics produces the largest cold pool fraction towards the end of the simulation, as a result of its strong sub-cloud rain evaporation and associated surface layer cooling

MicroHH and DALES (2-moment bulk microphysics, Seifert and Beheng, 2006) SCALE (1-moment bulk microphysics, Tomita, 2008)



0.65 0.70 0.75 0.80 0.85 0.90 0.95 1.00 1.05 Time (d since simulation start)



•No obvious grouping appears when considering the surface layer temperature and water vapor anomalies. This is in part caused by a disagreement between the temperature and water vapor anomaly extent and cold pool extent defined based on the mixed layer height. Cold pools are defined based on a mixed layer height \leq 400 m following Touzé-Peiffer et al. (2022), denoted by black or magenta contours in the spatial plots (poster center panel).

^o probability density functions of 10 m temperature and water vapor anomalies, mixed layer height

Cold pool definition based on mixed layer height and on 10 m temperature anomaly Dynamical quantities, such as horizontal wind speed and divergence inside cold pools, and vertical velocity along the cold pool periphery

Cold pool dissipation, in terms of surface layer temperature and water vapor recovery Include results from the ICON model (Giorgetta et al., 2018)